# FOSSIL ASSEMBLAGES, STRATIGRAPHY, AND DEPOSITIONAL ENVIRONMENTS OF THE CROUSE LIMESTONE (LOWER PERMIAN) IN NORTH CENTRAL KANSAS

by

ROXIE LYNN VORAN

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KANSAS STATE UNIVERSITY Manhattan, Kansas

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Approved by:

Major Professor

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#### INTRODUCTION

# Purpose of Study

This study examines fossil assemblages, stratigraphy, and depositional environments of the Crouse Limestone in north central Kansas, focusing particularly on lateral and vertical changes. The Crouse Limestone was chosen because (1) a previous study (West, et al., 1972) had proposed a depositional model for the Crouse, (2) sections N and O (fig. 1) in the northern part of the area did not fit the proposed model, (3) important new information was available at section Q in the southwest part of the area, and (4) it enabled me to apply my interests in palaeobiology, carbonate petrology, stratigraphy, and field geology. Examination of sections N, O, and Q along with other outcrops in this area was undertaken to refine, expand, and test the proposed depositional model.

#### Location

Ten measured sections, six roadcuts and four stream exposures, were studied in detail in an area approximately 50 by 20 miles (fig. 1). Three additional sections (E, J, and N) were studied in less detail. Parts of Riley, Geary, Wabaunsee, Pottawatomie, and Marshall Counties make up the area which includes the areas studied by Huber (1965) and West, et al. (1972). In addition to outcrops, I was fortunate to have the opportunity to examine the Crouse Limestone in a core (fig. 1) taken in 1975 by Amoco Production Company.

# Previous Investigations

Huber (1965) summarized the nomenclatural history of the Crouse Limestone as follows:

The Crouse Limestone was named by Heald (1917, p. 21-22) from an exposure at Crouse Hill in the northwest part of the Foraker

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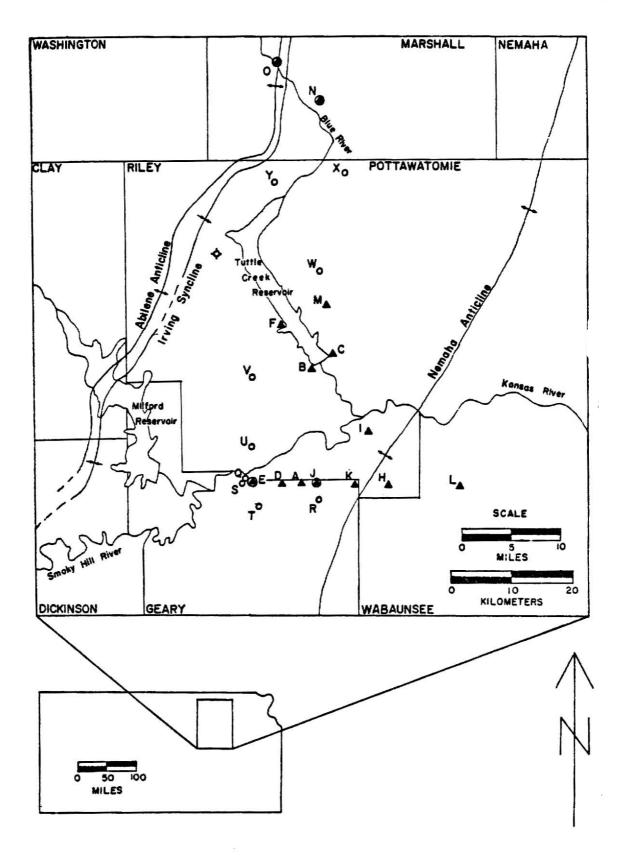


Figure 1. Location of Crouse Limestone Measured Sections.

quadrangle in Osage County, Oklahoma. The type locality is in the northeast quarter of section 23, T29N, R6E, about 2½ miles west of Frankfort, Oklahoma (Condra and Upp, 1931, p. 21). In 1927 Condra (p. 234) named this limestone the Sabetha Limestone and included it as one of three members in the Bigelow Limestone. The Bigelow consisted, in ascending order, of the Sabetha Limestone, the Blue Rapids Shale, and the Funston Limestone. The Bigelow Limestone was named for the town of Bigelow in Marshall County, Kansas. Later Condra and Upp (1931, p. 21) discarded the name Sabetha when they found that it was correlative with the Crouse Limestone.

In 1941 Jewett (p. 64-65) proposed dropping the Bigelow as a stratigraphic term and recognizing the Crouse Limestone, the Blue Rapids Shale, and the Funston Limestone as three separate formations.

Moore, et al. (1951, p. 46) described the Crouse Limestone in Kansas as follows:

The formation comprises an upper and lower limestone separated by a few feet of fossiliferous shale. The upper part displays platy structure and weathers tan to brown. The limestone beds locally are flinty. Thickness ranges from about 10 to 18 feet or more.

The three units described by Moore are recognizable in north central Kansas.

Throughout this report they will be referred to as "lower limestone,"

"middle shale," and "upper limestone."

Elias (1937) was among the first workers to deal with depositional environments of the Lower Permian rocks in Kansas. He believed that these rocks represented cyclic deposits and assigned relative distances from shore and absolute water depths to the different biotic and lithologic associations. He suggested that the Crouse Limestone was formed in a sea that was transgressive and later regressive with water depths ranging from 60 to 110 feet.

Mudge and Burton (1959) described the Crouse Limestone in Wabaunsee County, Kansas, adjacent to the study area on the southeast. They recognized three units and noted the abundance of fossils, some of them marine, in the lower limestone.

Huber (1965) studied six Crouse Limestone outcrops in north central Kansas (sections A-F, fig. 1). He believed the lower limestone was deposited in shallow, quiet water with low terrigenous influx and low turbulence. The middle shale and upper limestone were interpreted as deposits of a shallow, low-energy, brackish water environment. Such an interpretation implies that the Crouse is a regressive sequence.

Huber (1965, p. 24-25) described four stratigraphic units in the Crouse Limestone: (1) a lower limestone unit, (2) a middle "shaly" unit, (3) an upper "platy" unit, and (4) an upper "contorted" unit. In this report the upper "platy" unit and the upper "contorted" unit are considered together as the "upper limestone."

West, et al. (1972) examined 14 Crouse Limestone outcrops in the area (sections A, B, C, D, E, F, H, I, J, K, L, M, N, and O, fig. 1). They interpreted the Crouse as a regressive sequence and suggested that the lower limestone was formed in a shallow subtidal to low intertidal environment, the middle shale in a middle to high intertidal environment, and the upper limestone in a high intertidal to supratidal environment. They further interpreted the Crouse Limestone in this area as follows (p. 89):

...representing the eastern side of an embayed portion of a shallow open marine shelf. This lagoon was bounded on the east and west by the Nemaha and Abilene Anticlines respectively, and occupies the Irving Syncline. During Crouse deposition this lagoon probably gradually decreased in size. Drying up took place first in the north and northeast and last in the south and southwest. Such conditions would record a regressive sequence in rocks deposited during Crouse time.

Section Q is a roadcut made subsequent to previous investigations. Preliminary study indicated that this section tended to support the depositional model. If drying up did indeed take place as suggested by West, et al., the existence of a swamp (represented by a black shale at section Q) in the southwest part of the area would not be unexpected.

A primary objective of this study was to examine carefully section Q and nearby outcrops to determine if this information supports, negates, or requires refinement of the model.

Marine fossils typical of deeper water were found in the lower limestone at sections N and O and, therefore, these sections did not support the proposed model. Another aim of this study was to use careful stratigraphic correlation and examination of these and other outcrops in the central and northern parts of the area in an attempt to resolve this discrepancy.

#### METHODS OF INVESTIGATION

#### Field Procedure

Reconnaissance. -- Early in the investigation a reconnaissance was made to locate suitable outcrops. Elevations of previously studied Crouse Limestone outcrops were compared to United States Geological Survey topographic maps to determine areas of possible Crouse Limestone outcrops. Roads and streams in these areas were then checked. Eight exposures located in this way (sections R, S, T, U, V, W, X, and Y) were selected for study along with five previously known exposures (sections E, J, N, O, and Q).

Stratigraphic Description.--Each section was measured and described (Appendix I and West, et al., 1972, p. 46-49). Thicknesses were measured to the nearest one-half inch and later converted to metric units. Color (according to Goddard, et al., 1963), bedding, fracture, weathering characteristics, upper and lower contacts, fossils, and other features, such as bioturbation and sedimentary structures, were noted. Using these characteristics each measured section was divided into homogeneous units and

each unit was tentatively assigned a rock name.

Sampling. -- An effort was made to collect mappable bedding surfaces from the lower limestone. Such mapping would yield biotic data comparable to that collected by West and McMahon (1972) and facilitate lateral and vertical comparison of the lower limestone. To obtain the desired surfaces I collected slabs on which weathering had exposed and accentuated the fossils making them visible and identifiable. Such slabs were not always available, especially from outcrops where the base of the lower limestone was covered. When they were available they were seldom found in place and, therefore, could not be oriented with respect to compass direction.

Block and/or bag samples were collected, at least one sample from each described unit. More than one sample was collected from units in which some change in the fossil assemblage and/or lithology was noted.

Stratigraphic Correlation. -- Because study of lateral variation in the Crouse Limestone was an objective of this study, careful stratigraphic correlation was essential. The uppermost bed of the lower limestone is a distinctive marker in most parts of the area because it contains abundant gastropods, considerable iron oxide, and makes a prominent outcrop. The platy upper limestone is also distinctive. At sections N and O, however, the Crouse Limestone does not have these characteristics. This has led to erroneous stratigraphic correlation and considerable confusion regarding these two sections. At these localities a one- to three-foot thick limestone bed in the underlying Easly Creek Shale was a useful marker. This bed also makes a prominent outcrop and contains marine fossils, particularly ectoprocts and brachiopods, in the upper part; intraclasts occur locally in the lower part. This limestone bed was previously considered the lower Crouse Limestone, but examination of other outcrops, especially section X, helped

determine its proper stratigraphic position.

# Laboratory Procedure

General Statement. -- Each sample was examined using appropriate palaeontologic and petrologic techniques (fig. 2).

Washed Residue Examination. -- Argillaceous (bag) samples were disaggregated and examined for microfossil content. The procedure followed was the standard kerosene method described by Scott (1973, p. 9). This technique worked reasonably well on most samples, but some calcareous samples did not disaggregate well. To facilitate microscopic examination, disaggregated residues were sieved through 10 mesh (-1.0 phi), 18 mesh (0.0 phi), 35 mesh (1.0 phi), and 60 mesh (2.0 phi) sieves. This procedure was used to determine the types of microfossils (Appendix II) and the data were not quantified.

Bedding Surface Study.—Bedding surfaces from the lower limestone (fig. 3) were mapped using a technique similar to one described by Yarrow (1974). A one-eighth inch thick piece of plexiglas was laid on top of the surface and outlines of the fossils traced using grease pencils and rapidograph. Each map was transferred to a sheet of paper for permanence (Appendix III) and the plexiglas wiped clean with soft tissue paper and water. During mapping the following information was recorded for each fossil (Appendix III): (1) type of fossil (to genus if possible), (2) type of preservation, (3) direction of concavity (i.e. concave up or concave down), (4) articulated or disarticulated, (5) if disarticulated, which valve (i.e. right or left), (6) fragmentation, and (7) presence of attached epizoans. Length of Permophorus cf. P. subcuneatus shells was measured to the nearest millimeter. Size measurements were not made on other taxa because visual inspection of the surfaces indicated that whole specimens

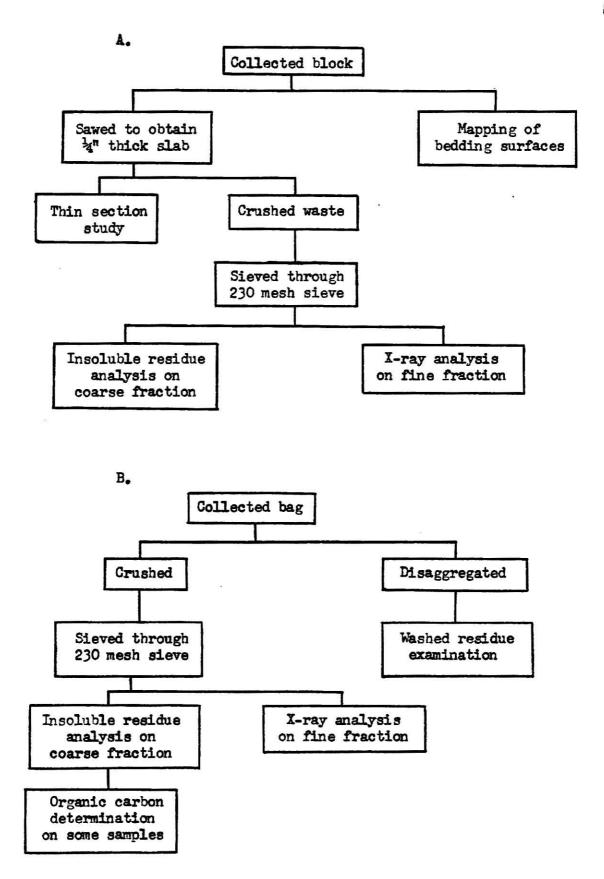


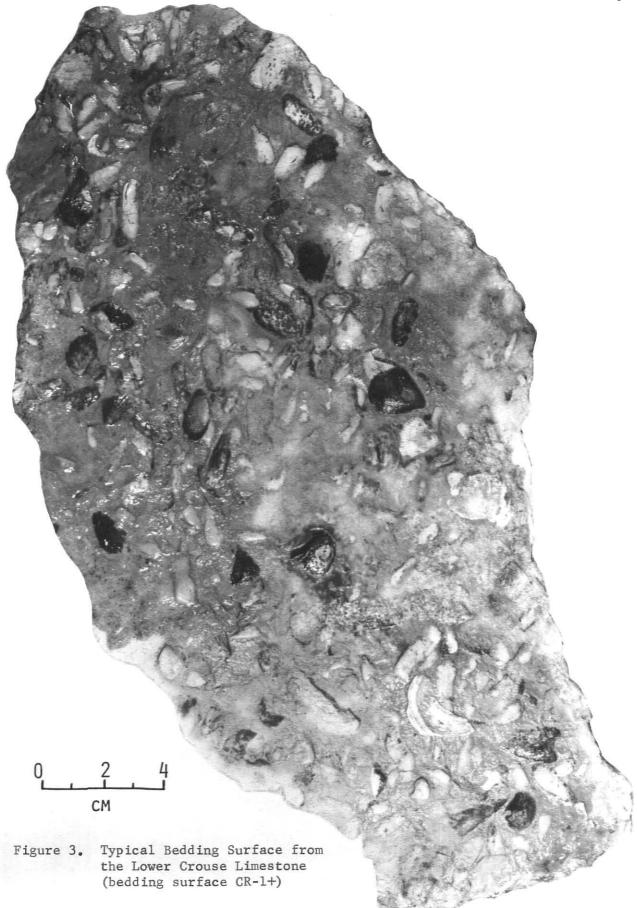
Figure 2. Flow Diagram of General Laboratory Procedure

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of other taxa were lacking on many surfaces.

Thin Section Study. -- Ninety-nine thin sections were made using a Hillquist thin section machine and Petropoxy 154 cement. All thin sections were examined with a petrographic microscope, briefly described, and any noteworthy features recorded. Following the method of Folk (1962), a rock name designating both grain size and composition was assigned to each thin section (e.g. medium calcarenite: immature recrystallized gastropod biomicrite) and is listed in Appendix IV.

Ten thin sections were selected to be point counted. Four were chosen to represent the different rock types and the other six to provide data for lateral and vertical comparison. Using the method described by Chayes (1949) approximately 1400 points were counted in an area 26 mm. by 19 mm. (1" x 3/4") on each slide (Appendix V).

<u>Insoluble Residue Analysis</u>.--Field observations showed that the entire formation was calcareous. Therefore, it was desirable to determine the weight percent insoluble residue.

Approximately one gram of crushed and dried sample was weighed to the nearest 0.0001 gram and placed in a beaker of known weight. Three percent hydrochloric acid was added slowly, allowing effervescence to diminish before adding more acid. After the reaction had ceased more acid was added and the sample was allowed to stand for several hours with periodic stirring. After washing and centrifuging (at least three times) to remove the acid, the residues were dried at 90°C. The residues and beakers were weighed and the weight percent insoluble residue calculated (Appendix VI).

The amount of organic carbon was determined for three samples of black shale and gray mudstone from section Q. The procedure followed is a modification of one described by Yarrow (1974). A known weight (between one

and two grams) of insoluble residue was placed in a beaker and covered with distilled water before 30 percent hydrogen peroxide was added. Hydrogen peroxide was added in 10 to 15 milliliter increments and the reaction allowed to diminish before adding more hydrogen peroxide. Beakers were placed in a drying oven at a temperature of 80°C to remove excess water and increase the reaction rate. When the samples were evaporated to dryness the process was repeated until no reaction occurred when hydrogen peroxide was added. The samples were then washed with distilled water until the pH was 6.5 to 7.0, oven dried, weighed, and percent weight loss calculated.

I-ray Diffraction Analysis. -- Samples were crushed and dry sieved through a 230 mesh (4.0 phi) sieve. The fine fraction was retained and used for X-ray diffraction analysis. Random powder mounts of each sample were prepared and run from 0 - 62 degrees two theta. The positions and heights of the principal peaks of quartz, calcite, and dolomite were recorded (Appendix VII).

# GEOLOGIC SETTING

#### Structure

The positions of the major structural features in the area (Nemaha Anticline, Abilene Anticline, and Irving Syncline) are in Figure 1 (Shenkel, 1959; Chelikowsky, 1972). All Crouse Limestone sections examined in this study except one (section 0) are on the eastern flank of the Irving Syncline with the Nemaha Anticline to the east and the Abilene Anticline to the west. Section 0 is on the Abilene Anticline. Three sections examined by West, et al. (1972) are either on the Nemaha Anticline (sections H and I) or east of its axis (section L). The Irving

Syncline is asymmetrical and plunges southward along with the Abilene Anticline. Neither feature can be traced south of Abilene.

# Stratigraphy

The Crouse Limestone is one of 14 formations in the Council Grove Group, Gearyan Stage, Lower Permian Series, Permian System (fig. 4).

It is above the Easly Creek Shale and is overlain by the Elue Rapids Shale. Both formations contain reddish shale representing very shallow, intermittently exposed or emergent environments (Elias, 1937; Russell, 1977). A six- to seven-foot thick bed of gypsum occurs in the Easly Creek Shale near section 0.

The Crouse Limestone in Kansas can be divided into three distinct units: an upper and lower limestone separated by a middle shale (Zeller, 1968, p. 47).

The lower limestone is "slabby" at the base to massive near the top. In northern and central Kansas molluscs dominate the fossil assemblage (West and McMahon, 1972), but in southern Kansas the alga Ottonosia is a conspicuous component (Bayne, 1962).

The middle shale in northern Kansas is really an argillaceous limestone because it contains more than fifty percent carbonate (Huber, 1965). In northern Kansas the fossil assemblage consists of ostracodes and foraminifers, but in central Kansas brachiopods, ectoprocts, and echinoderms occur (Mudge and Burton, 1959; O'Conner, 1953). In the southern part of the state the middle "shale" is a dense, locally cherty limestone (Bayne, 1962, p. 46).

The upper limestone is platy across the state. Algae and molluscs are in this unit in central Kansas (O'Conner, 1953). Previous workers in northern Kansas (Huber, 1965; West, et al., 1972) have reported

#### Speiser Shale Funston Limestone Blue Rapids Shale Crouse Limestone Easly Creek Shale Middleburg Ls. Mbr Hooser Shale Member Bader Limestone Eiss Limestone Member Stearns Shale Morrill Limestone Mbr. Council Grove Group Florena Shale Member Beattie Limestone Cottonwood Ls. Mbr. Eskridge Shale Neva Limestone Mbr. Salem Point Shale Mbr Grenola Limestone Burr Limestone Mbr. Legion Shale Member Sallyards Ls. Mbr. Roca Shale Howe Limestone Member Red Eagle Limestone Bennett Shate Member Glenrock Ls. Mbr. Johnson Shale Long Creek Ls. Mbr. Hughes Creek Sh. Mbr. Foraker Limestone Americus Ls. Mbr.

GEARYAN STAGE LOWER PERMIAN SERIES PERMIAN SYSTEM

Figure 4. Stratigraphic Section of the Council Grove Group

ostracodes, foraminifers, algae, and bivalves from the upper limestone.

In north central Kansas the Crouse Limestone averages 3.58 meters (11.8 feet) thick. It ranges from 2.22 meters (7.3 feet) at section N to 4.47 meters (14.7 feet) at sections S and T and, in general, thickens southward. The lower limestone averages 0.77 meters (2.5 feet) thick and ranges from 0.23 meters (0.8 feet) at section X to 1.21 meters (4.0 feet) at section S. The middle shale averages 0.92 meters (3.1 feet) thick. It is absent at sections H and I (West, et al., 1972), but ranges up to 1.71 meters (5.6 feet) at section S. The upper limestone averages 1.77 meters (5.8 feet) thick and ranges from 0.52 meters (1.7 feet) at section 0 to 2.37 meters (7.8 feet) at section C.

#### BIOTIC DATA

#### General Statement

Palaeoecologists are constantly concerned with fossil assemblages which may not accurately reflect life assemblages. In addition to the effects of transportation there is the problem of preservability. According to Johnson (1964) less than one organism in three that inhabits the benthic marine environment is preservable. Therefore, a fossil assemblage usually represents no more than one-third of the life assemblage. Often the information is even less representative because of post-mortem effects of physical and biological agents not to mention collecting biases.

The accuracy with which fossil assemblages reflect life assemblages has been discussed considerably. Most investigators agree that fossil assemblages are truly representative of the preservable parts of life assemblages. However, in terms of trophic structure fossil assemblages are not necessarily representative of life assemblages (Stanton, 1976) and

meaningful biological interpretation of relative abundance data cannot be made (Johnson, 1965). Consequently, fossil assemblages can be used only for palaeoenvironmental interpretations and, in fact, Warme, et al. (1976) found death assemblages more useful than life assemblages because of the removal of "noise." Care must be taken, however, that such interpretations are not based on the trophic structure of fossil assemblages or relative abundance data, but rather on information such as distribution, sediment relationships, and interspecific associations (Johnson, 1965; Stanton, 1976). In this study distribution and environmental tolerances are emphasized.

The mode of formation of fossil assemblages can be evaluated with criteria proposed by Johnson (1960). These criteria were applied to the bedding surface assemblages of the lower Crouse Limestone. The following evidence indicates similarity of the assemblages to Johnson's Model II: (1) ecologically coherent assemblage, (2) variation in sizes of fossils (from ostracodes less than 1 mm. long to bivalves more than 30 mm. long), (3) varying densities of fossils, (4) moderate proportion of fragmented fossils (average of 20.3 percent for all bedding surfaces), (5) orientation of the long axes of most fossils parallel to bedding, (6) nearly equal proportions of right and left valves (average of 49.3 percent right valves for all bedding surfaces), (7) algal coated shells (average of 8.8 percent for all bedding surfaces), and (8) sedimentary rocks consistent with tolerances of organisms and quiet water (carbonate mudstone). The low proportion of articulated fossils (average of 1.1 percent for all bedding surfaces) and hydrodynamically stable orientation of most specimens (average of 86.5 percent concave down for all bedding surfaces) favor Johnson's Model III. This evidence implies that some post-mortem movement of shells probably took place, but transportation was not a major factor in the formation of the assemblages. Most organisms probably lived very close to the spot where they were buried.

According to Johnson (1960), fossils in either Model II or Model III assemblages may show a preferred orientation in response to a prevailing current direction. Nagle (1967) demonstrated that shells may be oriented either by currents or by waves and he distinguished between the two types of orientation patterns. Pearson (1972) studied shell orientation on some lower Crouse bedding surfaces and concluded that wave action was the dominant factor responsible for the observed orientation.

#### Washed Residue Examination

Although no quantitative data were obtained from the washed residues, they were examined and microfossils identified. The following were identified in samples from mudstone intervals in the lower limestone: hollinid, bairdiid, and healdiacean ostracodes, pyramidellid snails, holothurian sclerites, and <a href="Earlandia-like">Earlandia-like</a> foraminifers.

Fossils in samples from the middle shale include hollinid, bairdiid, and healdiacean ostracodes and the foraminifers Ammodiscus? sp. and ammovertellids. Holothurian sclerites were also observed and a single productacean brachiopod spine was found at section Y. In addition, West, et al. (1972, p. 86) reported the ostracode Geisina sp. and the foraminifers Bathysiphon sp., Globivalvulina sp., and Eggerella-like foraminifers from the middle shale. Huber (1965, p. 34) reported Tetrataxis sp. from the same interval.

No washed residues from the upper limestone were examined.

# Bedding Surface Study

According to Dennison and Hay (1967), in order to be 95 percent certain of detecting a species of a particular size that occupies two percent or

more of the sea floor, an area 160 times as large as the area of that species must be examined. This rule was followed as closely as possible while mapping Crouse Limestone bedding surfaces. The area of the largest individual on the surface was multiplied by 160 and an area approximately that large was mapped. The size of some slabs prevented strict adherence to this rule. Mapped areas ranged from 100 to 565.5 square centimeters. Data obtained from mapped bedding surfaces were used to calculate diversity and equitability and to analyze size distributions of Permophorus cf. P. subcuneatus. Unidentifiable specimens and fragments that were less than half of the original specimen were not included in density (number of individual) tabulations. For bivalved taxa the number of individuals was considered to be the same as the greater number of valves. For example, if there were ten right valves and six left valves, the number of individuals would be ten. Some specimens were broken or partly covered and could not be identified with respect to right or left valve. Some or all unidentifiable valves were assumed to match unmatched right or left valves. If there were more unidentifiable valves than unmatched right or left valves, the remaining unidentifiable valves were assumed to be 50 percent right valves and 50 percent left valves. Therefore, the number of remaining unidentifiable valves was divided by two and the quotient added to the number of matched pairs to obtain the total (minimum) number of individuals.

## Diversity

Diversity is often considered to be the number of species in a given area. In this study, however, diversity refers to the Brillouin diversity index as modified by Illoyd, et al. (1968) and is expressed as:

 $H = \frac{c}{N}(\log_{10}N! - \Sigma \log_{10}n_{1}!)$ 

where H is diversity, N is the total number of individuals with ni individuals

in the ith species, and c is a scale factor for conversion of logarithm bases (for this study c = 3.322). Another commonly used measurement of diversity is the Shannon diversity index. It was not used in this study because the samples are not truly random. Every fossil in the population did not have an equal chance of being included in the sample because some have been eroded and others are not exposed (i.e. collecting biases). The diversity that was measured is really the diversity that exists on a very local scale and is in Table 1.

# Equitability

Equitability is a measurement of species evenness. The distribution of species is compared to MacArthur's (1957) theoretical distribution which is about as equitable as can be expected in nature. The equitability measurement used (Illoyd and Ghelardi, 1964) is:

$$E = \frac{S^1}{S}$$

where E is equitability, s is the number of species observed, and s' is the number of "equitably distributed" species needed to achieve the given diversity according to MacArthur's model. Equitability values are in Table 1.

## Size Analysis of Permophorus

Length of <u>Permophorus</u> cf. <u>P. subcuneatus</u> was measured on bedding surfaces from the lower limestone (Table 2). A more harsh environment in which the organism expended more energy to survive and less to grow is one possible cause of smaller size. To determine whether there was any lateral or vertical change in the size of <u>P. cf. P. subcuneatus</u>, t-tests were used. All possible lateral comparisons between stratigraphically equivalent surfaces were made along with vertical comparisons between surfaces from the same outcrop (Table 3). Lateral comparisons reveal no pattern.

Table 1

Diversity and Equitability

of Bedding Surface Assemblages in the Lower Crouse Limestone

Bedding Aurface Number	Diversity	Equitability	
CQ-2	2.070	0.857	
CQ-3	1.746	0.833	
CQ-1	0.843	0.500	
CQ-5	0.614	0.500	
CR-1	1.461	0.667	
CR-1+	1.628	0.571	
CT-1	1.481	0.571	
*CU-13	0.790	0.333	
CW-1	1.12և	0.750	
*CX-1	1.744	0.714	
J-1-1	1.474	0.499	

<sup>\*</sup> combines areas #1 and #2 at these horizons; see Appendix III.

Table 2

Length of Permophorus cf. P. subcuneatus on Lower Limestone Bedding Surfaces

Surface	Number of Individuals	Mean (mm.)	Standard Deviation
CQ-2	45	12.91	4.47
CQ-3	7և	10.32	3.32
CQ-1:	27	10.44	2.83
CQ-5	18	7.89	1.68
CR-1	150	10.74	4.98
CR-1+	150	11.26	3.54
CT-1	287	11.33	3.30
*CU-13	126	8.72	2.74
CW-1	10	11.60	4.62
*CX-1	79	10.77	4.09
J-1-1	119	10.05	3.36

<sup>\*</sup> combines areas #1 and #2 at these horizons; see Appendix III.

Table 3

Statistical Comparison of <u>Permophorus</u> of <u>P. subcuneatus</u> (valve length) between Bedding Surfaces in the Lower Crouse Limestone

Surfaces	Compared	t-test	Significant	F-test	Significant
lateral co	mparisons				
CW-1	CX-1	0.542		1.27	
*#CQ-2	CT-1	2.278	X	1.83	X
*CQ-2	CR-1	2.781	X	1.24	
CT-1	#CR-1	1.306		2.27	X
#CU-13	CQ-5	1.786		2.67	X
#CR-1	J-1-1	1.353		2.19	X
*#CQ-2	J-1-1	3.899	X	1.76	X
*CT-1	J-1-1	3.510	X	1.04	
vertical c	omparisons				
*#CQ-2	CQ-3	3.366	x	1.56	x
CQ-3	CQ-14	0.180		1.78	
*#CQ-L	CQ-5	3.789	X	2.19	X
#CR-1	CR-1+	1.040		1.98	X

<sup>\*</sup> sample with larger sized individuals

Critical t and F values are dependent on sample size. All tests for significance were made at the 0.05 alpha level.

<sup>#</sup> sample with larger variance

<sup>\*\*</sup> Surface listed on the left is lower stratigraphically.

but vertical comparisons indicate an upward decrease in size through the lower limestone. Field observations suggested that this trend exists throughout the area.

F-tests were performed to determine whether the variance of <u>Permo-phorus</u> cf. <u>P. subcuneatus</u> length changed either laterally or vertically (Table 3). Some lateral comparisons between stratigraphically equivalent surfaces show significant differences, but no definite trend. Vertical comparisons, however, show a decrease of the variance upward through the lower limestone.

# Autecology

General Statement. -- Autocology is the study of single species, their interaction with and adaptation to their environment (Odum, 1971, p. 6). The purpose of this investigation was not to explore these relationships in detail. Nevertheless, a brief examination of the autocology of the different taxa aids in the interpretation of the fossil assemblages.

Whittaker (1975) listed five trophic levels that make up the links to most food chains: (1) producer, (2) herbivore, (3) first carnivore, (4) second carnivore, and (5) third carnivore. Most food chains are considerably more complex because some organisms occupy several trophic levels. Because energy is lost at each step to a higher trophic level, productivity of each level is much less than that of the level below it. The number of individuals also decreases upward through most food chains (Whittaker, 1975, p. 215).

In the benthic marine environment the relationship of an animal to the substrate (mode of life) is an important part of its autecology. The mode of life categories used here were defined by West (1977) and include (1) epifaunal, animals living on the substrate, (2) epifloral, plants living on the substrate, (3) infaunal, animals living in the substrate, (4) semi-infaunal, animals living with one half or more of their length buried in the substrate, (5) quasi-infaunal, simulating an infaunal habit like some productacean brachiopods, and (6) nektic, swimming.

Among the first workers to recognize the importance of benthic marine detritus eaters was Turpaeva (1957). Her pioneering methods of trophic analysis have been applied, with modifications, to ancient communities by several workers (e.g. Walker, 1972; Rhoads, et al., 1972; and Toomey, 1976).

Walker and Bambach (1974) summarized previous work on feeding in benthic invertebrates and proposed a classification of feeding types. They emphasized (1) the type of food resource, (2) the location of the food resource, (3) the organism's food acquisition mechanisms, and (4) resource selection by the organism. Table 4 lists feeding groups defined by Walker and Bambach that are applicable to this study.

The trophic level, mode of life, and feeding behavior of all organisms encountered on bedding surfaces, in washed residues, and/or in thin section are in Table 5.

Algae.--Phylloid algae and encrusting algae were found in the lower limestone. Recrystallization of the phylloid algae prevented more specific identification. Osagia encrusted many fossils, especially bivalves and phylloid algae. Osagia consists of encrusting algae intergrown with sedentary tubiform foraminifers, especially Hedraites (Henbest, 1963). Bluegreen algae are represented by laminar stromatolites in the upper limestone. All algae are epifloral and, of course, primary producers.

Foraminifera. -- The encrusting sedentary tubiform foraminifers could have been either high- or low-level suspension feeders depending on their place of attachment. Toomey (1976, p. 11) considered them high-level suspension feeders because he found them attached to phylloid algae which

Table 4

Definition of Feeding Groups
(from Walker and Bambach, 1974)

Feeding Group	Definition
Low-level suspension feeders	Remove food from suspension in the water mass less than 3 mm. above the bottom
High-level suspension feeders	Remove food from suspension in the water mass more than 3 mm. above the bottom
Deposit feeders	Remove food from the sediment either selectively or non-selectively
Browsers	Scrape plant material from environ- mental surfaces or chew or rasp large plants
Carnivores	Capture live prey

were in growth position. In the Crouse Limestone these foraminifers commonly encrust bivalve shells, but only after death of the bivalves. Many bivalve shells are encrusted on both the inside and outside. Furthermore, Permophorus cf. P. subcuneatus shells are either encrusted almost entirely or are free of attached epizoans. If they had been encrusted while living, only the part of the shell protruding above the sediment-water interface would have been encrusted. Instead, P. cf. P. subcuneatus and other bivalves are completely encrusted which supports coating as a post-mortem event. Therefore, these foraminifers were attached to the substrate and are considered low level suspension feeders.

Free-living foraminifers are considered epifaunal suspension feeders and primary consumers (West, 1970; Toomey, 1976).

<u>Ectoprocta</u>.--A single ramose ectoproct colony was observed in thin section (sample CQ-h). Ectoproct colonies were epifaunally attached and

Table 5
Autecology of Fossils

Trophic Level	Mode of Life	Feeding	Type
	Ep-Epifloral Ea-Epifaunal I-Infaunal S-Semi-infaunal Q-Quasi-infaunal N-Nektic		vel suspension evel suspension feeder
TAXON	TROPHIC LEVEL	MODE OF LIFE	FEEDING TYPE
Algae		_	
Phylloid Algae	PR	Ep Ep	Ą
0sagia	PR	Ep	A
Hine-green Stromatolitic	Algae PR	Ep	A
Protozoa Foraminifera			
Tubiform encrusting forms	P	Ea	LS
Free-living mobile forms	P	Ea	HS/LS
Ectoprocta			
Ramose type	P	Ea	HS
Brachiopoda			
Inarticulata			
Lingulid	P	I	LS
Articulata			
Productacean	P	Q	LS
Mollusca			
Bivalvia			
Permophorus cf. P. subcune	eatus P	S	HS
Aviculopecten sp.	P	Ea	HS
Acanthopecten sp.	P	Ea	HS
Septimyalina sp.	P	Ea	HS
Pseudomonotis sp.	eatus P P P P P	Ea	LS
Edmondia? sp.	P	I	LS
Gastropoda Pollonophon sp	P	Ea	B/D
Bellerophon sp. Pyramidellids	P	Ea	B/D
97			
Armelida Polychaetia			

Table 5 cont.

TAXON	TROPHIC LEVEL	MODE OF LIFE	FEEDING TYPE
Arthropoda			
Cirripedia		í	
Acrothoracicans	P	Ĭ	LS
Ostracodes	P	Ea/I/N	B/D
chinodermata			
Crincid debris	P	Ea.	HS
Echinoid debris	S	Ea.	C
Holothurian debris	P	I/Ea	B/D
/ertebrata			
Fish debris	S/T	N	C

could not pursue food except to create water movements. Turpaeva (1957) classified ectoprocts as "awaiters," but because they use the same food resources as high level suspension feeders, subsequent workers have included them in that group. The trophic level of ectoprocts is primary consumer.

Brachiopoda. -- Brachiopods are represented by two fragmented specimens and are included only for completeness and do not represent a significant part of the assemblage. One is a lingulid from the lower limestone at section Q and the other is a productacean spine from the middle shale at section Y. Recent lingulids are infaunal suspension feeders. Their ancient counterparts are usually considered to have had the same mode of life and feeding type (Walker, 1972; Yarrow, 1974). Productaceans had a quasi-infaunal mode of life with supporting spines allowing them to live in soft substrates. This mode of life afforded them protection with only the commissure protruding above the substrate for food gathering (Rudwick, 1970).

According to Rudwick (1965) brachiopods feed chiefly on diatoms and dinoflagellates. McCammon (1969) has suggested that dissolved or colloidal organic nutrients may be an important food source for articulate brachiopods. Possibly brachiopods utilized both food sources. They are here considered primary consumers.

<u>Bivalvia</u>.--Bivalves are the most diverse group in the Crouse Limestone. They include infaunal, semi-infaunal, and epifaunal forms and both high- and low-level suspension feeders.

Aviculopectinids in the Crouse Limestone (Aviculopecten sp. and Acanthopecten sp.) possess an elongate anterior auricle and a byssal simus.

These features are indicative of an epifaunal, byssally attached mode of
life (Stanley, 1970, p. 32). Some living byssate pectinaceans are able to
release the byssus and swim, but they are not especially proficient swimmers
(Stanley, 1972, p. 193).

Permophorus cf. P. subcuneatus had a semi-infaunal mode of life (Voran, 1977; Stanley, 1972, p. 195). Because it protruded some distance above the bottom, it was probably a high level suspension feeder.

Pseudomonotis sp. is unique among Paleozoic bivalves because of its epifaunal cemented mode of life (Nicol, 1944; Mudge and Yochelson, 1962).

Newell (1960) and Newell and Boyd (1970) suggested that Pseudomonotis sp. may have given rise to the oysters. Because of its cemented mode of life, Pseudomonotis sp. is considered a low-level suspension feeder.

Septimyalina sp. is an epifaunal, byssally attached high-level suspension feeder and Edmondia? sp. is an infaunal low-level suspension feeder (Pearce, 1973, p. 21). The trophic level of bivalves is primary consumer (Scott, 1973).

Gastropoda. -- Gastropods were considered deposit feeders and/or browsers by West (1970, p. 83) and Toomey (1976, p. 7). Gastropods in the Crouse

Limestone may have obtained food by grazing on algae (primary consumers). West (1970, p. 100-101) inferred that gastropods are epifamal, however, there are infamal gastropods.

Polychaetia.—The polychaetes are represented by the genus Spirorbis which has been classified as a suspension feeder (Toomey, 1976). Idke encrusting foraminifers, Spirorbis sp. may be either a high- or low-level suspension feeder depending on its place of attachment. High organic carbon content of the rocks in which Spirorbis sp. occurs indicates a reducing environment of deposition. Such a reducing, anaerobic environment would not be favorable for habitation of the bottom. Plant debris is also in these units and although there is no direct evidence, it is inferred that Spirorbis sp. was epiphytic and should be considered a high-level suspension feeder. Spirorbis sp. is a primary consumer (Scott, 1973).

Arthropoda. --Ostracodes are the most common arthropods in the Crouse Limestone. Their modes of life may be epifaunal, infaunal, or nektic and they are considered deposit feeders and/or browsers (West, 1970, p. 101; Toomey, 1976, p. 7). Ostracodes are primary consumers.

Acrothoracican barnacle borings were observed in bivalve shells on several bedding surfaces of the lower limestone (fig. 5). Because they live in borings, acrothoracicans are considered infaunal. They feed by sweeping food, mainly plankton and organic detritus, from the water in the mantle cavity, in the boring, and near the boring (Tomlinson, 1969). Therefore, they are low-level suspension feeders and their trophic level is primary consumer.

Echinodermata.--Echinoderms are represented by crinoid, echinoid, and holothurian debris. Crinoids are epifaunal, high-level suspension feeders (Scott, 1973, p. 46) and their trophic level is primary consumer. Echinoids are considered epifaunal carnivores (West, 1970, p. 84) and are secondary



Figure 5. Acrothoracican Barnacle
Borings in a large Osagiacoated Bivalve in the
Lower Crouse Limestone
(bedding surface CR-1+).

consumers. Holothurian remains included a few sclerites found in washed residues. Holothurians are either infaunal or epifaunal deposit feeders and/or browsers (West, 1970) and their trophic level is primary consumer.

Vertebrata. -- Fish bones, teeth, and scales were found at section Q.

Fish are nektic and tooth morphology indicates that they were carnivores, either secondary or tertiary consumers.

Trace fossils.—Burrows and/or trails were observed on several bedding surfaces of the lower limestone and on both vertical and horizontal surfaces in the upper limestone (fig. 6A-C). Vertical surfaces of samples from the middle shale and upper limestone displayed disrupted laminations, mottling, and a reworked, homogenized lithology (fig. 7A-C). These features are indicative of bioturbation (Hardie and Ginsburg, 1977; Gebelein, 1971).

Worm tubes were identified in the lower limestone by West, et al. (1972)

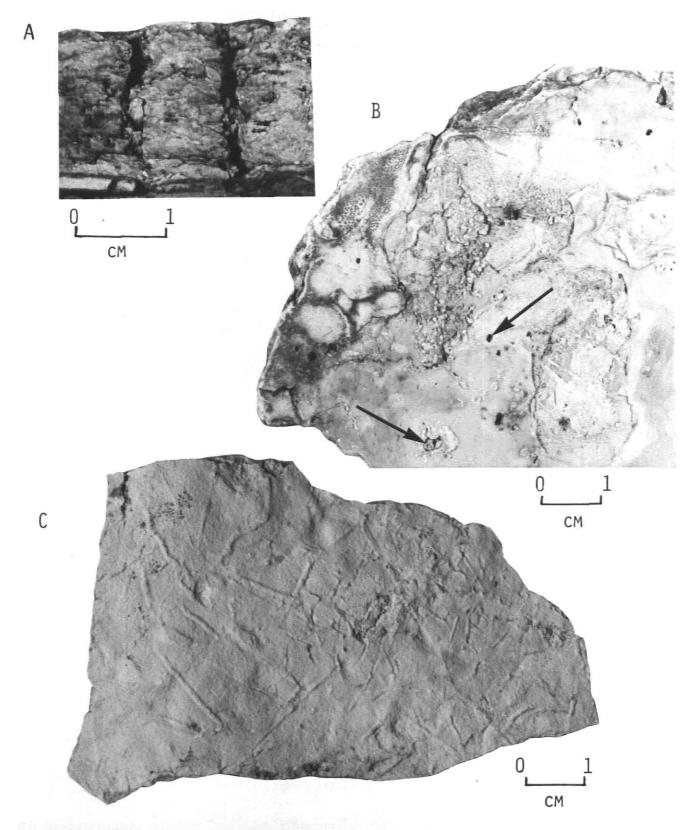
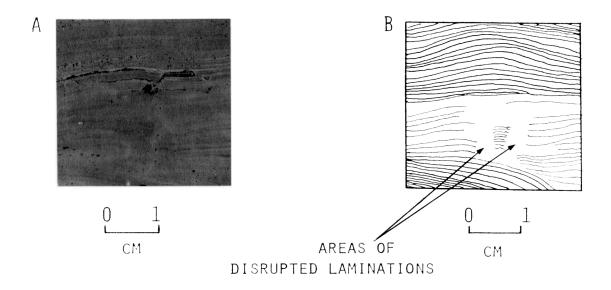


Figure 6. Burrows and Trails in the Upper Crouse Limestone. A. Vertical Burrows, Lateral View, Top is Up (sample CU-8). B. Same as 6A, Top View. C. Trails on Bedding Surface (sample CS-17).



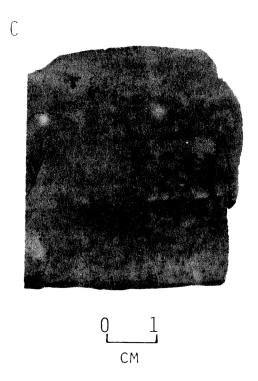


Figure 7. Evidence of Bioturbation in the Crouse Limestone. A and B. Disrupted Laminations, Top is Up (upper limestone, sample (CW-14). C. Mottled Lithology (upper limestone, sample CY-8).

and may represent the organisms responsible for at least part of the bioturbation. Trails (fig. 6C) could have been made by gastropods. Direct evidence of the organisms associated with the trace fossils is lacking.

#### PETROLOGIC DATA

#### General Statement

Insoluble residue, X-ray diffraction, and thin section analyses were made to assist in determining the depositional environment of the Crouse Limestone.

# Insoluble Residue Analysis

Insoluble residue analysis provided an estimate of the amount of terrigenous detritus. Insoluble residues may be introduced in three ways:

(1) as original constituents, (2) diagenetically, by metasomatic fluids, or (3) as a result of weathering. Samples from fresh surfaces were used to minimize the effect of weathering. Iron oxide replacing gastropods was observed in some lower limestone samples, but made up only a small part of the rocks. Therefore, it was assumed that the amount of insoluble residue approximates the amount that was deposited as an original constituent.

The average percents of insoluble residue for the three units are in Table 6 along with the values obtained by Huber (1965) and Twiss and Lee (1972). Agreement is close and the small differences probably result from the use of different outcrops in each study.

These data indicate that the middle shale is a shale only in its weathering characteristics. Most samples from this unit contained between 25 and 40 percent insoluble residue. As Huber (1965, p. 33) pointed out, this unit is really a limestone.

Huber (1965, p. 6) noted that the middle shale grades into the upper

Table 6
Insoluble Residue Percentages
for Crouse Limestone Units

Unit	This Study	Huber (1965)	Twiss and Lee (1972)
Upper Limestone	14.3	12.8	15.1
Middle Shale	29.6	27.5	33•2
Lower Idmestone	9.8	6.2	8.1

limestone. Field observations, especially at the less weathered outcrops, confirmed Huber's observations. Examination of the core (fig. 1) also emphasized the arbitrary placement of this boundary. Figure 8 shows the vertical changes in percent of insoluble residue at section Q. There is a sharp increase from the lower limestone to the middle shale and a gradual decrease from the middle shale to the upper limestone. This pattern is typical of the Crouse Limestone at most localities and further confirms the gradational boundary between the middle shale and the upper limestone.

The insoluble residues were not examined to determine their mineralogic composition. However, Huber (1965) examined the silt and coarse fractions from section A. He determined that the major minerals were quartz,
chert, and limonite with lesser amounts of chalcedony, orthoclase, volcanic
glass, gypsum, hematite, muscovite, and microcline. Twiss and Lee (1972)
analyzed samples from sections A, H, and L for clay minerals and found
illite, chlorite, and interlayered chlorite and vermiculite.

# X-ray Diffraction Analysis

All samples from sections Q, S, and Y and selected samples from all

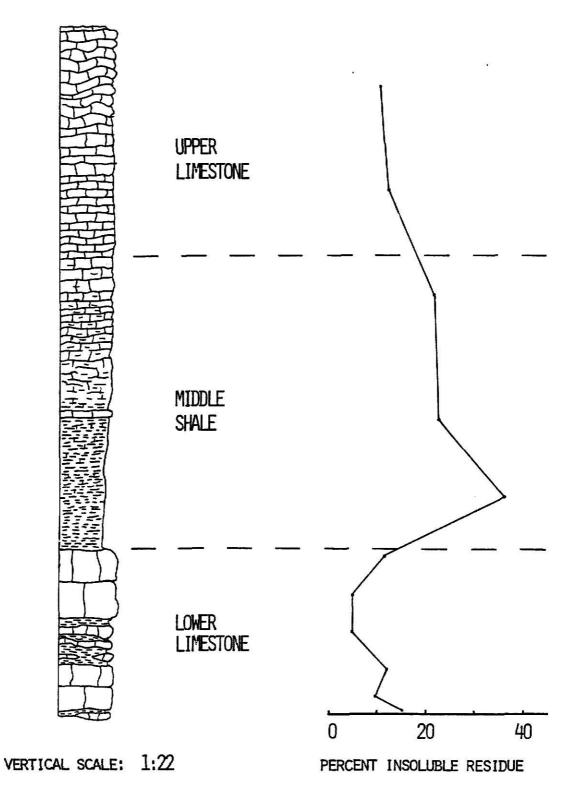


Figure 8. Insoluble Residue Percentages at Section Q

other sections were analyzed by random powder X-ray diffraction. The results indicated three major minerals: quartz, calcite, and dolomite. Calcite/quartz ratios and calcite/dolomite ratios were calculated using X-ray diffraction data (Appendix VII). This was done by comparing the peak heights of the principal peak for each mineral. This method is semi-quantitative at best, yet it provides an approximation of the relative abundances of these minerals.

Abundance of quartz can be roughly correlated with percent insoluble residue. Quartz content is low in the lower limestone, increases sharply in the middle shale, and decreases gradually through the upper limestone.

Dolomite had not previously been found in the Crouse Limestone and its occurrence may be described as erratic. Greatest amounts of dolomite are usually in the upper part of the middle shale. Dolomite occurs in both the middle shale and the upper limestone at sections Q, S, W, and Y. At sections O and E dolomite is only in the middle shale. Section T has dolomite in the upper limestone. No dolomite was found at sections R, U, v, and X.

Changes in dolomite composition may shift the position of the principal (2.88 %) dolomite peak. The most common variations in composition are the substitution of iron or calcium for magnesium in the dolomite structure. According to Miller (1967, p. 182), if large quantities of iron are present, copper radiation causes iron flourescence during X-ray diffraction. As a result the counts per second of the base line of the X-ray diffraction patterns increase. Thus, ferroan dolomite would cause an increase in the counts per second of the base line and greater amounts of ferroan dolomite would cause a greater increase in the counts per second of the base line. Calcium-enriched dolomite would not affect the base line. Scott (1973) compared the counts per second of the principal dolomite peak

with the counts per second of the base line and was able to show that ferroan dolomite was present in the Pennsylvanian Reading Limestone. This method was applied to the Crouse Limestone (fig. 7), but there appears to be no direct relationship between dolomite counts per second and base line counts per second. Therefore, it is assumed that there is no ferroan dolomite in the Crouse Limestone.

Absence of ferroan dolomite implies that any shift in the dolomite peak was due to calcium substitution in the dolomite structure. The mole percent calcium in dolomite can be determined from X-ray diffraction data. As calcium content increases, there is a linear increase in the d-spacing and a resultant shift in the position of the principal dolomite peak (Blatt, et al., 1972, p. 478). Mole percent calcium in dolomite from the Crouse Limestone averages 55.53 mole percent and ranges from 52.10 to 57.05 mole percent (Appendix VII).

Mole percent MgCO<sub>3</sub> in skeletal grains is related to three factors: mineralogy, taxonomy, and temperature (Bathurst, 1975, p. 235). Most calcite has either 0 - 5 mole percent MgCO<sub>3</sub> (low magnesian calcite) or 10 - 19 mole percent MgCO<sub>3</sub> (high magnesian calcite). High magnesian calcite is common in recent sediments, but has recrystallized to low magnesian calcite in most ancient limestones (Bathurst, 1975). MgCO<sub>3</sub> content of calcite can be determined from X-ray diffraction data. As MgCO<sub>3</sub> content increases there is a linear decrease in the d-spacing. This results in a shift in the position of the principal (3.03 %) calcite diffraction peak (Goldsmith, et al., 1955). This relationship was used to determine the mole percent MgCO<sub>3</sub> in calcite in samples from the Crouse Limestone (Appendix VII). Values ranged from 0.4 to 2.8 mole percent MgCO<sub>3</sub>-well within the range of low magnesian calcite. This may not have been true of the original sediment. Some skeletal fragments, especially some foraminifers, ostracodes, and

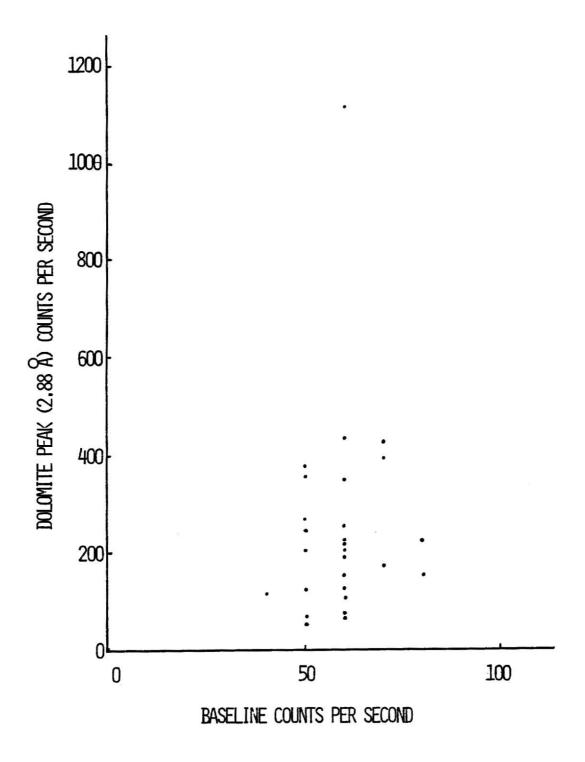


Figure 9. Diffractogram Baseline Counts Per Second Versus Dolomite CPS

calcareous algae are composed of high magnesian calcite. These taxa make up as much as 27.1 percent (sample CS-1, Appendix V) of the rock in the Crouse Limestone. Because of its relative instability, any high magnesian calcite in the Crouse was probably converted to low magnesian calcite during diagenesis.

# Thin Section Examination

A rock name was assigned to each thin section based on Folk's (1962) classification of limestones. Four major rock types were recognized: (1) recrystallized bivalve biomicrite, (2) recrystallized gastropod biomicrite, (3) recrystallized micrite, and (4) recrystallized laminated micrite. The lower limestone is a recrystallized biomicrite and belongs to Folk's Type II rocks (microcrystalline allochemical rocks). Most rocks from the middle shale and upper limestone are recrystallized fossiliferous micrites and belong to Folk's Type III rocks (microcrystalline rocks). Rhomb-shaped structures of dolomite and calcite along with rhombohedral pores indicate that many of these rocks were dolomitized and/or dedolomitized. Virtually all slides showed evidence of recrystallization of skeletal fragments and microspar and spar are commonly distributed patchily through the matrix (fig. 10). Percentages of orthochemical, allochemical, and terrigenous constituents were determined for the ten thin sections that were point counted (Table 7). Orthochemical constituents are divided into micrite (1 - 4 microns), microspar (4 - 10 microns), and spar (greater than 10 microns) according to Folk (1962). The lower limestone contains more allochems, more spar, and less micrite than the upper limestone. Percentages of microspar are nearly identical. The greater amount of spar in the lower limestone probably represents void fillings between allochems. Absence of voids in the upper limestone may be due to the absence of allochems.

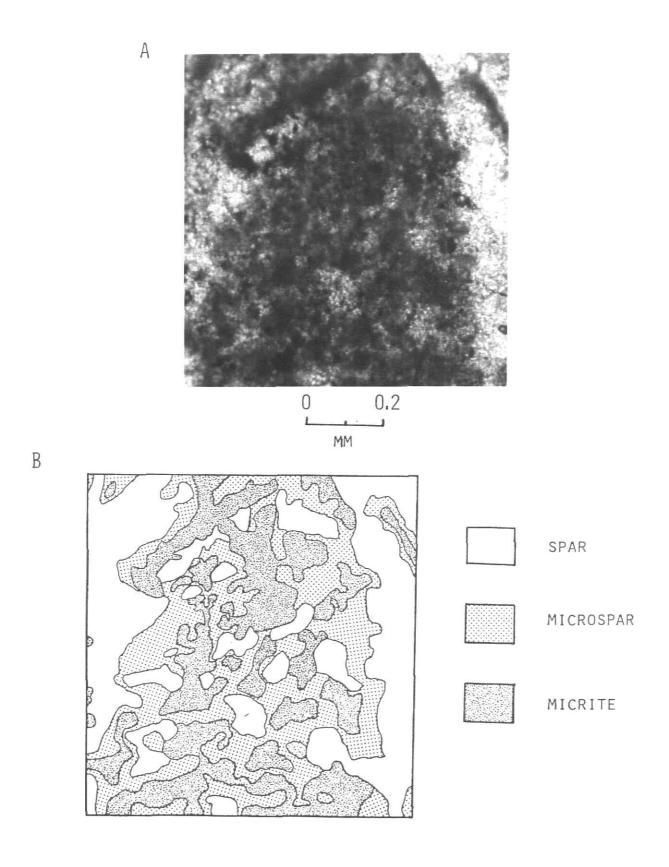


Figure 10. A and B. Patchy Distribution of Microspar and Spar in the Micrite Matrix of the Lower Limestone. Areas of Spar on the Left and Right are Recrystallized Skeletal Fragments (sample CR-2). Top is up.

Table 7

Percentages of Orthochemical, Allochemical, and Terrigenous Constituents in Thin Sections from the Crouse Idmestone

		Orthochemical Constituents		Fashandsof IA	Thermal recovers a
Sample	Morite	Microspar	æďS	Constituents	Constituents
Lower-ls. CO-1	62.14	8,9	1.8	26.5	0.3
CY-1a	28.3	27.0	6.8	36•4	<b>6</b> •3
GW=2	13.6	19.1	25.1	12.2	
*CV-1	38.6	3.7	13.0	0*1/1	0.1
CX-2	35.7	3.7	33.2	27.0	0.4
CS <u>−</u> 1	32.3	8.9	ን የ	52•3	†*0
*CS-6	16.7	11.0	23.9	7.444	1
Average	32•5	11.8	15,6	0*6€	0,2
Upper 1.s. *CS_15	85.3	6*6	1.1	6*0	1.0
*CS-16	79.1	15.5	3.1	1.1	6*0
0X-9	65.5	9.7	14.2	9.8	0.8
Average	9*91	11.7	1*9	ρ°7	6*0

\* the four major rock types

Similar amounts of microspar suggest that the effect of recrystallization was uniform throughout the formation.

Skeletal grains are the most abundant allochemical constituent, but some intraclasts and pelloids were recognized (slides CQ-5, CS-1, CT-5, CT-7, and CY-2; Appendix IV). The most abundant skeletal grains are ost-racodes, bivalves, Osagia, and gastropods (Table 8 and fig. 11).

Micrite, microspar, and spar were observed in each thin section. The patchy distribution of the microspar and some of the spar indicate that it is the result of recrystallization. An original micrite matrix suggests deposition in a low energy environment.

Silt-sized angular quartz is the most common terrigenous constituent observed in thin section (Appendix V). Small amounts of mica and clay minerals were also observed. These observations agree with those of Huber (1965) who found that quartz was the most abundant insoluble mineral in the sand and silt size fractions.

Enomb-shaped structures ranging from 30 to 150 microns were observed in most slides from the middle shale and the upper limestone (fig. 12).

Huber (1965, pl. XV) observed similar structures and concluded that they were not dolomite. In thin section some smaller rhombs appeared fuzzy making it difficult and sometimes impossible to determine their composition microscopically. X-ray diffraction indicated dolomite in 27 of 57 samples from the middle shale and upper limestone (Appendix VII), but in the remaining 30 samples calcite was the only carbonate mineral. Based on dedolomitization criteria (Evamy, 1967, fig. 1) these rhombs in the Crouse are interpreted as dedolomite. Three of the five stages of dedolomitization (Evamy, 1967) were observed in Crouse Idmestone thin sections—(1) rhombohedral pores, (2) rhombohedral pores partly filled with calcite, and (3) rhomb-shaped areas of equigranular mosaics of anhedral calcite (fig. 12).

# Allochemical Constituents of Thin Sections that were Point Counted

Type of Allochem	Number	Percent
Bivalves	913	22.3
Ostracodes	1197	29.2
Osagia	508	12.4
Gastropods	454	11.1
Encrusting foraminifers	137	3.3
Free-living foraminifers	118	2.9
Echinoderms	22	0.5
Phylloid algae	6	0.1
Unidentifiable skeletal fragment	s 672	16.4
Pelloids	15	0.4
Quartz silt	58	1.4
Fotal	4100	100.0

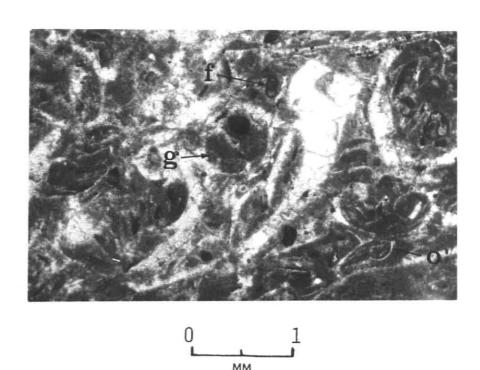
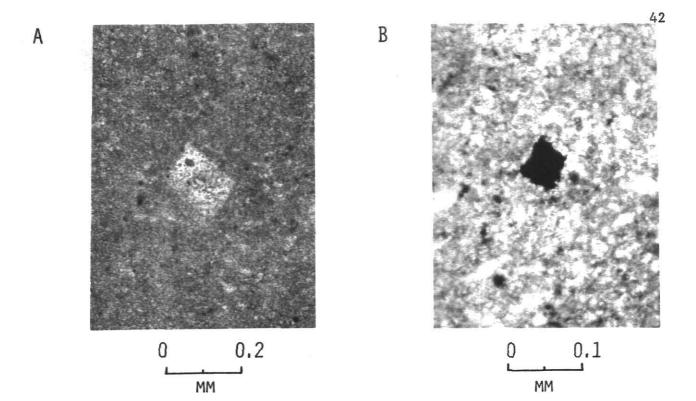


Figure 11. A Typical Thin Section from the Lower Idmestone (sample CS-1).

Large Areas of Spar are Bivalves. Arrows Indicate Foraminifer (f),
Gastropod (g), and Ostracode (o). Top is up.



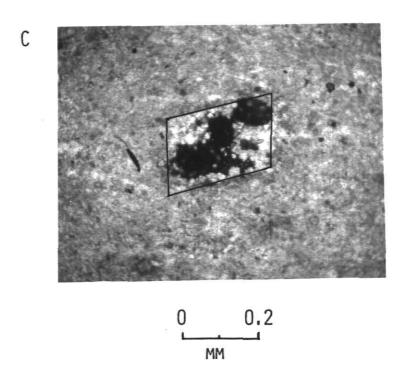


Figure 12. Dedolomitization in the Crouse Limestone. A. Rhomb-shaped Area Filled with an Equigranular Mosaic of Anhedral Calcite (middle shale, sample CV-7). B. Rhombohedral Pore, Crossed Nicols (upper limestone, sample CT-14). C. Rhombohedral Pore Partly Filled with Calcite, Crossed Nicols (upper limestone, sample CQ-11). Top is up on all photos.

Under proper conditions dedolomitization can proceed in the presence of sulfate ions which may be supplied by the dissolution of gypsum (Goldberg, 1967). In the Crouse Limestone gypsum crystal molds indicate that gypsum dissolution has taken place (fig. 13A), but the gypsum molds cannot be correlated stratigraphically with the supposed dedolomitization. Gypsum dissolution and dedolomitization in the Crouse may be related, but direct evidence is lacking.

Calcite also replaced gypsum in the Crouse Limestone. Replacement calcite forms areas of equigranular mosaics of anhedral calcite—a texture identical to that of some dedolomitization (fig. 13B). Some areas of gypsum-replacement calcite were rhomb-shaped in thin section and presented the problem of how to differentiate them from dedolomitization. The gypsum crystals and the pores resulting from their dissolution are elongate and when viewed in thin section they range from perfectly rhombic to elongate (fig. 13C). If in a given thin section replacement calcite was observed only as well-formed rhombs, it was considered dedolomite. If, however, there was considerable variation in the shapes of the areas of replacement calcite, it was assumed that calcite replaced gypsum.

#### DETAILED ANALYSIS OF SECTION Q

#### General Statement

As mentioned in the introduction, section Q is a unique exposure.

An erosional surface cuts through the upper limestone and part of the middle shale. Above this erosional surface are rocks unlike those exposed elsewhere in the Crouse Limestone. The rocks above the erosional surface at section Q will be considered in detail in this section.

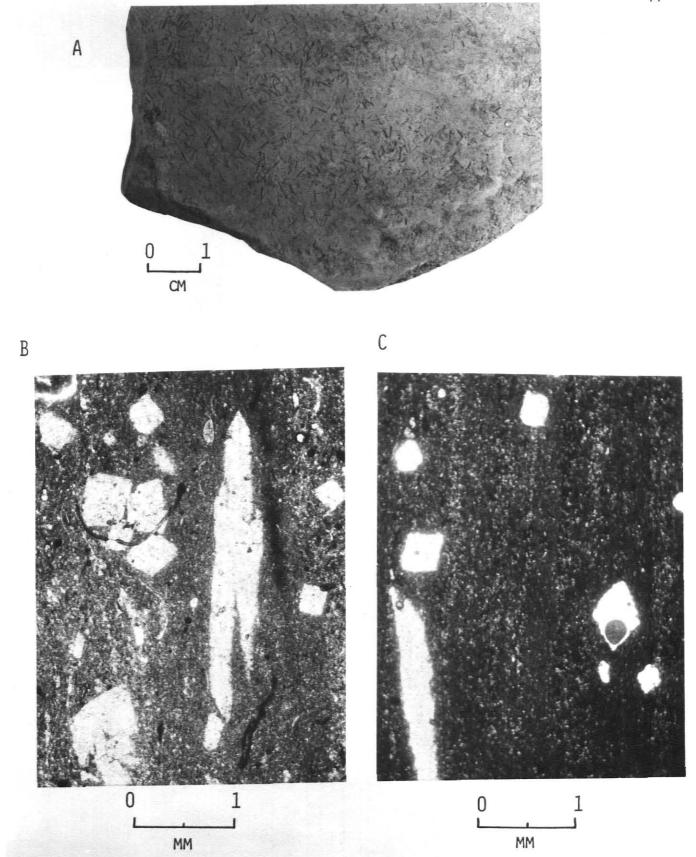


Figure 13. Evidence of Cypsum in the Crouse Limestone. A. Cypsum Crystal Molds from the Middle Shale (sample CW-8). B. Areas Filled with Equigranular Mosaics of Anhedral Calcite Interpreted as Replacement of Cypsum (upper limestone, sample CU-11). C. Pores from Cypsum Dissolution (upper limestone, sample CS-15). Left is up in B and C.

# Description

General Statement. -- The exposure at section Q and the four rock units between the erosional surface and the Elue Rapids Shale are in Figure 14. In ascending order these rock units are (1) a conglomerate, (2) a black shale, (3) a limestone, and (4) a gray mudstone.

Erosional Surface. The erosional surface cuts through the upper limestone and part of the middle shale at this locality. Excluding units above the erosional surface, the maximum thickness of the Crouse Limestone is 3.95 meters (13.0 feet) compared to 4.47 meters (14.7 feet) at section S less than a mile away. At its lowest point the erosional surface is 0.41 meters (1.4 feet) below the top of the middle shale. Assuming that the thicknesses of the middle shale and the upper limestone are the same at sections Q and S, between 0.52 and 1.96 meters (1.7 - 6.5 feet) of the Crouse was removed by erosion at section Q. The rocks immediately below the erosional surface have been altered from their typical color of yellowish gray (5Y7/2) to grayish orange (10YR7/4) or pale yellowish orange (10YR8/6).

To determine its configuration, a palaeotopographic map of the erosional surface was constructed (fig. 15). Elevations were determined (plane table and alidade) for 31 points (on both sides of the road) at the erosional surface. The data were adjusted to a datum 1.22 meters (3.9 feet) above the base of the middle shale to eliminate the effect of any post-Crouse deformation.

Conglomerate. -- Immediately above the erosional surface is up to 20 centimeters of conglomerate with a greater thickness where the erosional surface is lower (fig. 14). Figure 1 is an isopachous map of this unit made from data obtained while mapping the erosional surface. The conglomerate is composed of lithic clasts and a few pieces of charcoal. Where the unit is thickest the lower part is strongly cemented by sparry calcite, but

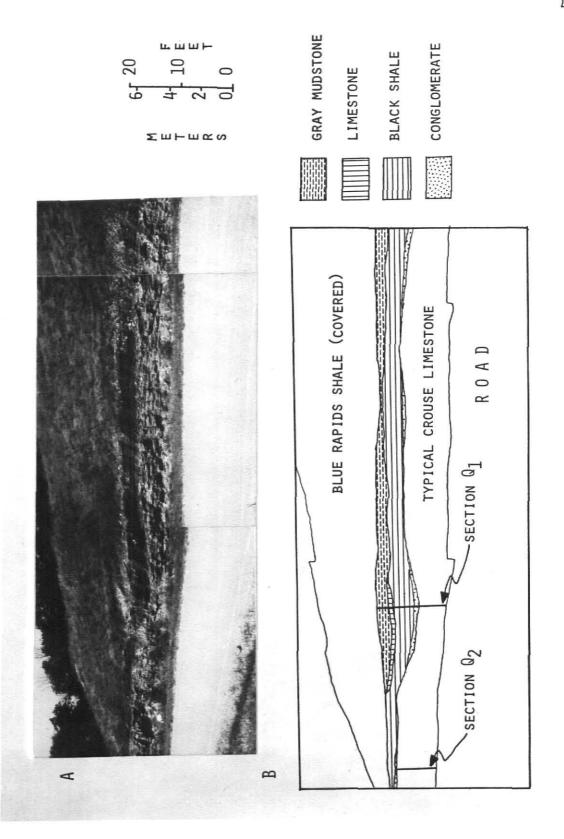


Figure 14. A and B. Crouse Idmestone at Section Q (north side of road).

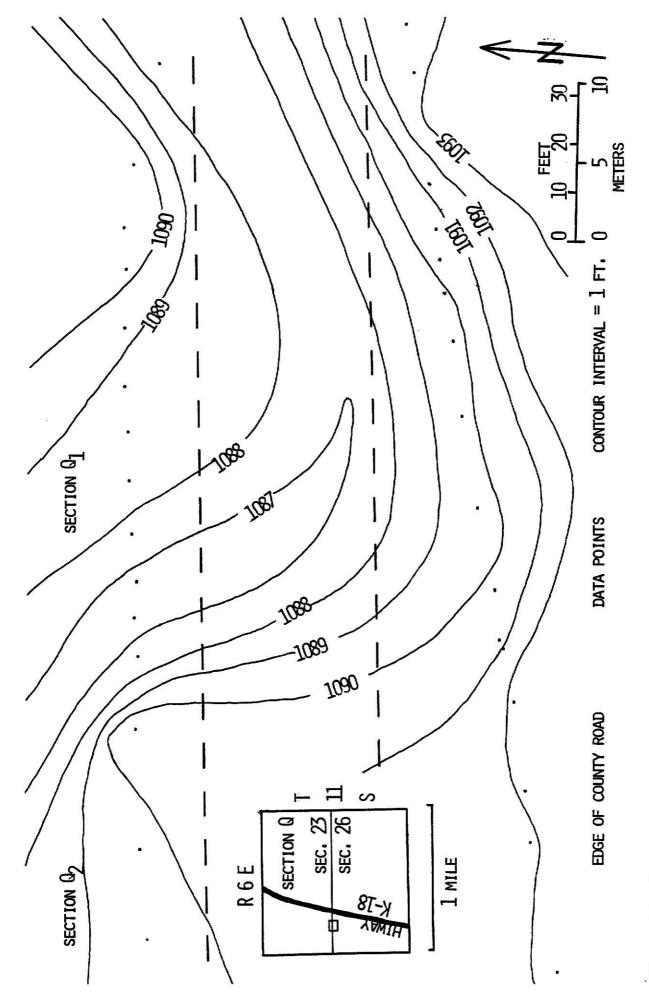


Figure 15. Palaeotopographic Map of the Erosional Surface at Section Q

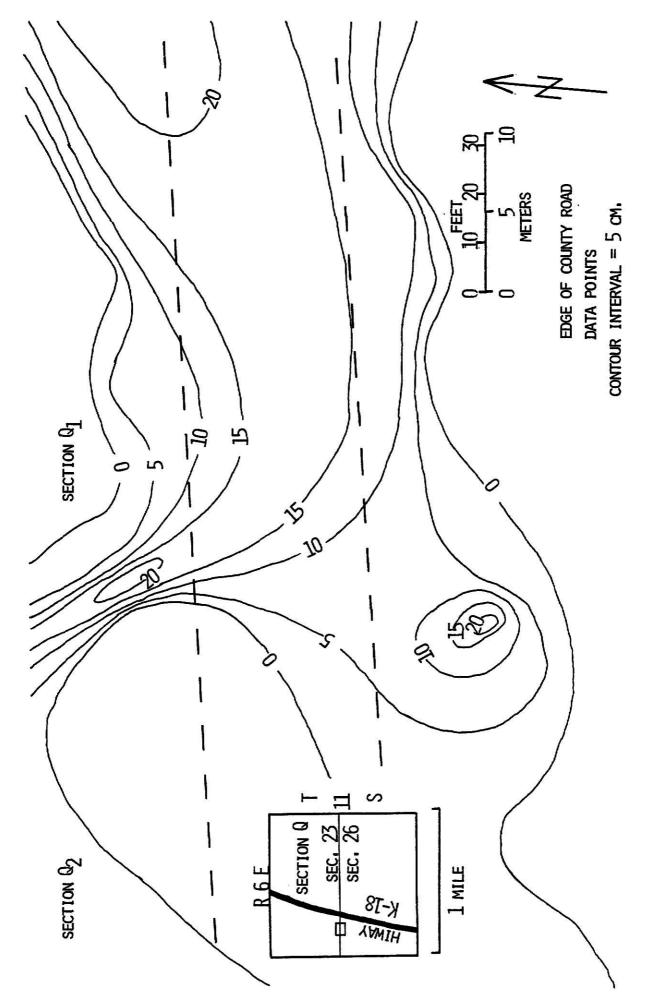
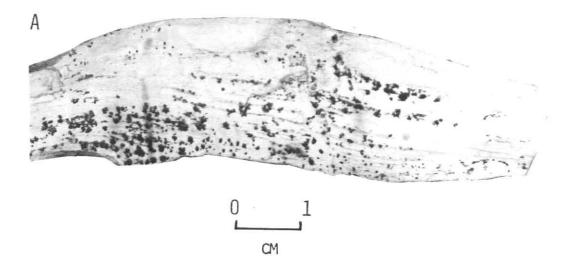


Figure 16. Isopachous Map of the Conglomerate at Section Q

elsewhere it is weakly cemented. The clasts range from 0.2 to 15.0 centimeters and are rounded to subrounded according to Fettijohn's (1975, p. 57) roundness classes. Visual inspection suggests that most of these thinly-laminated recrystallized micrite clasts are oriented with their long axes parallel to bedding. They are dark yellowish orange (10YR6/6) to grayish orange (10YR7/4) compared to the yellowish gray (5Y7/2) typical of the upper limestone. Lithologically the clasts and upper limestone are virtually identical. Based on these similarities it is reasonable to infer that the clasts were derived from this upper limestone (fig. 17). There are only a few examples of embayments of the clasts where they are in contact with each other indicating that most of them were lithified at the time of their deposition.

Hack Shale. -- Overlying the conglomerate is a black shale that ranges up to 1.37 meters (4.5 feet) thick and the greatest thickness is above the lowest point of the erosional surface (fig. 14). Although the unit is grayish brown (5YT3/2) near the base and the top compared to grayish black (N2) in the center, the contacts with the conglomerate below and the limestone above are sharp. Plant debris is abundant and there are several thin (less than five millimeters) discontinuous coal seams. Ostracodes are conspicuous and Spirorbis sp. and fish debris also occur.

Limestone.—Above the black shale is a discontinuous limestone (fig. 14) that is up to 15 centimeters thick and, like the underlying units, is thickest above the lowest point in the erosional surface. The limestone is subparallel to the erosional surface (fig. 14) and contains rounded clasts of micrite as large as six millimeters (fig. 18). These clasts are recrystallized laminated grayish orange (10YR7/4) micrite—a lithology identical to that of the upper limestone. The top of the unit contains more clasts but is less well cemented. Charcoal, fish debris, and ostracodes occur in this



В

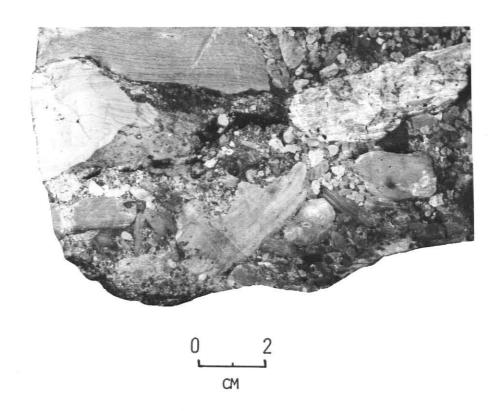


Figure 17. Similarity between Upper Crouse Limestone and Lithic Clasts in the Conglomerate at Section Q. A. Laminated Micrite from the Upper Limestone at Section Q (sample CQ-12). B. Conglomerate Composed of Laminated Micrite Clasts (sample CQ-13).

Top is up on both photos.

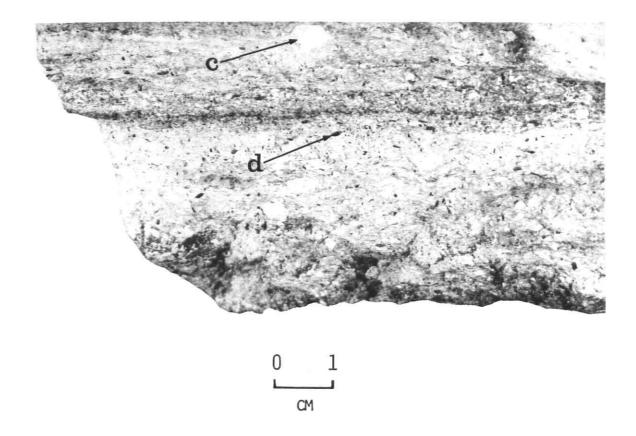


Figure 18. Limestone from above the Erosional Surface at Section Q (sample CQ-19) Arrows indicate Lithic Clast (c) and Fish Debris (d). Top is up.

unit (fig. 17).

Gray Mudstone. -- Overlying the limestone is a gray mudstone that grades upward into the Blue Rapids Shale. Again this gray mudstone is thickest above the lowest point in the erosional surface. It appears to fill low areas left after deposition of the underlying units. A few sand-sized grains of rounded micrite which are lithologically similar to the lithic clasts in the limestone and conglomerate units below occur in this mudstone. Fish debris and Spirorbis sp. were the only fossils found.

## Data

Biotic Data .-- Biotic data were obtained from these units by examining

washed residues and thin sections but were not quantified.

Except for a few pieces of charcoal, the conglomerate is devoid of any organic remains. The black shale contains abundant plant debris and smooth-shelled ostracodes along with <u>Spirorbis</u> sp. and some fish debris. Ostracodes could not be identified more specifically because nearly all were broken. The limestone contains abundant ostracodes as well as fish debris and charcoal and the gray mudstone contains <u>Spirorbis</u> sp. and fish debris.

Petrologic Data. -- Insoluble residue and X-ray diffraction analyses were made on samples from these four units. Organic carbon content was determined for the black shale and the gray mudstone.

Figure 19 shows the percent of insoluble residue, calcite/quartz ratio, and percent of organic carbon for each sample analyzed. The insoluble residues are low in the conglomerate and the limestone, and the calcite/quartz ratios are high. The mudrocks (shale and mudstone) have high insoluble residue content and low calcite/quartz ratios.

Organic carbon ranges from 16.2 to 18.1 percent of the black shale and the gray mudstone. Mudrocks may have up to 40 percent organic carbon, but less than three percent of all mudrocks contain more than ten percent organic carbon (Elatt, et al., 1972, p. 393). Therefore, the organic carbon content of the black shale and the gray mudstone is unusually high.

Based on thin section examination, the limestone (samples CQ-19 and CQ-19+) is a recrystallized ostracodal biomicrite or biomicrudite.

### Interpretation

Erosional Surface. -- The erosional surface cuts through the upper limestone and middle shale and, therefore, represents a post-Crouse event. More than two meters of relief is on this surface (fig. 15). In vast, shallow

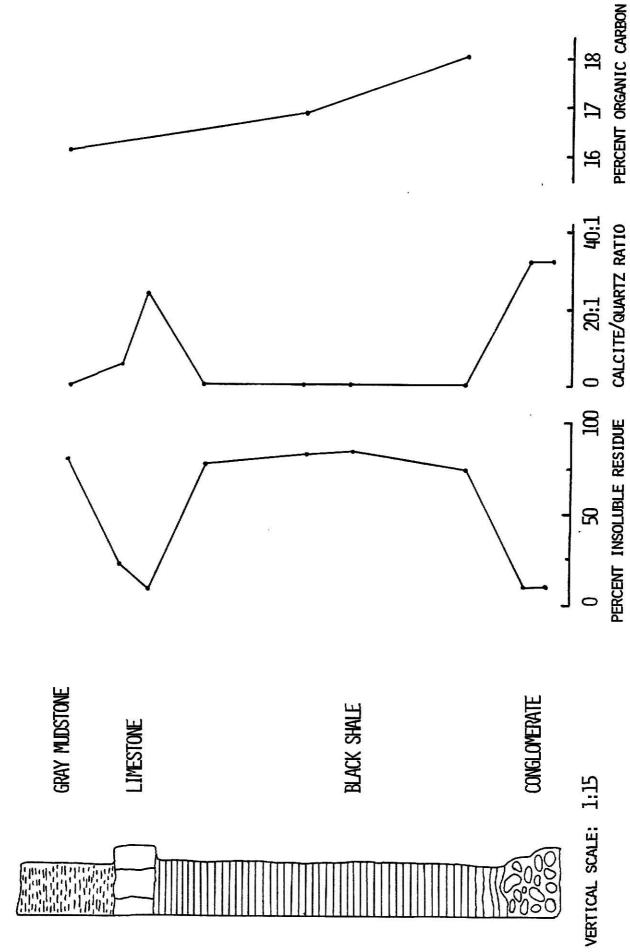


Figure 19. Petrologic Data from the Units Above the Erosional Surface at Section Q

epeiric seas with low energy, submarine erosion of this magnitude seems unlikely. The orange color of the rocks immediately above and below the erosional surface is probably a result of iron oxidation, suggesting that erosion took place in an oxidizing environment (<u>i.e.</u> subaerial).

The shape of the erosional surface (fig. 15) suggests that it may have been the bed of a channel that drained a subaerially exposed area. Its lateral extent could be due to meandering or lateral migration of a tidal channel. The dimensions of this channel (2 meters deep and 25 to 35 meters wide) are similar to those of modern tidal channels on Andros Island, Bahamas (Shinn, et al., 1969). Rapid downcutting during a period of abnormally high energy (i.e. storm) may have deepened the channel.

Conglomerate.—Sub-rounded to rounded lithic clasts with a wide range in size compose the conglomerate. The clasts are lithologically identical to the upper limestone and are believed to have been derived from its erosion. According to Pettijohn (1975, p. 515), limestone pebbles may become well rounded when transported 15 kilometers or less. Erosion of the Crouse Limestone took place soon after deposition of the upper limestone and probably before lithification was complete. Therefore, erosion and rounding of the clasts could have taken place even more easily and quickly. This coupled with poor sorting suggests a nearby source area.

The large size of some clasts and the absence of micrite suggest a high energy environment of deposition which supports the interpretation of the erosional surface as a channel bed. Laminated mud clasts and lithified dolomitic clasts are common in recent tidal channel deposits on Andros Island, Bahamas (Hardie and Ginsburg, 1977; Shinn, et al., 1969). The distribution of the conglomerate (fig. 16) is erratic and suggests deposition by an overloaded stream which supports the idea that maximum downcutting took place during a storm.

Black Shale. The black shale represents a change to a low energy environment and the fossils in it suggest brackish water. Spirorbis sp. is found in both marine and freshwater deposits (Scott, 1944). According to Lane (1964) smooth-shelled ostracodes occur in both marine and freshwater assemblages and fish debris is typical of a freshwater assemblage.

Organic carbon is an indication of a poorly oxygenated environment. In an oxidizing environment organic carbon is oxidized to carbon dioxide and water. High organic carbon content of the black shale along with coal and abundant plant debris indicate stagnant, anaerobic conditions. According to Zangerl and Richardson (1963) the absence of infaunal and benthic organisms also suggests a toxic bottom. The coal and plant debris indicate close proximity to a terrestrial environment and the characteristics of this black shale are typical of a swamp environment.

Reducing conditions at the bottom do not imply similar conditions in the water above. Spirorbis sp. and fish debris in the black shale suggest that parts of the water were well oxygenated and inhabitable. The stagnant, anaerobic conditions probably extended just a short distance above the bottom (Zangerl and Richardson, 1963).

Limestone. The limestone unit has a limited lateral extent (fig. 14). Ostracodes, fish debris, and charcoal suggest a freshwater or perhaps brackish environment, but other features common in freshwater limestones such as algal structures and nodular form are absent. The micrite matrix indicates deposition in a low energy environment, but the lithic clasts suggest somewhat higher energy than during deposition of the black shale. The shape of the unit (fig. 14) suggests deposition in a channel or may be due to differential compaction of the underlying shale. A channel deposit is more compatible with the lithic clasts and the unit's limited lateral extent.

Cray Mudstone .-- The gray mudstone appears to have filled low areas left

after deposition of the underlying units. It consists of over 80 percent insoluble residue and contains 16.2 percent organic carbon. Where the limestone unit is absent the contact between the gray mudstone and the underlying black shale is gradational. The upper contact with the Blue Rapids Shale is also gradational and the gray mudstone could, in fact, be considered part of the Blue Rapids. Spirorbis sp. and fish debris suggest a brackish to freshwater environment. Lane (1964) found a similar fossil assemblage in the Blue Rapids Shale. It included charophytes, which he believed represent either fresh or brackish water, the marine foraminifer Tetrataxis, the freshwater ostracode Carbonita, and fish debris. The high organic carbon content of the gray mudstone indicates reducing conditions of the bottom, unlike the oxidizing conditions during deposition of the reddish shales in the Elue Rapids Shale above. Although water conditions during deposition of the gray mudstone and the Blue Rapids Shale may have been similar, bottom conditions certainly were not. The slight decrease in percent of organic carbon through the black shale and gray mudstone (fig. 18) may be part of the change from reducing to oxidizing conditions.

Placement of the gray mudstone in the Crouse Limestone or the Blue Rapids Shale is not important. It is considered here as part of the depositional sequence and clearly represents a transition between the units above and below. This transitional relationship demonstrates the artificiality of certain stratigraphic boundaries and the need to study a rock unit as part of a depositional sequence.

#### Summary

A tidal charmel at section Q eroded parts of the upper limestone and middle shale. Rounded clasts from these units were deposited as gravel immediately above the erosional surface in a high energy environment.

Subsequently, a black, organic-rich mud was deposited in a swamp-like environment with low energy, fresh or brackish water, and reducing conditions at the bottom. A skeletal carbonate mud containing small (less than six millimeters) rounded lithic clasts was then deposited in a limited area. It formed in fresh or brackish water with more oxidizing conditions and somewhat higher energy than the mud below. Its shape and limited lateral extent suggest that this unit was deposited in a channel. The sequence was capped by a gray mud. It, too, was deposited in a fresh or brackish water, low-energy environment with reducing conditions on the bottom and represents a transition with the overlying Mue Rapids Shale.

The units above the erosional surface at section Q undoubtedly record a very local event. These units are unique in the Crouse Limestone and even nearby outcrops give no indication of a similar situation. The limestone and conglomerate units thin to a zero edge and the other units thin considerably toward the edge of the outcrop. The types of deposits also imply a restricted occurrence. In an area characterized by deposition in shallow epeiric seas, a high-energy environment like that represented by the conglomerate is not likely to be widespread.

# INTEGRATION AND INTERPRETATION OF DATA

## Vertical Integration and Interpretation

Rocks at each outcrop of Crouse Limestone were formed as a result of a sequence of events and conditions. To best understand these events and conditions it is necessary to study the entire sequence of rocks because each event and each set of conditions did not occur exclusive of those preceding and following it. Furthermore, many factors affect the formation of carbonate rocks. Although it is impossible to consider each factor, the

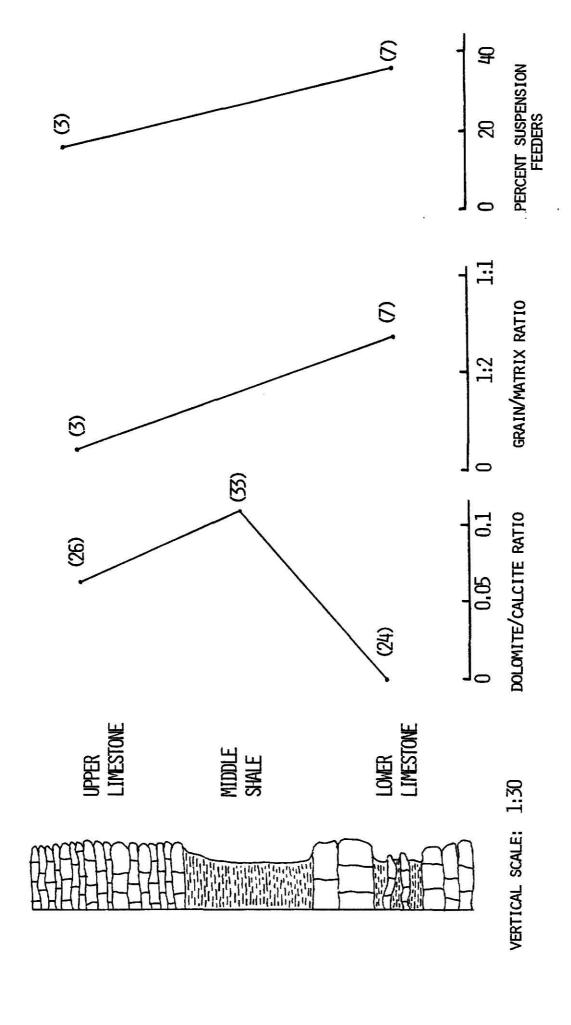
best interpretation is the one that takes into consideration the most evidence.

Figures 20 and 21 are the integration of petrologic and biotic data respectively for all Crouse Limestone sections. Petrologic data (from thin sections and X-ray diffraction) show the following trends (fig. 20): (1) decreasing percent of suspension feeders upward, (2) decreasing grain/matrix ratio upward, and (3) dolomite in the middle shale and upper limestone. Biotic data (from lower limestone bedding surfaces) indicate upward (fig. 21): (1) increase in fragmentation, (2) decrease in diversity, (3) decrease in equitability, (4) decrease in size of Permophorus cf. P. subcuneatus, and (5) increase in deviation of the percent right valves from 50 percent. These trends were not tested for statistical significance.

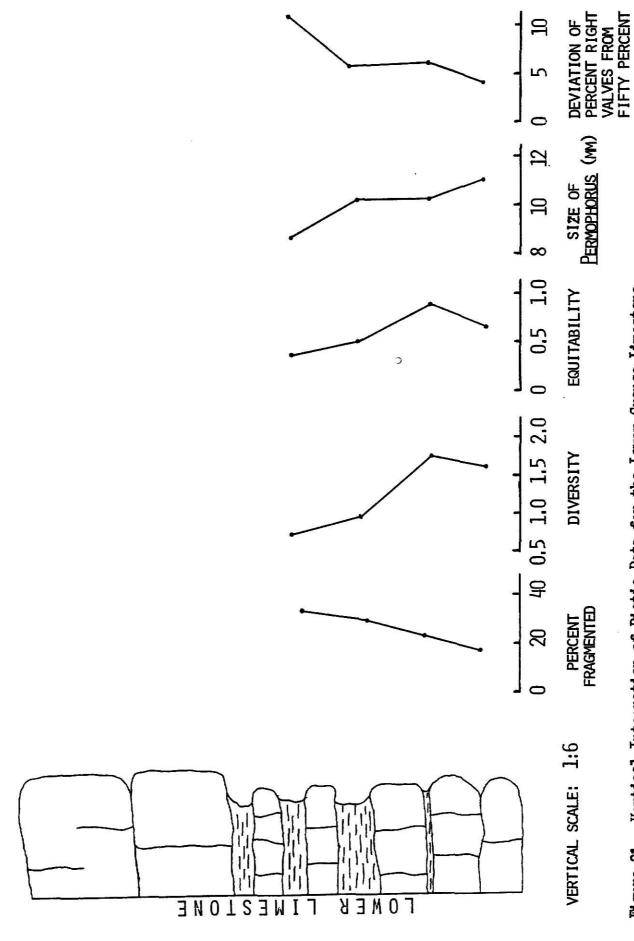
West and McNahon (1972) examined changes in lower Crouse bedding surface assemblages and found (1) an upward increase in concave downward shells, (2) an upward increase in fragmentation, (3) an upward decrease in relative bivalve abundance, and (4) an upward decrease in percent right valves. They interpreted the assemblages as representative of a regressive sequence with shallow subtidal conditions at the base and low intertidal conditions at the top. The trends noted were considered to be mainly a result of higher energy conditions in the upper part.

Increases in fragmentation and deviation of the percent right valves from 50 percent (fig. 21) agree with the observations of West and McMahon (1972) and suggest an increase in energy (<u>i.e.</u> higher energy).

Trends shown in Figures 20 and 21 suggest an increasingly harsh environment during Crouse Limestone deposition. Such environmental changes would place more and more stress on the organisms and could result in (1) fewer organisms (decrease in grain/matrix ratio because grains are dominantly skeletal), (2) a less diverse assemblage (decrease in diversity)



Vertical Integration of Petrologic Data for the Crouse Limestone (composited from all sections). Points plotted are averages for each unit and the number of values averaged are in parentheses. Figure 20.



Vertical Integration of Biotic Data for the Lower Grouse Idmestone (composited from all mapped bedding surfaces, Appendix III) Figure 21.

(3) dominated by fewer species (decrease in equitability), and (4) smaller sized individuals (decrease in size of <u>Permophorus</u> cf. <u>P. subcuneatus</u>). A shallow regressive sequence like that proposed by West, <u>et al.</u> (1972) would indeed be such an environment. Shallower water and more exposure would place more stress on the organisms. The trend of decreasing suspension feeders also supports a regressive interpretation. Regression would have a particularly adverse effect on shallow subtidal to low intertidal suspension feeders because of greater turbulence and more suspended sediment.

In general, data collected from the lower limestone, the middle shale, and the upper limestone support a regressive interpretation.

Bedding surfaces often mark a change in conditions of sedimentation (Hills, 1972, p. 7). Most bedding surfaces in the lower limestone are irregular and there is usually a thin mudstone between limestone beds. Irregular bedding surfaces represent interruptions in normal sedimentation which were probably caused by periods of increased water movement (i.e. storms). A period of increased agitation would result in an irregular bottom because of submarine erosion and increased wave base. It would also be responsible for an increase in terrigenous sedimentation represented by the thin mudstones. The wave produced orientation of shells (Pearson, 1972) is further evidence of increased turbulence.

The lower surfaces of limestone beds signify a return to normal carbonate sedimentation. Bedding surface data indicate that 88.4 percent of the shells on the upper surfaces of limestone beds are concave downward (hydrodynamically stable). Only 71.8 percent of the shells on lower surfaces are concave downward. Concave up shells are typical of quiet water environments (Johnson, 1965; Clifton, 1971). Although the difference between shell orientations may not be statistically significant, it does

suggest that the lower limestone was deposited in a relatively low-energy environment which was periodically interrupted by higher energy conditions.

The conditions described above are typical of a shallow nearshore (subtidal) environment. Horizontal burrows on some bedding surfaces and relatively high diversity of organisms are further evidence favoring subtidal deposition (Walker and Laporte, 1970).

Fossils in the middle shale are a mixture of organisms typical of marine environments and those typical of freshwater environments (Lane, 1964). As pointed out by West, et al. (1972, p. 87) this is evidence in favor of a nearshore environment.

Disrupted laminations, mottled texture, and reworked, homogenized beds--features indicative of bioturbation (Hardie and Ginsburg, 1977, p. 119)--are common in the middle shale. Extensive bioturbation is an indication of intertidal deposition (Gebelein, 1971).

A conspicuous feature of the upper limestone is its fine laminations which are very similar to those caused by high intertidal to supratidal algal mats in recent tidal flat environments (Hardie and Ginsburg, 1977; Logan, 1974; and Hagan and Logan, 1974). The laminated fabric of the upper limestone is similar to typical stromatolitic fabric described by Monty (1976, p. 194). Flat laminated stromatolites such as those in the upper limestone are indicative of high intertidal deposits (Walker and Laporte, 1970). Laminations give the unit its platy character which is also typical of high intertidal and supratidal deposits (Gebelein, 1971).

Hardie and Ginsburg (1977) considered three possible origins for the supratidal laminated sediment on Andros Island. These are (1) chemical precipitation, (2) normal diurnal tidal flooding, particularly high spring tides, and (3) flooding by storms. Chemical precipitation was ruled out because of the well-sorted, open framework fabric of many of the laminae.

Tidal flooding was eliminated as a possibliity because the normal tidewaters carry very little suspended sediment. In fact, clear water is a feature of all recent carbonate environments (Bathurst, 1975, p. 94). Hardie and Ginsburg were able to verify empirically the storm-related origin of the supratidal sediment and they concluded that storm deposition was the exclusive agent responsible for laminated sediment. Black (1933) also emphasized the importance of storms on Andros Island.

At Cape Sable (Florida) 62 percent of the calcareous sediment is deposited by storms. In the supratidal zone all sediment is derived from waters driven over the area during storms (Gebelein, 1971). Storms probably played a role in the deposition of the Crouse Limestone. The irregular bedding surfaces in the lower limestone may well be storm-related and much of the upper limestone sediment was probably storm-derived. As suggested previously, storms may be responsible for the downcutting of the tidal channel at section Q.

Mudcracks, trails, and vertical burrows were observed in the upper part of the upper limestone. Oscillation ripple marks and possible "birdseye" structures were reported from this unit by West, et al. (1972, p. 88).

All these features are typical of a high intertidal to supratidal environment (Gebelein, 1971; Walker and Laporte, 1970).

Dolomite in the middle shale and the upper limestone may be further evidence of a high intertidal to supratidal environment of deposition for these units (Pray and Murray, 1965).

The origin of dolomite has long been a mystery to geologists. Its common occurrence in the geologic record and the inability to produce it in the laboratory have compounded the mystery. Among the most widely accepted hypotheses of dolomite formation are that it forms (1) by seepage-refluxion (Adams and Rhodes, 1960), (2) from hypersaline brines in a sabkha

environment (Shinn, et al., 1965), or (3) by secondary replacement of calcite.

The seepage-refluxion hypothesis proposes that in restricted waters salinity increases because of evaporation. Gypsum may precipitate raising the Mg/Ca ratio in the brine. The brine becomes heavy and seeps downward displacing lighter connate water. During seepage magnesium-calcium exchange and dolomitization occur.

According to the hypersaline brine hypothesis, pore waters in supratidal sediments become increasingly saline until gypsum precipitates.

This raises the Mg/Ca ratio and leads to the replacement of calcium by magnesium and the formation of dolomite.

The secondary replacement hypothesis proposes that dolomitization is accomplished by metasomatic fluids during late diagenesis.

The erratic distribution of dolomite seems to support a secondary origin. However, the occurrence of dedolomite indicates that at one time dolomite was in both the middle shale and upper limestone at virtually every locality (except sections 0, R, and T). It is dedolomitization that is erratic.

Recent primary or early diagenetic dolomite has been found in a variety of environments. Several workers have found dolomite forming from hypersaline water (Illing, et al., 1965; Shinn, et al., 1965). Thompson, et al. (1968) reported dolomite from a deep sea environment with normal marine salinity. Freshwater dolomite has also been reported (Müller, 1970; Sherman, et al., 1962).

Folk and Land (1975) related the formation of dolomite to two factors:

(1) Mg/Ca ratio of the water and (2) salinity (fig. 22). Dolomite may be produced from calcite by increasing the Mg/Ca ratio or by decreasing salinity. According to Folk and Land, dolomite can form better in less saline waters

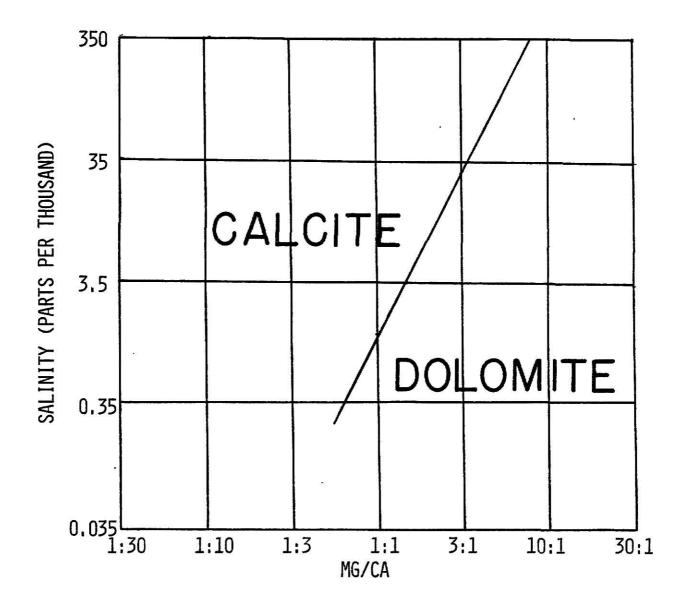


Figure 22. Calcite and Dolomite Stability Fields (after Folk and Land, 1975, p. 61)

because of slower crystallization rates and lower concentrations of competing ions. Perhaps the best way to form dolomite would be to introduce fresh water into marine or hypersaline water. A high Mg/Ca ratio would not be necessary. This does not deny that dolomite may also form in hypersaline environments with high Mg/Ca ratios. In fact, the two processes compliment each other nicely. Hypersaline water with a high Mg/Ca ratio is most likely to be found in a nearshore environment. The effect of freshwater runoff is also greatest there. Therefore, a nearshore environment provides two possible mechanisms for dolomite formation.

The absence of dolomite in section R and in the middle shale at section T could be expected if the dolomite formed in an intertidal or supratidal environment. Sections R and T are the southernmost exposures and occupy the lowest structural positions. Therefore, during regression these sections would be the last to "dry up"--i.e. reflect high intertidal to supratidal conditions favorable for dolomite formation. Section 0, which had dolomite in the middle shale only, is on the Abilene Anticline (fig. 1) where different conditions could be expected.

Dolomite commonly occurs with gypsum in recent supratidal flats (Sherman, 1963; Deffeyes, et al., 1965). Gypsum formation may raise the Mg/Ca ratio and lead to dolomitization. In the Crouse Limestone elongate gypsum crystals were observed at sections Q and S. Cavities the size and shape of those crystals were found at all other outcrops, and they are interpreted as molds of gypsum crystals. These molds were found in different beds of the Crouse Limestone, from the base of the lower limestone at sections W and I to the upper limestone at section T. Their greatest abundance is in the middle shale and the upper limestone. Similar cavities have been described by Summerson (1966) from the Silurian of Ohio, Randazzo (1969) from the Devonian of North Carolina, and Hoffman (1967) from the Precambrian of British

Colombia. All have interpreted the cavities as molds of evaporite crystals.

There is no correlation between the stratigraphic occurrences of gypsum and dolomite in the Crouse Limestone. Hence there is no direct evidence that they are related. Nevertheless, these two minerals are not incompatible in a nearshore sedimentary environment. Their occurrence together in the middle shale and upper limestone is evidence in favor of a nearshore hypersaline environment of deposition (embayed lagoon) for these units. Russell (1977) found evidence suggesting an arid climate during deposition of the Council Grove Group. Such a climate would indeed be conducive to formation of both gypsum and dolomite.

In summary, the results of this investigation support the interpretation of the Crouse Limestone as a regressive tidal flat sequence (West, et al., 1972). Depositional environments ranged from subtidal during lower limestone deposition to supratidal during upper limestone formation.

This regression need not have involved any change in sea level. A prograding tidal flat model (Lucia, 1972, p. 160-163) explains such a sequence without any sea level change. Carbonate mud produced on an adjacent carbonate platform migrates shoreward. This sediment is trapped and builds up on the tidal flat. Regression and progradation take place not because sea level is lowered, but because the tidal flat becomes higher through the accretion of sediment (Ginsburg, 1971; Matti and McKee, 1976). The units above the erosional surface at section Q are interpreted as tidal channel and swamp deposits and are not inconsistent with a prograding shore model.

## Lateral Integration and Interpretation

North-south and east-west cross sections of the Crouse Limestone are in Figures 23 and 24 respectively. Total thickness of the Crouse increases to the south and west. The lower limestone and the middle shale thicken to

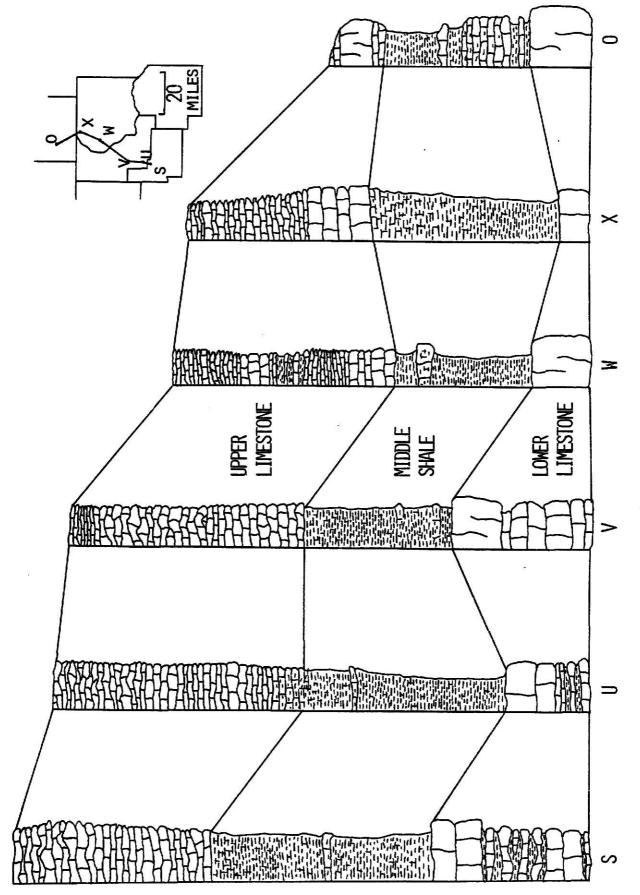
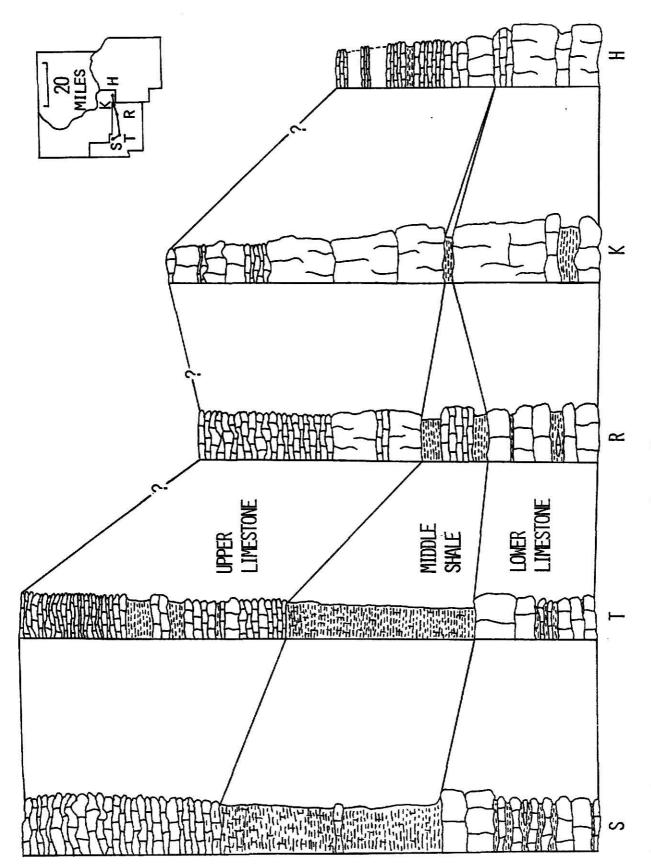


Figure 23. North-South Cross Section of the Crouse Idmestone



Mgure 24. East-West Cross Section of the Grouse Idmestone

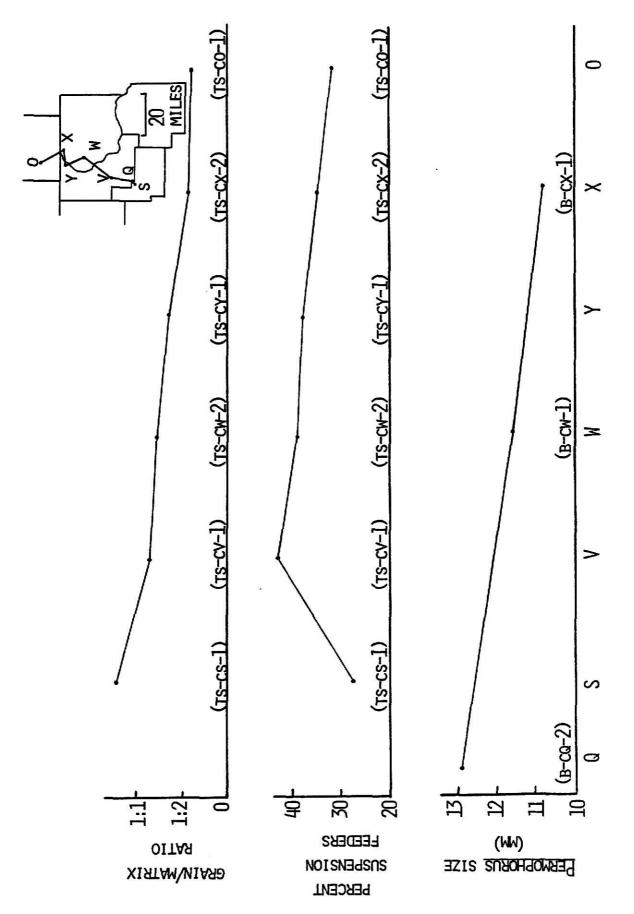
the west and all three units thicken southward. The area occupies the east flank of the south-plunging Irving Syncline (fig. 1). Therefore, the Crouse Limestone thickens toward structurally low areas.

Figures 25 and 26 are the lateral integration of data in north-south and east-west directions respectively. These data were obtained from stratigraphically equivalent bedding surfaces and thin sections, mostly from the lower limestone. There are northward decreases in (1) the grain/matrix ratio, (2) percent of suspension feeding organisms, (3) size of Permophorus of. P. subcuneatus along with eastward decreases in (1) the size of Permophorus of. P. subcuneatus, (2) diversity, and (3) equitability. These trends suggest a more harsh environment possibly due to shallower water in the northern and eastern parts of the area. An eastern increase in percent of suspension feeders does not agree with the other trends and may be due to north-south distance between sections.

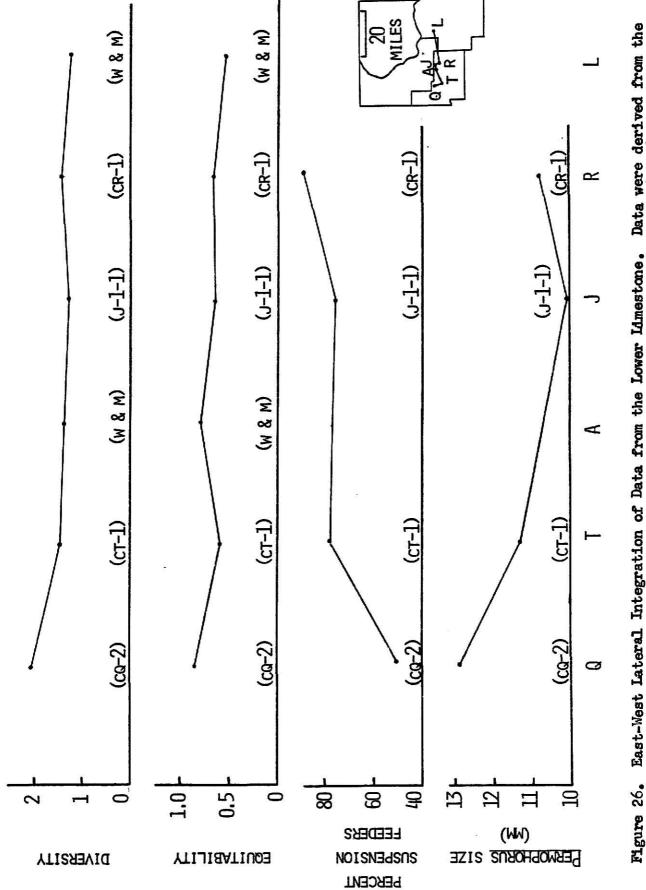
Although the lateral trends examined by West and McMahon (1972) were different from those used in this study, the conclusions are the same. Bedding surface assemblages in the lower limestone suggest shallow subtidal conditions in the west and south and low intertidal conditions to the east and north. Areas of supposed deeper water coincide with the structurally low areas (fig. 1).

The area was probably protected from the open ocean; the Abilene Anticline to the west may have served this purpose. It need not have been an exposed land area, but could have protected the area equally well as a submarine feature. A protected environment is suggested by several lines of evidence.

Permophorus cf. P. subcuneatus is a semi-infaunal, endobyssate bivalve. Endobyssate bivalves are restricted to stable substrates (Stanley, 1972, p. 204). If not protected from the open ocean, the carbonate mud substrate



North-South Lateral Integration of Data from the Lower Idmestone. Data were derived from the samples indicated in parentheses; "B" indicates a mapped surface, "IS" indicates thin section. Figure 25.



East-West Lateral Integration of Data from the Lower Idmestone. Data were derived from the bedding surfaces indicated in parentheses. Data from West and McMahon (1972) are indicated by "W & M."

of the lower limestone would likely undergo frequent periods of instability.

Stability of the substrate is a determining factor in the growth form of stromatolites (Hardie and Cinsburg, 1977, p. 57). An unstable substrate allows stromatolite growth only at isolated spots which results in stromatolite heads similar to those at Shark Bay, Australia (Logan, 1961). Laminar stromatolites like those in the upper Crouse Limestone form on a uniform stable substrate which is more likely if nearby waters are calm and protected.

This is not to imply that the water was completely still and motionless. The Abilene Anticline cannot be detected south of Abilene and could not have caused complete isolation of the area. Connection with the open ocean must have existed and with it some movement of water. Furthermore, some agitation of the water is required by Osagia (Moore, 1964, p. 343).

Clark (1976) in studying populations of <u>Argopecten gibbus</u> in Harrington Sound, Bermuda found a distinct difference in shell convexity between populations from different water depths. Shells from depths greater than ten meters had a lower convexity (flatter shells). Clark considered possible environmental causes of this difference and concluded that turbulence was the factor most likely controlling shell convexity.

Argopecten gibbus and Aviculopecten sp. are similar morphologically (Cox, et al., 1969). Both are members of the Superfamily Pectinacea and their modes of life are the same (epifaunal byssally attached).

Convexity was measured on two specimens of <u>Aviculopecten</u> sp. from bedding surface CQ-2 which were the only specimens not fragmented, distorted, or partly covered. These data are compared to Clark's (1976) measurements (p. 607, fig. 2) in Table 9. The convexity/height ratio of <u>Aviculopecten</u> sp. is very similar to that of the deep water form of <u>Argopecten</u> gibbus.

Several possible interpretations of these results are: (1) The lower Crouse Limestone was deposited in water more than ten meters deep.

Table 9
Shell Convexity in Pectinaceans

Taxon	Height (mm)	Convexity (mm)	Convexity Height
#Aviculopecten sp.	28	6	0.214
#Aviculopecten sp.	18	4.5	0.250
*Argopecten gibbus	56	12	0.214
**Argopecten gibbus	62	24	0.387

<sup>#</sup> from bedding surface CQ-2

(2) The lower Crouse was deposited in non-turbulent water less than ten meters deep. (3) Aviculopecten sp. was transported into the area from a deep water and/or non-turbulent environment and has little direct bearing on the interpretation of Crouse Limestone depositional environments in this area. (4) Despite hard part morphological similarities Argopecten gibbus and Aviculopecten sp. differ biologically and cannot be meaningfully compared. The first two interpretations depend on Aviculopecten sp. being buried at or near the spot where it lived. The conclusion was reached earlier in this report that most organisms on lower limestone bedding surfaces probably lived very close to the spot where they were buried. It seems unlikely that a single species could be transported into an area while the remainder of the assemblage was not noticably affected. Nevertheless, high fragmentation of Aviculopecten sp. (41.7 percent compared to 20.3 percent for the entire assemblage) may be an indication of transportation (Johnson, 1960).

<sup>\*</sup> Clark's (1976) deep water form

<sup>\*\*</sup> Clark's (1976) shallow water form

Orientation data indicate that Aviculopecten sp. along with other taxa was wave-oriented (Pearson, 1972). However, mere shifting of the shell's position may produce orientation without any significant transportation. A shallow water environment with low turbulence is consistent with other lines of evidence and, I believe, is the correct interpretation.

# SUMMARY

The Crouse Limestone is interpreted as a regressive carbonate tidal flat sequence. Characteristics of the lower limestone fossil assemblages are (1) high diversity, (2) wide size range of individuals, and (3) many macroscopic invertebrates compared to the overlying parts of the Crouse. All these are typical of subtidal environments. Bioturbation, a mixed marine and freshwater assemblage, and dolomite are evidence for a low to high intertidal environment of deposition for the middle shale. A high intertidal to supratidal environment for the upper limestone is suggested by (1) laminated algal stromatolites, (2) low species diversity, (3) low abundance of individuals, (h) mudcracks, (5) dolomite, and (6) vertical burrows.

This tidal flat environment was probably protected from the open ocean to the west, perhaps by the Abilene Anticline. It was certainly a low energy environment as shown by (1) laminar growth form of stromatolites, (2) the occurrence of <u>Permophorus</u> cf. <u>P. subcuneatus</u>, an endobyssate bivalve, and (3) carbonate mud.

Dolomite rhombs or evidence of dedolomitization and gypsum crystals were found in the middle shale and upper limestone at all localities. Both minerals occur in a high intertidal to supratidal environment of deposition and they may have formed contemporaneously. The processes of dedolomitization

and gypsum dissolution took place during late diagenesis and may be related.

Apparent inconsistencies in the northern part of the area (sections N and O) were found to be due to erroneous stratigraphic correlation. Examination of the outcrops and establishment of the proper stratigraphic relationships revealed that they support a regressive carbonate tidal flat environment.

The units above the erosional surface at section Q are also consistent with a regressive tidal flat sequence. They represent deposition in a tidal charmel and tidal pond or swamp environments which are certain to exist on a tidal flat.

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### APPENDIX I

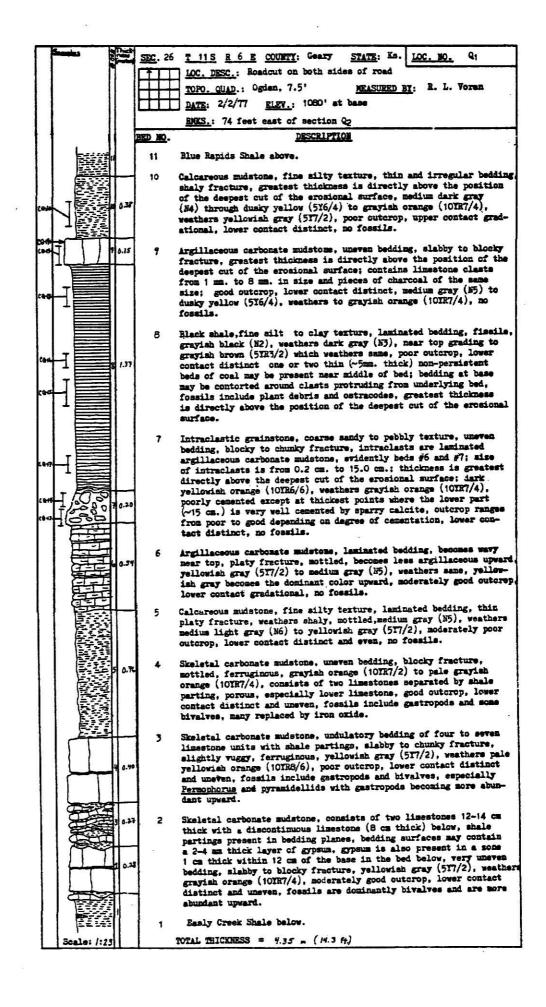
This appendix contains measured sections each of which includes (1) descriptions of each unit measured, (2) thickness of each unit, (3) graphic section, and (4) the locations of all samples collected. The location of each sample is indicated on the graphic section by an arrow or bracket.

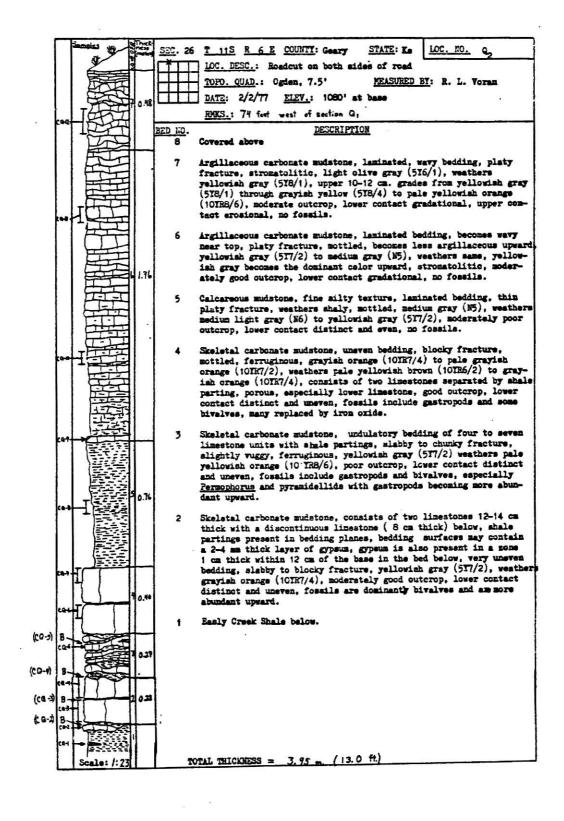
The positions of slabs that were collected for bedding surface mapping are shown by a "B" with the number of the surface shown in parentheses.

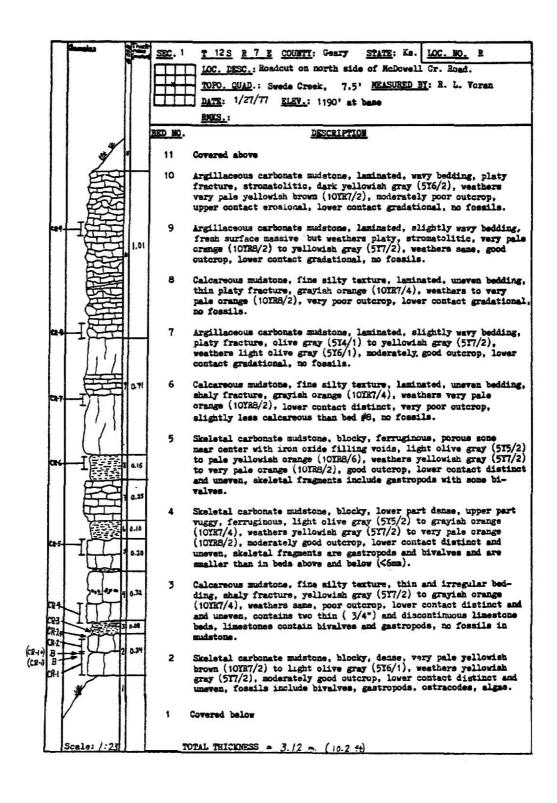
Example of numbering system:

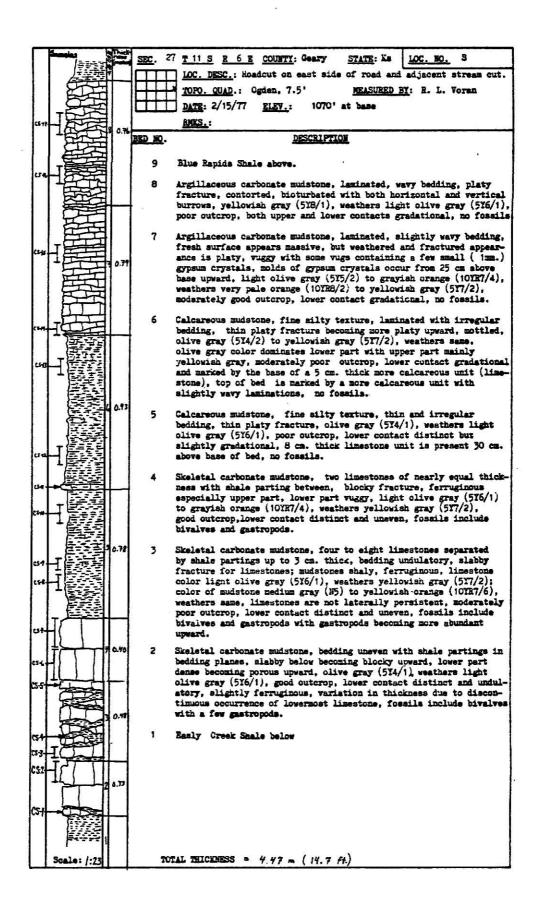
CX-3 = Crouse Limestone, Section X, the third sample collected at that exposure.

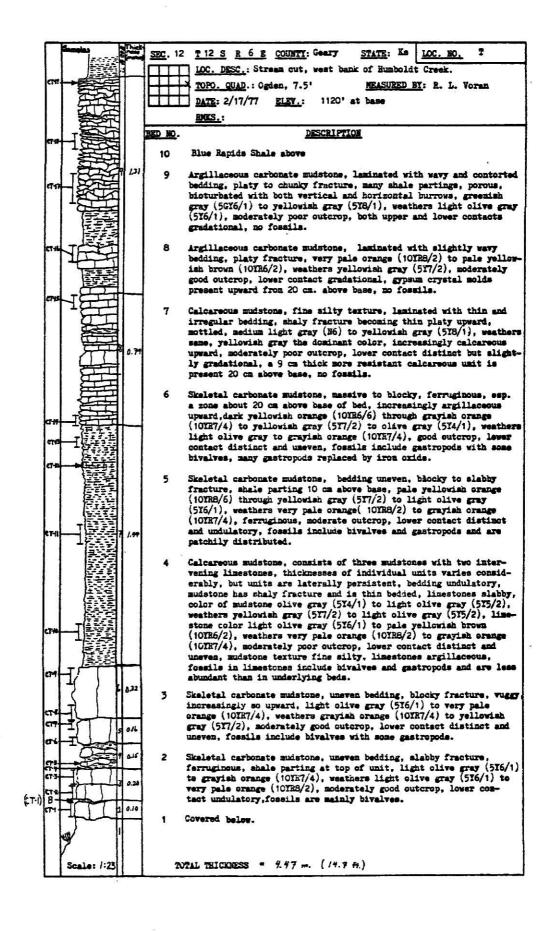
Some samples were split into an upper and lower part after being collected and the upper was designated by a "+" (e.g. CQ-19+).

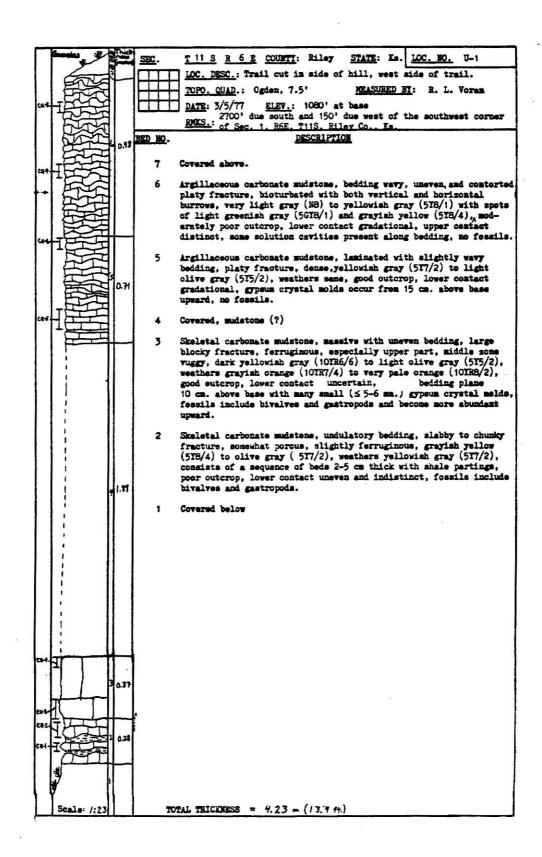


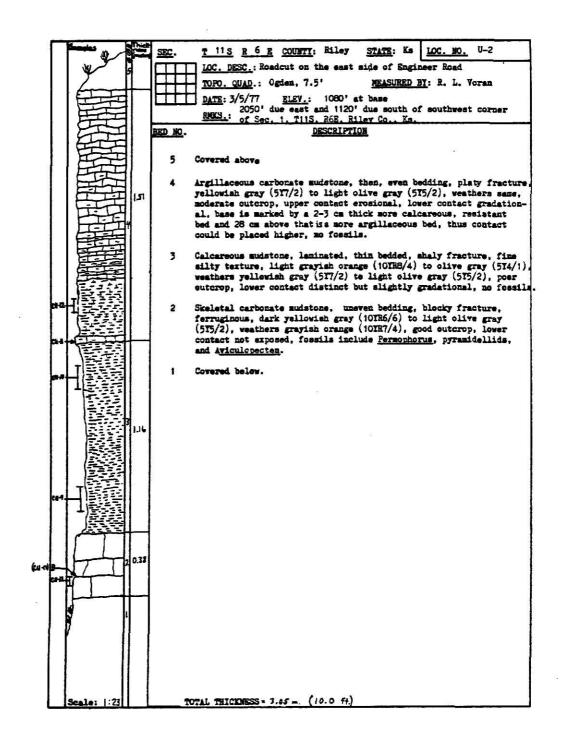


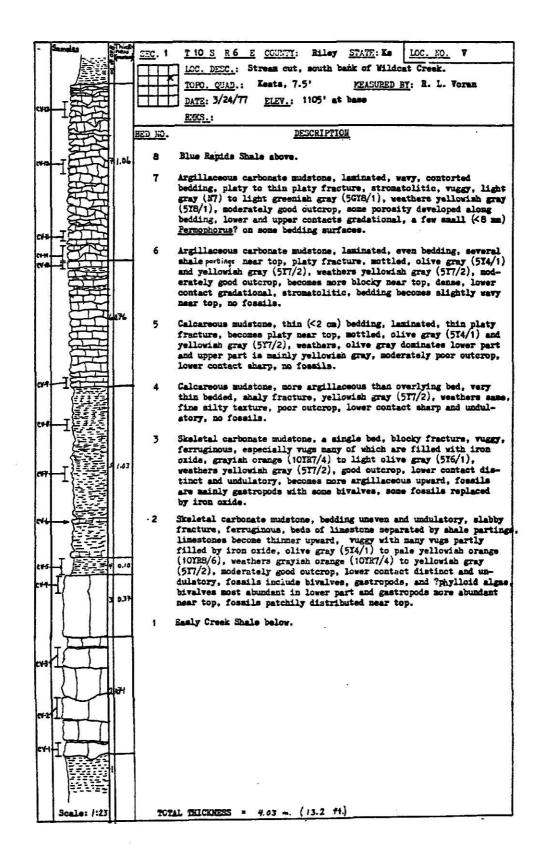


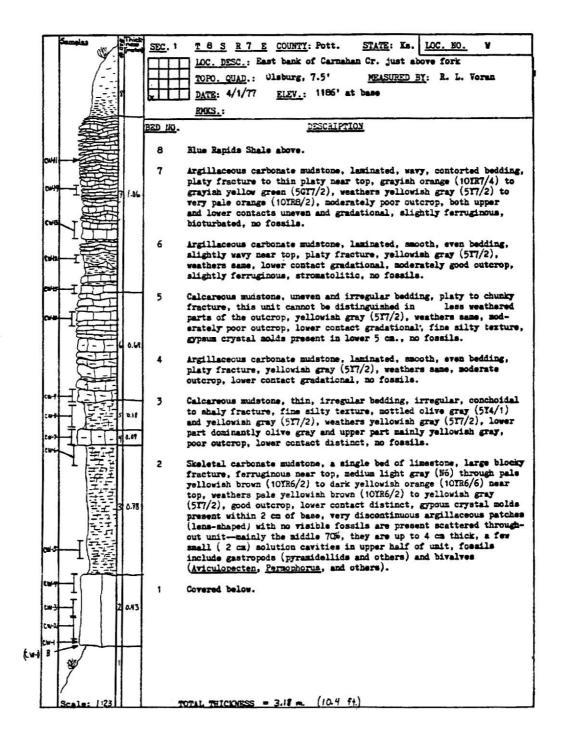


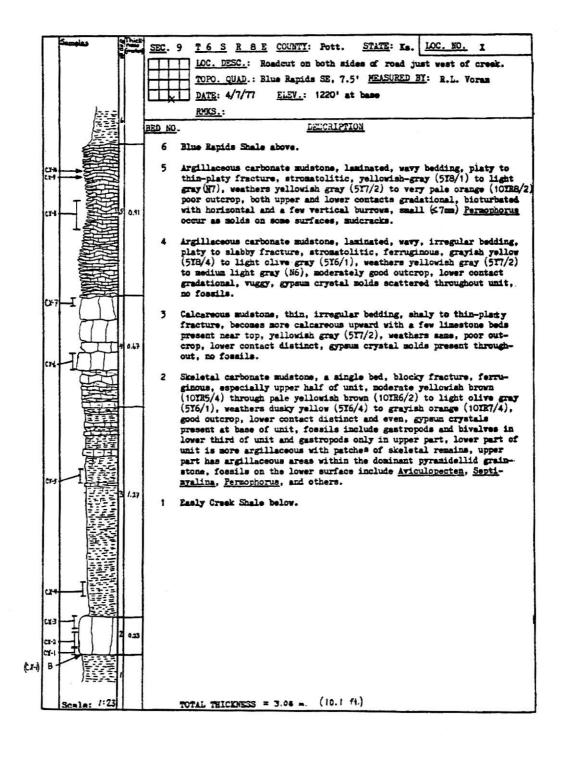


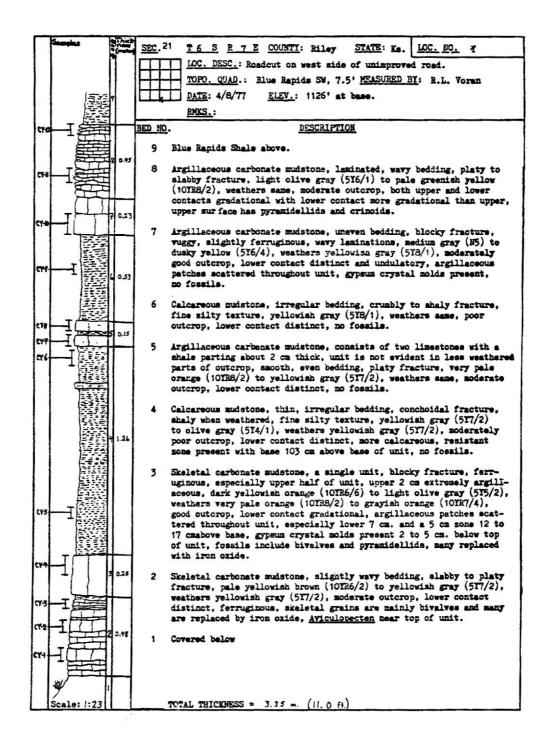


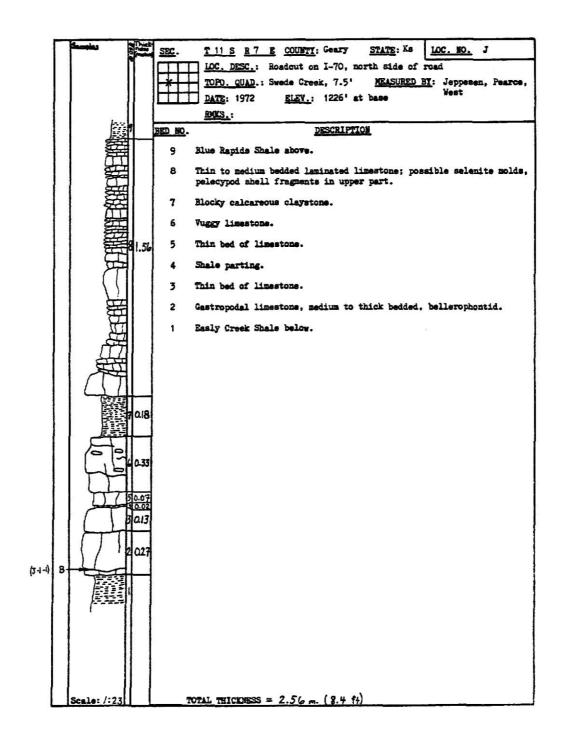


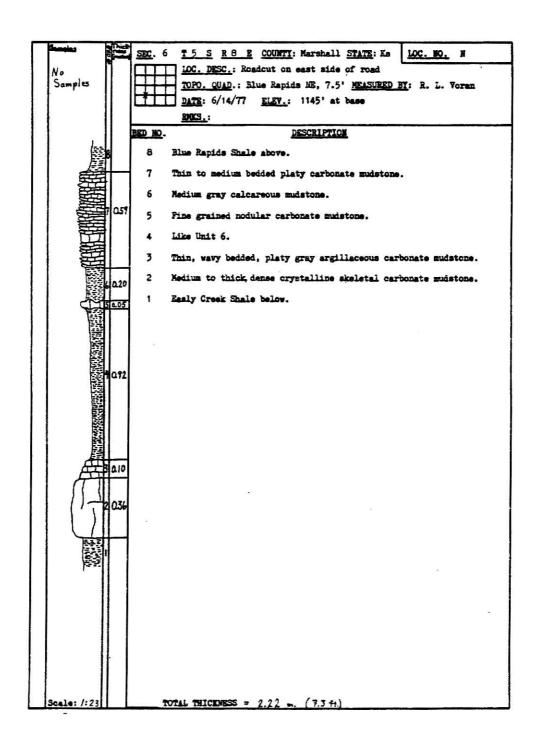




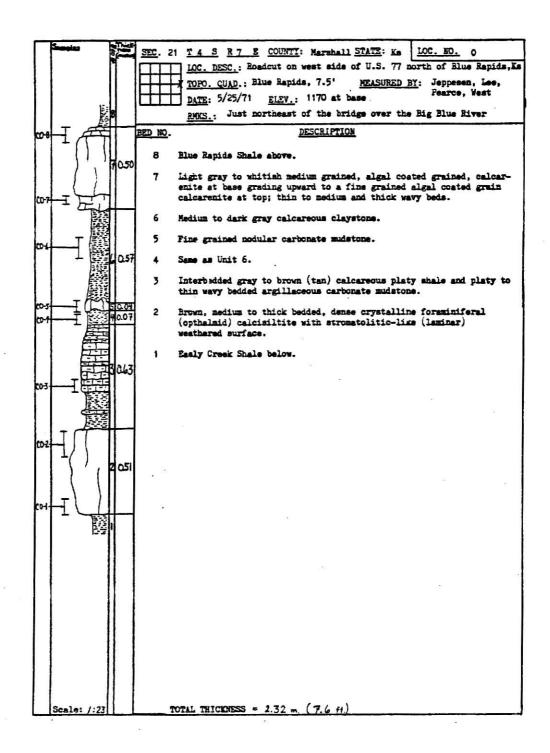








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# APPENDIX II

Argillaceous (bag) samples were disaggregated and the residues examined for microfossils. The following pages contain lists of the microfossils identified in the washed residues. Locations of the samples are shown in Appendix I.

Sample	Taxa
Section O CO-4 CO-6	healdiaceans, bairdiids healdiaceans
Section Q <sub>1</sub>	smooth-shelled ostracodes, plant debris, coal smooth-shelled ostracodes, plant debris Spirorbis sp., plant debris, charcoal, fish debris smooth-shelled ostracodes, plant debris fish debris, Spirorbis sp.
Section Q <sub>2</sub> CQ-8	None
Section R CR-3 CR-6	hollinids, healdiaceans, bairdiids, pyramidellids, Earlandia-like foraminifers, holothurian sclerites healdiaceans, hollinids, bairdiids?
Section S CS-l <sub>1</sub> CS-8 CS-10 CS-12 CS-13	hollinids, healdiaceans, bairdiids?, pyramidellids, holothurian sclerites, <u>Earlandia</u> -like foraminifers hollinids None None healdiaceans
Section T CT-4	pyramidellids, hollinids, healdiaceans, bairdiids, holothurian sclerites, <u>Earlandia</u> -like foraminifers
CT-10 CT-11 CT-13	healdiaceans None healdiaceans
Section U-2 CU-9 CU-10 CU-12	hollinids, bairdiids?, healdiaceans, ammovertellids bairdiids? None
Section V CV-5	hollinids, healdiaceans, bairdiids?
Section W CW-6 CW-13	hollinids, healdiaceans None
Section X CX-4 CX-5	bairdiids, hollinids, healdiaceans healdiaceans

Sample	Taxa
Section CY-5 CY-6 CY-9	healdiaceans, bairdiids? None productacean spine

### APPENDIX III

This appendix contains data from mapped bedding surfaces from the lower Crouse Limestone. It consists of (1) descriptions of fossils observed on these surfaces and (2) maps of the bedding surfaces showing the relative positions of fossils. The stratigraphic positions from which the slabs were collected are designated by a "B" and the number of the surface is given in parentheses on the measured sections (Appendix I). Two separate areas were mapped on surfaces CU-13 and CX-1 to obtain sufficient data. They are designated areas #1 and #2 (e.g. CX-1 Area #1). Some data from surface J-1-1 were lost and not all specimens could be identified by number. Some are identified only by a letter.

Fossil names are abbreviated and are listed below:

### Abbreviation

# Brachiopoda ling = Lingulid Bivalvia Per = Permophorus cf. P. subcuneatus Avic = Aviculopecten sp. Acan = Acanthopecten sp. Pseu = Pseudomonotis sp. Sept = Septimyalina sp. Edm = Edmondia? sp. Gastropoda Bell = Bellerophon sp. pyra = Pyramidellid Echinodermata crin = Crinoid debris Arthropoda ost = Ostracodes alfr = Algal-coated fragment burr = Burrow unid = Unidentifiable

### Fossil Size

The number in this column is the length in millimeters of <u>Permophorus</u> cf. P. subcuneatus. Size of other taxa was not measured.

### Articulation = art

An x in this column indicates that the specimen was articulated.

### Valve = val

The letter in this column indicates which valve of the organism was observed.

r = right valve

1 = left valve

? = valve indeterminate

# Preservation = pres

The following abbreviations are used for type of preservation:

o = original or altered shell

i = internal mold

e = external mold

### Orientation = orient

Orientation of single valves was recorded. The following abbreviations are used to indicate orientation:

cu = concave up

cd = concave down

i = inclined to bedding

# Fragmentation = frag

An x in this column indicates that the fossil was fragmented.

# Epizoans = epiz

An x in this column indicates that epizoans were attached to the shell. The only epizoans observed were tubiform sedentary foraminifers and encrusting algae, usually intergrown in Osagia colonies.

# THIS BOOK CONTAINS NUMEROUS PAGES THAT WERE BOUND WITHOUT PAGE NUMBERS.

THIS IS AS
RECEIVED FROM
CUSTOMER.

1 Avic ? 0 cd x 2 Avic ? 0 cd x 3 unid 0	x x
2 Avic ? 0 cd x 3 unid 0 x 4 Per r i cd x 5 ost r 0 cu 6 pyra i x 7 Per 10.0 ? i cd 8 unid 0 x 9 Avic 1 0 cd 10 Per ? i cd x 11 Per 14.0 r i cd 12 pyra	x
11 Per 14.0 r i cd 12 pyra i	
11 Per 14.0 r i cd 12 pyra i	
11 Per 14.0 r i cd 12 pyra i	
11 Per 14.0 r i cd 12 pyra i	
11 Per 14.0 r i cd 12 pyra i	
11 Per 14.0 r i cd 12 pyra i	
11 Per 14.0 r i cd 12 pyra i	x
11 Per 14.0 r i cd 12 pyra i	
12 pyra i	
13 Per 8.0 r i cd 14 unid o x 15 Per 18.0 l i cd 16 ost l o cu	
14 unid 0 x 15 Per 18.0 1 i cd 16 ost 1 o cu	
15 Per 18.0 1 1 cd 16 ost 1 o cu	
16 ost 1 o cu	
10 050	
17 Per 18.0 1 i cd	
18 pyra i x	
19 Avic r o cd x	x
20 pyra i	•
21 Avic r o cd	x
22 Avic r o cd	x
23 Not a fossil	
23 Not a fossil 24 unid o x	
25 unid o x	x
26 unid o x	
27 Per ? i cd	
28 Acan ? o cd	
29 Avic 1 o cd	X
30 Per 8.0 r i cd	
31 Per 10.0 ? i cd	
32 Per 13.0 r i cd	
35 Per 10.0 r i cd 36 pyra i	
36 pyra i	
37 pyra 1 x	
38 pyra i	
39 pyra i	
10 Per 15.0 r i cd	
1 o cu 12 ost 1 o cu	
42 ost 1 o cu 43 ost r o cu	
lil Per ? o cd	
lis ost r o cu	
45 ost r o cu 46 ost 1 o cu 47 Sept 1 o cd	
47 Sept 1 0 cd	
48 ost 1 o cu	
48 ost 1 o cu 49 ost 1 o cu	
50 Sept r o cd 51 Per ? i cd	

CQ-2 Cont.

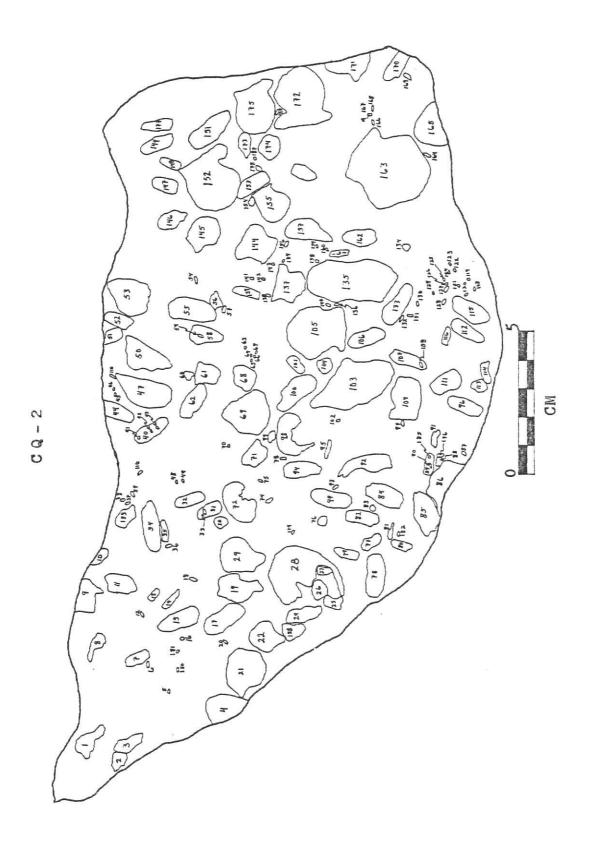
#	name	size	art	val	pres	orient	frag	epiz
52	Per	14.0		r	1	cd		
53	Avic			r	0	cd		x
54	unid				0		X	
53 54 55 56 57 58 59 60	Per	18.0		r	1	cd		
56	unid				0		X	
57	pyra	50.00 <b>=</b> 1 (448			ī			
58	Per	14.0		1	1	cd		
59	pyra				i		e e	
60	unid			52001	0		x	
61	Avic	70.19 <b>4</b> FG		l	0	cd		
62	Per	15.0		1	1	cd		
63 64 65 66	ost			r	0	cu		
<i>6</i> L <sub>1</sub>	ost			r	0	cu		
65	ost			r 1 ? 1 1	0	cd		
66	ost			?	Ο,	cd		
67	ost			1	0	cu		
68	Avic			1	0	cd		x
69	Avic				0	cđ		
70	ost			r	0	cu		
71	Per			r	0	cd	X	
72	Avic			?	0	cd	x	x
73 74 75 76	pyra				i			
74	unid				o i		x	
75	pyra				i			
76	unid				i	cđ		
77 78	Per	11.0		l	i	cd		
78	Per	17.0		r	i	cd		x
79	Per	9.0		1	i	cd		
80	Per	8.0		1	i	cd		
81	unid				0		X	
82	Per	13.0		1	i	cd		
83	Avic			1	0	cđ	X	
84	Per	15.0		l	i	cd		
85	unid				0		x	x
86	Avic			?	o i	cd		x
87	pyra			22		<u>120</u>		
88	Avic			1	0	cd	X	
89	Per	9.0		r	i	cd		
90	pyra				i			
91	unid			<u>226</u> 0	0		x	
92	Sept			1	0	cd	x	
93 94	pyra				i	100		
94	Per	15.0		r	<u>1</u>	cd		
95 96	unid	8 32 88		2:	0	2	X	
96	Per	15.0		1	i	cđ		
97 98	pyra			_		200-750		
98	Avic			? ?	0	cd	x	x
99	Avic		10		0	cd	x	
100	Per	18.0		r	i	cd		x
101	Per	9.0		r	i	cđ		
102	ost			r	0	cu		

CQ-2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103	Sept			1	0	cd		
104	unid				0		X	
105	Avic			1	0	cd		X
106	Sept	F023.		1	0	cd	x	
107	Per	14.0		1	i	cđ		
108	pyra				i		-55	
109	Avic			1	0	cđ		x
110	pyra	328			i			
111	Per	16.0		r	1	сф		
112	Per	13.0		1	i	cd		
113	unid				0		x	x
114	unid	92			0		x	x
115	Per	14.0	X		1	cd		
116	Per	6.0		?	i	cd		
117		fossil				<u>e</u>		
118	ost			r	O	cd		
119	ost			1	0	cd		
120	ost			r 1 1 1	0	cu		
121	ost			l	0	cu		
122	ost				0	cu		
123	ost			r ?	0	cu		
124	ost			?	0	cd		
125	crin				0	2		
126	ost			r	0	cd		
127	pyra				i		X	
128	pyra			440	i			
129	ost			1	0	cu		
130	pyra				i i i			
131	pyra				i			
132	pyra				1			352
133	Per	16.0		r ?		cd		X
134	Avic				0	cu	X	
135	Per	31.0		1	i	cđ		X
136	pyra			<b>=</b> :	i			
137	Avic			1 1	0	cd	x	X
138	ost				•	cd		
139 140	Per	9.0		r	i	cd		
140	pyra			4	1			
141	ost			1	0	cu		
142	ost			+	٥	cu		
143	ost			1 1 1	0	cd		
1 44	Avic				0	cđ		x
145	Avic	0.0		r	o i	cd		x
146	Per	9.0		r ?	1	ed		x
1),2	Per Per	11.0 8.0		r	i	cd		x
147 148 149	rer Per	12.0	•	I.	i	cd cd		-
150		fossil	x		-	eu		x
150 151	Per	18.0		r	i	cd		-
152	Avic	10.0		r	0	cd		x
153	Per	14.0		r	1	cd		
. , ,	T CT	14.0		•	-	Gu		

CQ-2 Cont

<del>#</del>	name	size	art	val	pres	orient	frag	epiz
154 155	unid Avic pyra			1	0 0 1	cd	x	x
156 157	Sept			1	0	cd		
158 159	ost				0	cu		10
159	ost			r	0	cu	2	
160	ost			1	0	cu		
161	Per	9.0		r	i	cd		
162	Per	12.0		r	i	cđ		
163	Avic			1	0	cd		
164	pyra				i	- 4		
165 166	Per			r	i	cd		x
167	oyra ost			r	0	cu		
168	pyra			1	i	Cu		
169	unid				0		x	
170	unid				o		x	
171	Avic			?	0	cd	HODES.	x
172	Avic			r	0	cd		x
173	Per	11.0		1	i	cd		
174	unid				0	cd		x
175	Avic			?	0	cd	X	
176	unid				0		x	
177	Per	10.0		1	i	cđ		
178	Avic			r	0	cd		x
179	pyra			_	i			
180	ost			1	0	cu		
181	ost			1	0	ca		
182	pyra			-	i			
183	Per	6.0		1	1	cd		
184	pyra				i			
185 186	pyra				0		*	
187	unid				i		x	
101	pyra				-4-			



(7

#	name	size	art	val	pres	orient	frag	epiz
1	Avic			?	0	cd		
	unid				0	-	x	
234567890	unid						x	x
L	unid				0		x	
3	Per			?		cd		
6	pyra				: ::::::::::::::::::::::::::::::::::::		x	
7	pyra				i			
8	Per	11.0		r	i	cd		
9	pyra				1			
10	Per			?	i	cd		
11	unid						x	x
12	pyra				i			
13	Per			?	0	cd		
14	unid						x	
15 16	pyra				i i i			
16	pyra				i			
17	Per			?	i	cd		
18	Avic			?	0	cu	X	
19	ost			1	0	cu		
20	ost			1	0	cd		
21	unid				0		x	
22	pyra				i			
23	Per	7.0		r	* * * 1 * 1 * 1	cd		
24	pyra	*****			i			
25 26	pyra				i			
26	Per			?	i	cd		
27	ling			?	0	cd	x	
28	unid				0		x	x
29	Per	17.0		1	0	cu		
30	Avic			r	0	cd		
31	unid				0		x	
32	Per	-		?	i	cd		
33	unid				0			X
34	Per	8.0		1	i	cd		
34 35	Avic			?	0	cd		
36	Avic			r	0	cd		x
37	pyra							
38	pyra				i i i		x	
39	Per	13.0		1	i	cd		
40	Avic			?	0	cd	X	x
41	Avic			?	0	cd		X
42	Avic			7 ? 1 ?	0	cd		
43	Avic			?	0	cd		x
44	unid				0		X	x
45	pyra				i			
46	Per	11.0	x		i	cd		
47	Avic			?	0	cd	X	
48	Per	7.0		r	0 i	cd		
49	pyra			46	i			
33783904423456789051	ost			1	o i	cu		
51	pyra				7			

CQ-3 Cont.

#	name	size	art	val	pres	orient	frag	epi.z
5234567890 61	Per Per ost Per Per pyra	6.0 16.0 9.0 8.0		1 1 1 1 1	i 0 i i	cd cd cu cd	x	x
62	pyra pyra pyra Per Per pyra	12.0 12.0		r	को रने को को को को को को को को को	cd cd		x
63 64 65 66	Per pyra pyra	11.0		r	11111	cd		
67 68 69 70	Per Avic Avic Per	12.0		? ? ? 1	i 0 1	ed ed ed ed	x	
71 72 73 74	Per pyra pyra Per	9.0		r	: : : :	cd cd		
75 76 77 78	pyra Avic Per	6.0		r 1	0	od on	x	x
79 80 81	pyra Per pyra pyra	13.0		r	i i i i i i	cd	x	
82 83 84 85	Per pyra unid Per	17.0		1	i : o i	ed ed	x	x
86 87 88	pyra Avic pyra	12.0		?	i	cd	x x	
89 90 91 92	pyra Per Per pyra	12.0		?	11111	cd cd	x	
93 94 95	Sept pyra pyra			?	0 i i	cđ	x	
96 97 98 99	Per unid Avic unid			r	i 0 0	ed ed	x	x
100 101 102	unid Avic pyra			?	i o i	cd	x	Δ.

CQ-3 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103	Per	10.0	x		1	i		
104	ost	10.0		70	0	cu		
105	Avic			r ?	0	cd	x	
106	pyra			•	i	ou	x	
107	unid				0		X	
108	ost			r	0	cu		
109	Per	13.0		r ?	0	cu		
110	unid	. 5.0			0		x	
111	unid				7.7		X	X
112	Per	10.0		r	0	cu		
113	pyra				i		x	
114	Per			?	i	cd		
115	pyra				i			
116	ost			1	0	cd		
117	Avic			?	0	cd		
118	Per			?	0	cu		
119	unid				0		x	
120	ost			r	0	cu		
121	ost			r	0	cu		
122	ost			r	0	cu		
123	unid				0	cd	x	x
124	Per	13.0		r	i	cd		
125	pyra				i			
126	ost			1		cd		
127	pyra				o			
128	Per			?	0	cd		
129	pyra				i			
130	ost			1	0	cu		
131	Avic			?	0	cu	x	
132	Per	6.0		1	i	cd		
133	ost			1 ? 1 ? 1	0	cu		
134	Avic			?	0	cd	X	
135	ost			1	0	cd		
136	ost			r	0	cd		
137	unid				o i			
138	pyra							
139	ost			r	0 1 1 1	CII		
140	pyra				i			
141	pyra				i			
142	pyra							
143 144 145	ost			r	0	cd		
144	ost			r	0	cu		
145	pyra				i			
146 147	unid				0 1 1 1 1 1 1		x	
147	pyra	, -		_	1		x	
148	Per	6.0		I	į	cd		
149	Per	10.0		r ?		cd		
150	Avic			?	0	cu	X	
151	Avic	0 -		?	0	cd	x	
152	Per	8.0		?	i	cd		
153	Per	10.0		?	1	cd		

CQ-3 Cont.

#	name	size	art	val	pres	orient	frag	epiz
154 155 156 157 158	Per Per Per ost ost			? ? ?	0 0 0 0	cu cd cd cd		
159 160 161 162	Per Per Per Per	5.0  10.0		r 1 ? ? r ?	0 1 1 1 1 1	cd cd cd		
163 164 165	Per Avic Per	12.0		? r	i 0 i	ed ed ed		
166 167 168	Per Per unid	21.0		1	o i	cu	x	x
169 170 171	pyra unid Per	12.0		1	o i o i	ed	x	x
172 173	pyra				i i i			
174 175 176	Per Per Per	11.0 6.0 13.0		r r r	i i	ed ed		
177 178 179 180	Per pyra pyra pyra	8.0		1	약리 막리 막리 막리 막리 막리 막리	cđ		•
181 182 183 184	Per Per unid	14.0 8.0		r	0	cđ cđ	x	x
185 186	pyra Per Per	10.0		r ?	i i i	cd cd		
187 188 189	pyra Per Per	12.0		1	i	cd cd		
190 191 192	Per Per unid	8.0 7.0		r	i i	cd cd		
193 194 195 196	pyra pyra pyra pyra				જો ની ધો ની ની ની ની તો			
197 198 199	ost Avic ost			1 ? 1	0	cu cd cd		
200 201 202 203	pyra Per Per Per .	9.0 10.0		? ? 1	i i i	cd cd cd		
201	unid	5.0±.1♥.15.		_	0	·	x	x

CQ-3 Cont.

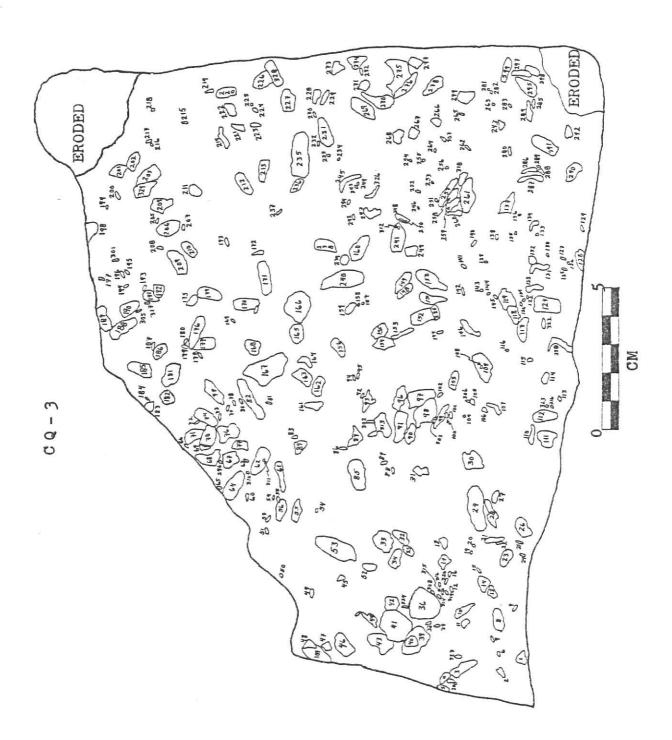
#	name	size	art	val	pres	orient	frag	epiz
205 206 207	pyra Per pyra			1	111111	cđ	x	
208	pyra			_	i	-	X	
209	Per	11.0		1	i	cd		
210	Per	8.0		r	1	cd.	15 223	
211 212	unid	8.0		7	i	- 4	X	
213	Per Per	8.0		1 1 ?	i	cd cd		
214	Avic	0.0		2	0	cd	x	
215	pyra				i	ou	and a	
216	ost			r	0	cd		
217	pyra				i		x	
218	ost			r	0	cd		
219	pyra				i			
220	Per	9.0		1 ?	i	cd		
221	Avic			?	0	cu	x	
222	Sept			l	0	cd	x	
223	Per		X		i	i		
224	unid				0		X	x
225	ost			l	0	ou.		
226	Per	9.0		1	i	cd		
227	unid			-	0		, X	x
228	Per	6.0		7	i	cd		20
229 230	Avic			*	0	cd	x	
231	pyra Per	10.0		1	1 1	ed		
232	pyra	10.0		-	-	Cu		
233	ost			ı	0	cu		
234	ost			r	0	cd		
235	Per	18.0		r	i	cd		
236	Per	7.0		r	i	cd		
237	pyra				i			
238	unid				0	2	x	
239	Avic			r	0	cd		
240	Per	17.0		r	i	cd		X
241	Per	11.0		1	i	cd		
242	pyra			7		-		
243 244	ost			1 1 ?	0	cu		
245	Per	12.0		2	0	cu		
245	ost	12.0			0	cu		
247	Per	11.0		r	i	cd		
248	unid				0	-		
249	unid						x	
250	pyra				0			
251	ost			r	0	cd		9
252	pyra				i		x	
253	ost			r	0	cu		
250 251 252 253 254 255	ost			r	0	cu		
255	ost			1	0	cd		

CQ-3 Cont.

#	name	size	art	val	pres	orient	frag	epiz
256 257	ost Per	11.0		1 r	0	cu		
258	Per			? ?	0	cd	x	
259	Per				0	cd	x	
260	Per			?	0	cd	x	
261	Per	15.0		1	0	cu	X	
262	pyra				i			
263	ost			r	0	cu		
264	ost			l	0	cu		
265	pyra				i		221	
266	unid				0		x	
267 268	unid Per			?	0	cd	×	
269	Per	12.0			4	cd		
270	Per	12.0		i	4	cd		
271	Per	4.0		r 1 1	11111	cd		
272	pyra	4.0		and i	ร้	04		
273	unid				0		x	x
274	unid				0		x	X
275	unid				0		x	x
276	Per	14.0		?	i	cđ		
277	Per	am 90		?	i	cd		
278	unid				0		X	
279	Per			?	0 '	cu		
280	pyra			-	i	er g		
281	ost			1	0	cd		
282	ost			1	0	cd		
283 284	ost unid			1	0	$\operatorname{cd}$	12	198
285	pyra				o i		x	x
286	Avic			?	0	cu	x	
287	Avic			?	0	cu	x	
288	Avic			?	0	cd	x	
289	ost			1	0	cu		
290	Per	8.0		1	i	cd		
291	Per	12.0		r	е	cu		
292	Per	5.0		1	i	cd		
293	unid				o i i		x	x
294	pyra				i			
295 296 297	Per			?		cd	x	
290	ost	7 0		r 1 1	0	cu		
298	Per Per	7.0 9.0		<u>+</u>	i i	cd cd		
299	Per	7.0		7	4	cd		
300	unid	1.0		-		Ca	Z	
301	pyra				i			
302	pyra				0 i i i i i i			
303	pyra				i			
304	pyra				i			
305	pyra				i			
306	pyra				i			

CQ-3 Cont.

#	name	size	art	val	pres	orient	frag	epiz
307	pyra				i i			
308	pyra				i			
309	unid				0		X	x
310	ost			r	0	cu		
311	ost			r	0	cd		
312	ost			r 1 ?	0	cd		
313	Sept			?	0	cd	X	
314	pyra				i		x	
315	pyra				i			
316	ost			r	0	cd		
317	ost			r 1 ?	0	cu		
318	Avic			?	0	cd	X	
319	pyra				i i i			
320	pyra				i			
321	pyra				i			
322	pyra				i			
323	ost			1	0	cu		
324	pyra				i			
325	Per			r ?	i	cd		
326	Per			?	0 1 1 1	cd		
327	pyra				i			
328	Per			? ?	0	cu		
329	Per			?	0	cu		
330	Avic			?	0	cd	x	



#	name	size	art	val	pres	orient	frag	epiz
1	pyra				i			
5	unid Per			?	0	cu	x	
23456789	Per	12.0		i	id id id id	cd		
3	pyra	F		_	i	200 (C.E.)		
6	pyra				i		x	
7	Avic	10.0		?	0	cd	x	
8	Per	10.0		?	i	cd		
10	pyra	10.0		1	i	ed		
10 11	Per pyra	12.0		7	o	Ga	x	X
12	Per	7.0		1	i	cd	-2-	
13	pyra	. • •			i		X	
13 14	Per	8.0		?	i	cd		
15	pyra				i			
16	pyra				i		x	
15 16 17 18	pyra	10.0		***	i	cd		
19	Per	10.0		r	1 1	ca		
20	pyra pyra				i			
21	pyra				ī			
22	Per	16.0		r	i	cd		
23 24	pyra	mr w			i			
24	Per	9.0		1	i	cd		
25 26	Per	5.0		r	저 저 저 저 저 저 저 저 저 저 저 저 저 저 저 저 저 저 저	cd		
27	pyra				4		x	
28	pyra pyra				i		Δ.	
29	pyra				i			
30	pyra				i			
31 32	pyra				i		x	
32	pyra	0 0			i	5150 <b>4</b>		
33	Per	8.0		r l	1	cd		
3E 3H	Per Per	11.0				ed ed		
36	Per	9.0		? r	i i	cd		
37	Per	9.0 16.0		ī	i	cd		
38	Per	9.0		r	i	cd		
39	ost			r		cu		
40	pyra				1			
41	pyra	40.0		-	i	cd	X	
1.2	Per pyra	10.0		1	1	ca	x	
1,1,	Per	12.0		?	i	cd	of Ba	
45	Per	11.0		? ? ?	1	cd		
46	Per			?	i	cd		
47	pyra				i		700000	
48	pyra			9	1	cd	X	
47	Per Per			?	0 게 귀 게 거 귀 게 귀 귀 쥐 하 귀 쉬	ed		
356789012345678901	pyra	( <del>-101</del> )		: <b>*</b> )	i			
	10							

CQ-4 Cont.

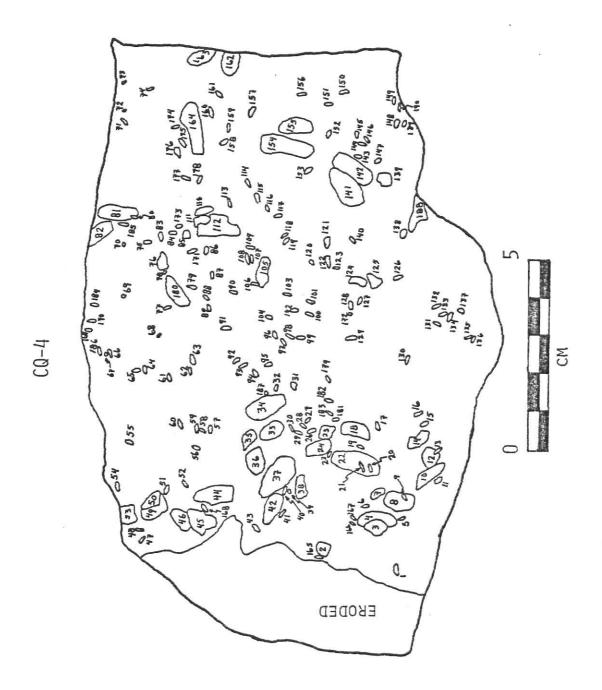
#	name	size	art	val	pres	orient	frag	epiz
52 53 55 55 55 57 58	pyra Per			?	위하시키위하시카위위에 위하기	cd	x	
24	pyra				4		x	
22	pyra pyra				4		X	
57	pyra				i			
58	pyra				1			
59	pyra				i			
60	pyra				i		x	
59 60 61 62 63 64 65 66 67 68 69	pyra				i		X	
62	pyra				i			
61.	pyra pyra				1 4		x	
65	pyra				i		x	
66	pyra				i			
67	ost			1	0	cd		
68	ost			1 1 1	0	cu		
69	ost			ļ	0	cd		
70	ost			1	0	CII		
71 72	pyra			**	i	-	X	
73	ost			r	0	cd		
73 74	pyra			1	i	Cu		
75	pyra				i		x	
75 76	Per			r		on	x	
77	pyra				o i		x	
78	ost			r	0	Cu		
79	pyra				i			
80	pyra	44 0		7	1	cd		
81 82	Per Per	11.0		l r	1 1	cd		
83	pyra	1877		-	i	20.		
83 84	pyra				0 i i i i i			
85 86	unid				0		x	
86	pyra				i			
87	pyra				i			
88	pyra				i		x	
89	pyra				1		x	
90 91	pyra pyra				÷		Δ.	
92	pyra				i			
93	pyra				i			
94	pyra				<u> </u>			
95	pyra				i		X	
93 94 95 96 97	pyra				1		x	
98	pyra				4		x	
99	pyra pyra				i		x	
100	pyra				i		1 <del>571</del> 1	
101	pyra				커 너 커 커 커 커 커 커 커 너 너 너 커 커 커 커		x	
102	pyra				i			

CQ-4 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103 104 105 106 107 108	pyra pyra Per pyra pyra pyra pyra	9.0		1	거 더 너 커 쉬 너 다	cd	X	
109 110 111 112 113 114 115 116	pyra Per Avic Per pyra pyra pyra pyra pyra	 5.0		? r	44044444444444	cd cd cd	х	
117 118 119	pyra pyra pyra				i i		x	
120 121	pyra pyra				i		x	
122 123	pyra pyra						x	
124	unid			_	0	cu	x	
125	Avic			?	0	cd	X	
126	pyra				i		X	
127	pyra				i			
128	pyra				1		x	
129	pyra				1		X	
130 131	pyra pyra				11111111111			
132	pyra				1			
133 134	pyra						x	
135	pyra pyra				4			
136	pyra				i		x	
137	pyra							
138	pyra				i		x	
139	unid						X	
140	pyra	200		_	i		X	
141	Per	14.0		ı	i	cd		
142 143	Per	12.0		r	ļ	cd		
144	pyra pyra				<u> </u>			
145	pyra				÷		X	
146	pyra				i			
147 148	pyra				i			
148	pyra				i			
149	pyra				i			
150 151 152	pyra				i		X	
151	pyra				1			
153	pyra pyra				0 서서워서서시시시시시시		x	
ررا	Mira						A.	

CQ-4 Cont.

#	name	size	art	val	pres	orient	frag	epiz
154 155 156 157	Per Per pyra pyra	14.0 10.0		r 1	i i i	cd cd		
158 159 160	pyra pyra pyra				111111111		x	
161	pyra				i	100	X	
162	Per			?	i	cd		
163 164	Per	8.0 14.0		1 r	i 0	cd cd		7.5
165	Per pyra	14.0		r		Ga		x
166	pyra				1111111111		x	
167	pyra				i		2550	
168	pyra				i			
169	pyra				i			
170	pyra				i			
171	pyra			-				
172 173	ost ost			1 r	0	cd cu		
174	pyra			T		Cu	x	
175	pyra				i		x	
176	pyra				i		W-2 C-16	
177	pyra				i		x	
178	pyra				11111111		x	
179	pyra			12	i	2		
180	Per	10.0		?		cd		
181	ost			r	0 i i	cd		
182 183	pyra				1 2			
184	pyra				4			
185	pyra pyra				i		x	
186	ost			r	0	cu		
187	ost			r	0	ca		
188	unid			_	0	ou	x	x
189	pyra				i		X	
190	pyra	¥.			i			



#	name	size	art	val	pres	orient	frag	epiz
1 2 3 4	pyra pyra ost pyra			r	i i o i	cd	x x	
234567890	pyra ost ost pyra			?	0 0 1	cu		
9 10 11 12	ost pyra pyra pyra			1	0 i i	cd	х	
13 14 15 16 17	Per Per pyra pyra pyra	9.0 10.0		r	0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	cd cd	35	
18 19	Per pyra	7.0		1	4	cd	nosta.	
20 21 22	pyra Per ost	5.0		l r		cd cu	X	
23 24 25 26 27	pyra Per pyra pyra pyra	6.0		1	0 111111111	cu	x	
28 29 30 31 32	pyra pyra Per ost	6.0		r 1	i i o	cd	x	
33	pyra pyra pyra				0 1 1 1 1		х	
36 37 38	ost Avic pyra			?	0	cd		
39 40 4 <b>1</b>	Per pyra pyra	8.0		ı	i i	cd	x	
42 43	pyra pyra pyra				i i		x	
34536789901 4434456789901	pyra pyra pyra				O 마시에 마시아시 아시아시 아시아시아시아		x	
48	pyra				<u>i</u>		x	
50	pyra pyra				i		x	
51	pyra				i		x	

#	name	size	art	val	pres	orient	frag	epiz
1234567890	pyra pyra ost pyra pyra			r	i i o i i	cđ	x x	
7678	ost ost			?	001	cq		
9 10 11	pyra ost pyra pyra			ı		cd		
12	pyra Per Per	9.0 10.0		r	0 4 4 4 4 4 4 4 4 4 4 4 4 4	cd cd	x	
13 14 15 16 17 18	pyra pyra pyra Per	7.0		1	i i i	cd		
19 20 21	pyra pyra Per	5.0		1		cd	x	
22 23 24	ost pyra Pe <b>r</b>	6.0		r	o i i	cu	х	
23 24 25 26 27 28	pyra pyra pyra				0 개寸거러러거러		x	
29 30 31	pyra Per ost	6.0		r	i 0	cd cu	x	
32 33 34 35	pyra pyra pyra pyra				0 1 1 1 1 1		х х	
356 788 90 44 24 34 54 54 54 55 51	ost Avic pyra			?	0	cd		
39 40 4 <b>1</b>	Per pyra pyra	8.0		1	<u>i</u> i i	ed	x	
42 43 44	pyra pyra pyra				i i i		x	
45 46 47	pyra pyra pyra				이 매 하 더 더 때 때 때 때 때 에 에 해 해 해 하		x	
48 49 50	pyra pyra pyra				i i		x x	
51	pyra				i		x	

CQ-5 Cont.

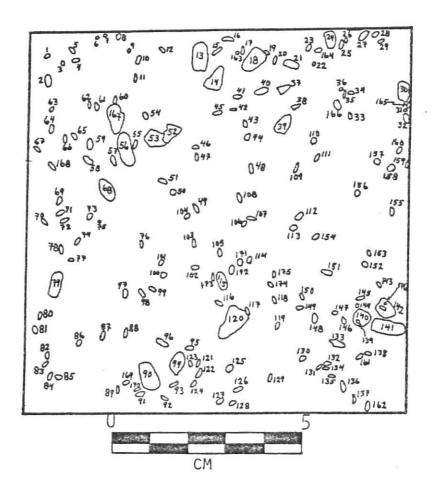
52	#	name	size	art	val	pres	orient	frag	epiz
60 pyra	53	Per pyra				i i i			
60 pyra	55 56 57	Per	10.0	ы	ı	i i	cd	X	
60 pyra	58 59	pyra				i			
75 ost 1 o cu	60	pyra				i		x	
75 ost 1 o cu	62	pyra				i		x	
75 ost 1 o cu	63 64					i i			
75 ost 1 o cu	65 66	pyra				i			
75 ost 1 o cu	67	pyra	7-0		7	i	55	x	
75 ost 1 o cu	69	pyra	1.0			i	- Cu		
75 ost 1 o cu	71	pyra				i			
75 ost 1 o cu	73					i		x	
76 pyra i x 77 pyra i 78 pyra i x 79 Per 8.0 l i cd 80 pyra i	74	pyra			1		cu	X	
78 pyra i x 79 Per 8.0 l i cd 80 pyra i	76	pyra				<u>i</u>		x	
80 pyra i	78	pyra	2 0		-	i	- 3	x	
	80	pyra	0.0		Τ.	i	ea.		
81 pyra i x 82 pyra i x	82					i			
83 pyra i x 84 pyra i	83 81 <sub>1</sub>	Fig. 1				i		X	
	85	pyra							
87 pyra i x 88 pyra i x	87	pyra				i		X	
88 pyra i x 89 pyra i x	89	pyra				i			
86 pyra i	91		8.0		1	i	cd		
92 pyra i x 93 pyra i x	92 93					i i			
	94	Per	7.0		1		cd		
96 pyra i x	96	pyra				- -		x	
94 Per 7.0 l o cd 95 pyra i x 96 pyra i x 97 pyra i x 98 pyra i 99 pyra i 100 pyra i 101 pyra i x 102 pyra i x	98	pyra				i		Α.	
99 pyra i 100 pyra i	100	pyra				i			
101 pyra i x 102 pyra i x						i			

CQ-5 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103 104 105 106 107 108 109	pyra pyra pyra pyra pyra pyra pyra pyra				선생선선생생 하는 생생선생생선 시작 시생선 시생선 시생선 시생선 시생선 시생선 시생선 시생선		x x x	
111 112 113	pyra pyra pyra				i i		x x	
11½ 115 116	pyra Per	6.0		1	1 *1 *1 *	cd		
117 118	pyra pyra pyra				i		x x	
119 120 121	pyra Per pyra	11.0		1	1111	cd	x	
122 123 124	pyra pyra pyra				i		x	
125 126	pyra pyra				i		x	
127 128 129	pyra pyra pyra				i i		x x	
130 131 132	pyra pyra pyra				i i		x	
133 134 135	unid pyra				0 i i	cu	x	
136 137	pyra pyra pyra				ન જન જન જન		x x	
138 139 140	pyra ost Per			1 r		cu	x	
141 142 143	Per Per pyra	9.0 7.0		r	0 1 1 1 1	cd cd		
144 145 146	ost pyra pyra			ı		ed.	x	
147 148	pyra pyra				i i		x	
149 150 151 152 153	pyra pyra pyra				0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		х	
152 153	pyra pyra				i i		X	

CQ-5 Cont.

#	name	size	art	val	pres	orient	frag	epiz
154 155	pyra pyra				i		x	
155 156	pyra				i		x	
157 158	pyra pyra				111111111111		x	
159	pyra				i		x	
160	pyra				i		x	
161	pyra				i		x	
162	pyra						x	
163	ost			?	0	cd		
164	pyra				i		x	
165	pyra				i			
166	pyra	40.0		-	i	- 1	X	
167	Per	10.0		l	1	cd		
168	pyra				j		X	
169	pyra				ļ			
170	pyra				1			
171	pyra				1			
172 173	pyra				1		30	
174	pyra						x	
175	pyra pyra				0 서거귀러서거러서러러			
176	pyra				3		x	
110	52 T C				-		4	



$\frac{\mu}{\pi}$	name	size	art	val	pres	orient	frag	epiz
1	unid				i			
	unid				0		x	20
2	Per	13.0		r	4	cd	25	X
23456789	Per	12.0		r	i	cd		x
7	Per	17.0		i	i	cd		A.
6	unid	11.0		-4-	0	Cu		*se
7				1	i	43		X
Ó	Per	11. 0				cd		
0	Per	14.0		I.	•	cd		2000
9	Per	18.0		r	i	cd		X
10	Per			1	i	cd		
11	Per	8.0		r	i	cd		
12	unid				0		X	
13 14	unid			-	0			X
14	Avic	CASSA PRIVI		?	0	cđ		
15 16	Per	10.0		?	i	<b>c</b> d		
16	Per	am 140			0	cd		
17 18	Per	10.0		2*	i	cd		
18	alfr							
19	Per	16.0		r	0	cd		
20	Sept			1	0	cd		
21	Per	11.0		1 1 1	i	cd		$\propto$
22	Per	11.0		1	1	cd		
23	unid				0			X
23 24	Avic			r	0	cd		
25	Per	13.0		r	0	cd		
26	alfr	. 500		-	•	- Cu		
27	Per	17.0		79	i	cd		
28	Avic	1100		r? l	0	cd		
29	Per	11.0		4	0	cd		
30	Per	11.0	59	T.	i			
	unid	11.0		7.		cd		
31				2	0	i		
32	Per	10.0		? 1 ?	0	i		
33	Per	12.0		7	i	cd		
34	Per	9.0		ţ	i.	cd		
35 36 37 38	unid				0		x	
30	unid			_	0		X	
37	Avic			?	0	i		
38	Avic	121 920 1221		1	0	cd		
39 110 141 142 143 144	Per	12.0		r	0	cu		
110	unid				0		X	
41	Avic	53		r	0	cd		X
42	Per	6.0		r	0 i	cd		
43	Per	10.0			1	cđ		
11/1	Avic			r r ?	0	cd		
45	Avic					cd		x
46	Per	19.0		r	o i	cd	475	x
47	unid	\$556						X
46 47 48	Per	7.0		r	o i	cd		200 <u>000</u>
49	unid	A #400		(=)	0	1700701	X	
50	Per	8.0		r	o i	cd	40%50.4	
51	Avic	17(RA ) <del>3</del> (4 (R))		r ?	0	cd	X	
				1075	-	-		

CR-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
5234567890	unid Per unid	6.0		1	0 1 0	cu cd	x	
55	Per unid	10.0		1	i	cd		
57	Per	12.0		1	o i	cd		
59	Per	16.0 10.0		1 ? 1	o i	cd		x
61	Sept Avic			r	0	cd cd		
62 63 64	Per unid	7.0		r	i 0	od od		
64 65 66	Per unid			r	i e	cd		
67	unid unid				0		x	х
68 <b>69</b>	Avic ost			? r	0	cd cd	X	
70 71	Per Per	12.0 15.0		l r	i i	cd cd		ж
72	Per Avic	10.0		r ?	i	cd	x	
73 74 75 76	Per Per	9.0		l r	i	cd cd		
76 77	Per Per	9.0 9.0		ī	i i 0	cd cd		
78	Per Per	17.0 6.0 5.0		1 ? ?	i	cd cd		
79 80 81	Per	10.0		l		cd		
82	Avic Per	10.0		r	o i	cd cd		
83 84	alfr unid				0			x
85 86	unid Per			1	0 <u>i</u>	cd	K	
87 88	Per Per	13.0 13.0		r	i i i	ed ed		
89 90	Avic Avic			r 1 1	0	ed ed	x	
9 <b>1</b> 92	Per Per	13.0		1 r	o i	cd cd		
93 94 95 96 97	Avic Avic			? ?	0	ed ed	x	
95 96	unid Avic			ı	0	cd	x	
97 98	Per Per	12.0			i i	ed ed		X
99 100	Avic Per	10.0		r ? 1 1	o i	cd cd		
101	Per Avic	13.0		1 ?	i	ed ed		x x

CR-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103	Per	7.0		r	i	cd		
104	Per	18.0		1 -	0	cd		x
105	Avic			r 1 2	0	cd		x
106	Avic			?	0	cd		12.2
107	unid				0		x	x
108	Per			?	0	cd		x
109	Per	18.0		?	i	cd		
110	Per	7.0		r	i	cd		
111	Per	9.0		?	i	cd		
112	Per	19.0		r	i	cd		
113	Per	35.0		r	i	cd		x
114	unid				0			
115	unid				0			
116	Per			r	i	cd		
117	Per	6.0	X		0			
118		fossil				-		
119	Avic	20.0		?	0	cd		X
120	Per	20.0		?	i	cd	1000000	
121	Avic			7	0	cd	X	
122	unid				0			nes
123 124	unid				0		X .	x
125	unid unid				0	i	X	x
126	unid				0		x	x
127	Per	22.0		71	i	cd		X
128	Avic	22.0		r ?	0	cd	x	Δ.
129	alfr			•	0	CC	77	
130	Avic			70	0	cd		
131	Per	10.0		r ?	i	cd		
132	Per	11.0		1	i	cd		
133	unid	31 5 5 5		( <del>1.5</del> )	0	3.55		
134	Per	11.0		?	0	cd		
135	Per			?	i	cd		
136	pyra				i			
137	Avic			?	0	cd	x	
138	unid				0			
139	unid				0			
140	Avic			?	0	cd		x
141	Per	12.0		1	i	cd		
142	alfr							
143	alfr							
144	Per	27.0		r	j	ed		x
145	Per	0 0		r	1	ed	X	X
146	Per	8.0 8.0		r ? ?	1 1 1 1 i	cd		
147 148	Per Per	5.0		i	-	ed ed		
149	Avic	2.0		r	0	cd		
150	unid			-	0	CG		
151	Per	8.0		r	i	cd		
152	unid	0.0			0	-	x	x
153	Per	7.0		1	i	cd	10.70°	
		100						

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CR-1 Cont.

11		-•	a b	7			2	•
$\frac{H}{H}$	name	size	art	val	pres	orient	frag	epiz
154	unid				0			X
155 156	Per	7.0		r	i	cd		
157	Per Per	12.0 14.0		r 1	1	cd cd		x
158 159	Per	9.0		r	러 너 너 너 너 너 너 너 너 너 너	cd		
159	Per	10.0		r	i	cd		x
160 161	Per Per	7.0		7	i	cd cd	X	
162	Per			r ? ! ?	i	cd		
163	Per	12.0		1	i	cd		
164	pyra				į			
165 166	pyra Per	7.0		r	1	cđ		
167	pyra	7.0		-	i	ou.		
168	Per	6.0		r	i	cd		
169	Per	12.0		1 ?	o	cd		
170 171	Per unid	11.0			0	cđ		
172	unid				o			x
173	Avic			?	0			x
174	Per	6.0		r	i	cd	1337	
175 176	Sept unid			r	i	cd	x	
177	Per	11.0		r	i	cd	45	
178	Per	9.0		1		cd		
179 180	unid Per	7.0			0 1	2.5		
181	Avic	7.0		r ?	0	cd i		x
182	Per			?	i	cd		
183	Per	14.0		r	0	cu		
184 185	Not a Avic	fossil		?	0	cd	75	
186	Per	9.0		i	0	cd	x	
187	unid	35, 30			0		x	
188	Avic			1	0	cd		x
189 190	unid unid				0	i		
191	unid				0	-	x	
192	unid				0			
193 194	Sept Avic			r	i	cd cd		3.0
195	unid			1	0	CCL	x	32
196	unid				0		x	
197	Avic			?	0	cd	X	
198 199	Per Avic	9.0		r ?	i	cd cd		~
200	Per	10.0		i	0	ou		x
201	Per	7.0		l l	i	cd		
202	Per	6.0		r	i	cd		
203 204	Per Per	11.0 7.0		i	i	cd cd		
	4 04	1.00		-44	•	Cu		

CR-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
205 206 207 208	Per Per Avic Avic	6.0 6.0		1 ? ?	i 0 0	cd cd cd		x
209 210 211	unid unid pyra				0 0 i		Ж	x
212 213 214	unid Per Avic	12.0		1	0 i 0	ed ed		x
215 216 217	Per Per Pseu	13.0 7.0		1	0 i 0	cu cd cd		
218 219 220	Per Per Per	18.0 10.0		r ? 1	i	cd cu cd		x
221 222 223	Per Per Per	9.0	x	r ? r	i i i	cd cd i		
22h 225 226	unid unid Per	8.0		r	0 0 i	i i cd	₹	
227 228 229	Per Per Per	10.0		î r r	0 i	ed ed		
230 231 232	Per Per unid	7.0 7.0		r	i 0 0	cd cd		
233 234 235	Per Per Per	10.0		1 ?	i o	cd i cd		
236 237 238	Per pyra Per	6.0 7.0		r ? r	i i i	cd cd		
239 240 241	Per pyra ost	38.0		r	i i	cd cd		x
242 243 214	pyra ost ost			l r	i 0 0	ed ed		
245 246 247	pyra Per Per	12.0 9.0		r r	i i o	ed eu	ж	
248 249	ost burr ost	, • •	x	r	0	ed		
250 251 252 253 254 255	Per ost ost	9.0		r 1 1 2	i 0	od od ou		
254 255	Avic Per	7.0		ī ?	o i	cd cd		X

CR-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
256 257	Avic Per	11.0		r 1	o i	ed ed		
258 259	Per Per	10.0 13.0		r	i	cd cd		
260 261	Per Per	5.0 26.0		r	i	cd cd		ne:
262	ost	20.0		r	0	cu		x
263	ost			r	0	cu		
264 265	ost Per	6.0		r ?	o	ed ed		
266	Per	7.0			i i	cd		
267	Per	-		?	i	cd		x
268 269	ost Per	13.0		r ?	0	ed eu		
270	ost	15.0			0	cd		
271	Per	10.0		r	i	cd		
272	Per		x		0	i		
273 274	Per Per	9.0 8.0		r	i i	cd cd		
275	Per	13.0		r	i	cd	S.	
276	Per	16.0		1	0	cu	5	
277	Per	15.0		1	i	cd		
278 279	Per unid	11.0		r	i	cd cu		
280	Per	14.0		r	. 0	cđ		x
281	Per	12.0		r 1 1	i	cd	840	
282 283	Per	4.0		1	i	cd		
284	unid unid				0			
285	Per	10.0		r ?	i	cd		
286	Per	8.0			i	cd		
287 288	Avic pyra			r	o i	cd		
289	Per	12.0		r	i	cd		
290	Per	10.0		r	i	cd		
291 292	Per Per	12.0 9.0		1 1 ?	i i i i	cd cd		
293	Per	7.0			i	cd		
294	Per	7.0		1	i	cd		
295 296	Per	11.0		1"		cd		
297	unid Avic			r	0	cd		x
298	pyra			_	i	-		
299	pyra	0 0		_	i			
300 301	Per Per	8.0 7.0		1	i	cd cd		32
302	Per	7.0		i	0 1 1 1 1 1 1 1 1	cd		X
303	Per	-			i	cd		
30T	Per			r ? 1	i	cd		
305 306	Per Per	9.0 8.0		?	i	cd cd		X
100000000000000000000000000000000000000	CE . 15175	등 등 등			2003	4 <b>7</b> 44517		

CR-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
307 308 309 310 311 312 313	Avic Per unid Per Avic Avic unid Per	6.0 4.0		r 1 r ?	0 1 0 1 0 0 0 1	cd cd cd cd cd	x	х
315 316 317 318 319 320 321 322 323 324	unid Per Per unid Per Per Per Per Per Per	5.0	x	? ? ? 1	0 1 0 0 1 1 0 1 1 1	cd cu cd cd cu i	x	

ć

#	name	size	art	val	pres	orient	frag	epiz
1 2 3 4 5 6 7 8 9 0 10	Avic Per Per Per	14.0 11.0		? 1 r 1	0 1 1	cd cd cd cd	x	x
6	unid				0		x	x
8	Per Avic			1 ? 1	0	cd cd		x
9	Per Per	10.0 10.0		1	i i	cd cd		x
11	Per	7.0			444444	cd		
12 13 14	Per Per			?	ı i	cd cd	x	
14	Per Per	12.0 10.0		r ? ? ?	<u>i</u>	cd cd		
15 16	unid				0		x	x
17 18	Per Per	9.0 14.0		r 1 ? 1	에 너 너 너 너 너 너 너	cd cd		x
19 20	Per	8.0		1	i	cd		
21	Per Per	13.0			i	cd	x	
22	Per Per	12.0 15.0		r 1 ?	i	cd cd	121	
23 24	Per	12.0		?	1	cd	3	
25 26	Per unid			?	i o	cd cu	x	x
27 28	Avic			?	0	cd	x	
29	Avic Per				i i	ed ed		x
30 31	Avic Per	12.0		r ? 1 ?	o i	cd cd		
31 32	Per	11.0		1	0	cu		x
33	Avic Per	14.0		r	o i	cd cd	х	X
35	Per	9.0		1	i i i	cd	ж	
33 34 35 36 37 38	pyra pyra			_			Α.	
38 39	ost Avic			1 ?	0	ed eđ	x	
39 40	Per	16.0		r ?	0	cd	x	
41 42	Avic Sept			r	0	cd cd		X
43 44 45 46 47 48	Per Per	13.0		r ? r ?	o i	cd cd	x	
45	Per				i	cd	x	
40 47	Avic Per	14.0		r	0	cd cd		x
48 49	Avic Avic			r ?	0	cd cd		
50 51	ost			r 1	0	cd		x
51	ost			l	0	cd		

CR-1+ Cont.

#	name	size	art	val	pres	crient	frag	epiz
52	pyra				i			
53	Avic			l	0	cu	x	
51.	Avic			r	0	cd	. <del>11 - 1</del> - 1	
55 56 57	Per	10.0		i	i	cd		
20	Per	7.0		i	i			
20					i	cd		1999
21	Per	17.0		r	1	cd		x
58	Per	20.0		r	i	cd		
59	Per	12.0		1	i	cd		
60	Avic	0 -		?	0 i i	cd	X	
61	Per	8.0		r	i	cd		
62	Per	9.0		1	i	cd		
63	Per	11.0		r	i	cd		
64	Sept			r	0	cd		
65	Per	14.0		r	0	cu		
66	Avic			?	0	cd	X	
67	unid				0		x	
68	Per	8.0		1	i	cd		
69	Per			1 ?	i	cd	x	
70	unid				0	-	x	x
71	Per	10.0		r	i	cd	32	
72	Per	20.0		ī	i	ed		x
73	ost	20.0		r	0	cd	9	
74	Per	11.0		1	i,	cd		
75	Per	9.0		1	i			
76		18.0		1		cd		
70	Per			r ?	i	cd		X
77	Per	14.0			0	cu		
78	Per	15.0		r	i	cd		X
79	Per	14.0		r	i	cd		
80	unid				0	12	х	X
81	Per	12.0		r	i	cd		
82	Avic			?	0	cd	x	X
83	Per	10.0		r	0	cd		
84	Per	7.0		1 ?	i i	cd		
85	Per	3.0		?	i	cd		
86	Per	17.0		r	0	cd		
87	Avic			r	0	cd		x
88	Per	16.0		1	i	cd		
89	Per	14.0		r	1111111111	cd		
90	Per			r ?	i	cd	x	
91	Per	10.0		r	i	cd		x
92	Per	13.0		T	i	cd		
93	pyra			70	1	5.75	x	
91,	Per	13.0		7	i	cd		
95	Per			?	i	cd	x	
93 94 95 96 97	ost	3000 CMG		1 ? 1 ? r	0	cu	46	
97	Per	7.0		?	i	cd		
98	Avic	1.0		*	0	cd		
99		16.0		Τ.	4			45
77 100	Per			1,	i	cd		X
100	Per	8.0		1	1	cd		
101	Per	13.0		r 1 1	i	cd		
102	Avic			1	0	cd		

CR-1+ Cont.

#	name	size	art	val	pres	orient	frag	epiz	
103 104	Per Per	16.0		1 ?	o i	ed ed			
105	Per	22.0		r	i	cd		x	
106	Per	10.0		r	i	cd			
107	Per	13.0		?	i	cd		x	
108	Per	14.0		?	i	cd		X	
109	Avic			?	0	cd	x		
110	Avic			?	0	cu	x		
111	unid			•	0		X		
112	Per			?	i	cd			
113 114	Avic	0.0		r	0	cd			
115	Per ost	9.0		r	i	cd			
116	ost				0	cu cu		*	
117	Avic			r	0	cd		x	
118	Per	13.0		1	0	cd		x	
119	Per	6.0		ī	i	cd		2	
120	Per	17.0		1	0	cu			
121	Avic	N =		r	0	cd	x		
122	Per	13.0		r		cd			
123	Per			?	i i	cd	x		
124	Per			r ? ? 1	i	cd	x		
125	Per	6.0		ļl	i	cd			
126	Sept			1	0	cd			
127	unid				0	cu	X		
128	unid	10.0		-	٥	cu	x	X	
129 130	Per Per	10.0		1 ? ?	° i	cu			
131	Avic	0.0		?	0	cd cd	x		
132	Per			?	i	cd	x		
133	unid			-	0	04	x		
134	Per	-		?	i	cd	x		
135	unid				0		x		
136	Avic			1	0	cd		x	
137	unid				0		x		
1 38	Per	11.0		r	i i	cd			
139 140	Per	-		r??!rr!r1r1		cd	x		
140	Avic	40.0		?	0	cd	X		
141	Per	13.0		1	i	cd			
142	Per Per	12.0 6.0		r	0 i i	cd	x		
143 144	Per	9.0		r	-	cd cd			
145	Avic	7.0		1	0	cd			
146	Per	16.0		7	0	OIT OCT			
147	Per	13.0		ī	0	cd			
148	Per	8.0		1	0	cd			
11.9	Per	12.0		r		cd			
150	Per	6.0		?	i i i	cd			
151	Per	5.0		r		cd			
150 151 152 153	Avic	2 2		r ? r ?	0	cu	x		
153	Per	9.0		?	i	cd			

CR-1+ Cont.

#	name	size	art	val	pres	orient	frag	epiz
154	Per	18.0		r	0	cu	x	
155	Per	12.0		1	•	cu	X	
156	Per	9.0		1 1 ?	i	cd		
157 158	Avic Per	10.0		i	o i	cd cd	x	
159	Per	18.0			0	cu		x
160	Per	8.0		r 1	i	cd		•••
161	Per	8.0			i i	cd		
162	Per	8.0		r l	i	cd		
163	Per			?	i	cd	x	
164	unid				0		x	
165	Per	18.0		r	0	cd		x
166	unid	0 5			0	cu	X	
167	Per	8.0		?	i	cd		
168 169	Per	11.0		l	i i	cd		
170	pyra Per	11.0		1	i	cd		x
171	Per	14.0		ī	o	cu		A-
172	Per	9.0		ī	i	cd		
173	Per	16.0		1 1 ?	i	cd		x
174	Per			?	i	cd	х .	
175	Per	12.0		1 ?	i	cđ		
176	Per			?	i i i	cd	x	
177	Bell		**	•	i			
178 179	Per Per	10.0		?	i	ed	X	
180	Avic	10.0		1 1 1	0	cd cu		
181	Per	18.0		ī	4	cd		
182	Per	9.0		r	i i	cd		
183	Per	15.0		ī	i	cd		
184	Per	9.0		1 ? ?	i	cd		
185	Avic			?	0	cd	x	x
186	Per				<u>i</u>	cd		
187	Per	11.0		r	i	cd		
188	Sept			r	i	cđ		
189 190	Avic Per	11.0		<u> </u>	o i	cd		
191	Avic	11.0		r 1 1 ?	0	ed ed		
192	Avic			?	0	cd		
193	Per	13.0		i	i	cd		x
194	Per	8.0		r	i	cd		32
195	unid				0		x	x
196	unid				0		x	
197	unid				0		x	x
198	Per	11.0		?	i	cd		X
199	ost			r	0	cu		
200	ost				0	cd		
201 202	ost ost			r	0	cd cu		
203	ost			r 1	0	cu		
204	ost			ī	0	cu		

CR-1+ Cont.

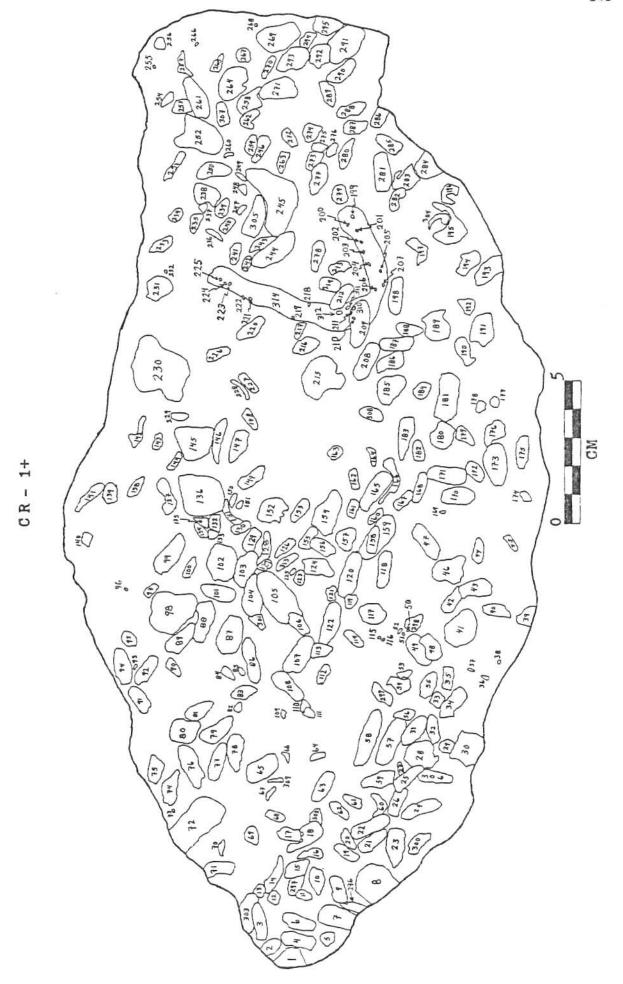
#	name	size	art	val	pres	orient	frag	epiz
205	ost			1 ?	0	cu		
206	ost				0	cd		
207	ost	4 C O		r	0	cu		
208 209	Per	15.0		1	0	cu	-	
210	unid ost			7	0	cd	x	
211	ost			1	0	cu		
212	Per	12.0		r	i	cd		
213	Per	6.0		?	0	cu		
214	Per	10.0		r	i	cd		
215	Avic			r	0	cd		x
216	unid	+			0		x	
217	Per			?	i	cd		
218	ost			l	0	cu		
219	ost			1	0	cd		
220	Per	11.0		1	i	cd		
221	ost			r	0	cd		
222	ost			r	0	ca		
223	ost			1	0	cd		
224	ost			1	0	cd		
225	ost	40.0		ı	•	on	9	
226 227	Per	12.0		1 ?	i	cd		
228	Per unid			•		cd		
229	Per	9.0		?	o i	cu cd	X	
230	Avic	7.0		?	0	cd		
231	Avic			r	o	cd		x
232	ost			ī	0	cd		43
233	Per	11.0		ī	i	cd		x
234	Per	9.0		r	1	cd		7.5
235	Per	9.0		r	i	cđ		
236	unid				0		x	x
237	Per			1	i	cd		
238	Avic			1	0	cd		
239	Per	8.0		r	i	cd		
240	Per	7.0			i	cd		
241	Per	9.0		r 1	1 1 1 1	cd		
242	Per	7.0		Ţ	i	cd		
243 244	Per	10.0		1	i	cd		
244	Per Psue	18.0		1	i	cd		
246	Per	8.0		r l	0	cd cd		
247	Per	5.0		r	i	cd		
248	Per	5.0		i	i	cd		
249	unid	7.0		-	0	~~ <u>~</u>	x	
250	Per	12.0		1	i	ed	-14	
250 251 253 253 250	Per	13.0		ī	0	ou		
252	unid				0	5000/501		x
253	Per	7.0		1	i	cd		
254	Per	7.0		1	1	cd		
255	ost			1	0	ed		

CR-1+ Cont.

name	size	art	val	pres	orient	frag	epiz
unid			2	0		x	x
						x	
	18.0		r				X
			?			x	
	44.5						
	13.0						
	-		r		cu	X	x
	( 0		221		•		x
	6.0		r				
			Τ.		cu		
					nara:		
			r				
	0.0		-				
							x
	0.0		ŗ		ca		X
			2		. 3		
			2	9		X	
	12.0		,				
	15.0		2				
			?				
	15 0						
		8					
	7.0		_		Ca		x
			2		60	~	_
	11.0		i			~	
			1				
			ī	i			
				i			
			?				
	12.0			7			
			?				
			?				
			?				
			?				
			?				
						x	x
							x
							x
	9.0		r	i	cd		
			1	i			X
Per			?	0			
Per	10.0		r	0	cu		
unid				0	cu	x	
Avic				0	cd	x	
Per			l	0	cu		
Per	14.0		1	i	cd		
	unid Avic Per Avic Per unid ost do unid Per Per Bell avic Per Per Bell Avic Per Per Avic Per Avic Per unid unid unid Per Per unid Avic Per Uni	unid Avic Per 18.0 Avic Avic Per 13.0 Pseu Per unid Per 6.0 ost unid Per 9.0 Per 8.0 Per 8.0 Per 8.0 Per 15.0 Per 9.0 Rell Avic Per 15.0 Per 9.0 unid Avic Per 15.0 Per 9.0 unid Avic Per 15.0	unid Avic Per 18.0 Avic Avic Per 13.0 Pseu Per unid Per 6.0 ost unid ost unid Per 9.0 Per 8.0 Per 8.0 Per 8.0 Per 8.0 Per 9.0 unid Avic Per Per 15.0 Per 9.0 unid Avic Per 15.0 Per 9.0 unid Avic Per 12.0 Avic Per 12.0 Avic Per 12.0 Avic Per 10.0 Per 9.0 Indid Unid Unid Unid Unid Unid Unid Ver 9.0 Per 10.0 Per 10.0 Indid Avic Per 10.0	unid Avic Per 18.0 r Avic ? Per 13.0 r Pseu r Per 13.0 r Pseu r Per r unid Per 6.0 r ost unid ost r unid Per 9.0 r Per 8.0 r Per 8.0 r Per 8.0 r Per 15.0 r Per 8.0 r Per 9.0 r Per 9.0 r Per 9.0 r Per 15.0 r Per 9.0 r Per 15.0 r Per 15.0 r Per ? Per 13.0 r Per ? Per 15.0 r Per 9.0 r unid Avic ? Per 11.0 r Avic ? Per 12.0 r Avic ? Per 10.0 r Per 10.0 r Unid Avic ? Per 10.0 r	unid Avic Per 18.0 r Avic Per 13.0 r Per 13.0 r Per 15.0 r Per 8.0 r Per 15.0 r Per 13.0 r Per 15.0 r Per 1	unid Avic Per 18.0	Unid Avic

CR-1+ Cont.

#	name	size	art	val	pres	orient	frag	epiz
307 308 309 310 311 312 313 314	Per Per Avic pyra pyra pyra Per burr	8.0 13.0		1 ? ?	i 0 0 i i i i	cd cu cu	x	



#	name	size	art	val	pres	orient	frag	epiz
12345678	Per Per Per	13.0  5.0	x	?	i i	cd cd cd		2
4	Not a Per			r	1	cd		
6	Per	7.0		r	- -	cd		
7	Per	11.0		ī	i	cd		
8	Per	7.0			************	cd		
9	Per			r ?	i	cd		
10	Per	11.0		?	i	cd	x	
11	Per	6.0		r	i	cd		
12	Per	15.0		1	i	cd		
13	Per	10.0		1	i	cd		
14	pyra			<u> </u>	i	_		
15 16	Per	13.0		1	i	cd		
16	Per	10.0		1	i	cd		
17	pyra	44 0		-	i	2222 <b>9</b>		
18	Per	11.0		1	ï	ed		
19 20	Per Per	12.0		r	i	cd		
21	unid	17.0		r	0	cd cu	~*	x
22	Per	9.0		r	i	cd	x	
23	Per	11.0		r	i	cd		
24	Per	12.0		ī	0	cd		x
25	Per	12.0		r	i	cd		
26	Per	8.0			i	cd		
27	Per	9.0		?	i	cd		
28	Per			r ? ? ? ? ?	ì	cď		
29	Per			1	0	cu		x
30	Acan			?	0	cu		
31	Avic	20.02		?	0	cd		
32	Per	5.0		1	i	cd		
33	Sept			?	0	cd	x	
34	Acan	44 0		1	0	cd	x	
35	Per	11.0		r	<u>,</u>	cd		
27	Per Per	8.0 13.0		r	i i i	cd cd		25
38	unid	15.0		22	0	Cu	x	x
30	Avic			T*		cđ		
10	Per	14.0		r 1 1 r	o i i	cd		
41	Per	10.0		ī	i	cd		
1,2	Per	15.0		r	i	cd		
43	Avic	57		2		cd		
1,1,	Per	4.0		r	o i i	cd		
45	Per	12.0		r		cd		
46	Per	13.0		1	0	car.		
47	Per	13.0		r	i	cđ		x
48	Per	16.0		1	i	cd		
49	Per	15.0		1	i	cd		
3567899442345678901	Per	12.0 13.0 16.0 15.0 13.0		41 411 41	0 i i i	cd		
וכ	Per	11.0		7	1	cd		

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
52	Per	16.0		1	i	cd		
53	pyra				i i			
53 55 55 55 57 58	Per	13.0		r ?		cd	828	
22	Avic Per	10.0		í	0	cđ cd	x	
57	Per	8.0			<u>.</u>	cd		
58	Per			r ? 1 1 ?	i	cd		
59	Per	16.0		1	i	cd		
59 60	Per	13.0		l	i	cd		
61	Per	15.0			ì	cd		
62	Per	12.0		r	i	cd		
63	pyra			?	i i i i i i i i i i i i i i i i i i i	2.3		
64	Per	~~			1	cd	x	
65 66	pyra Per	12.0		1	÷	cd	Х	
67	Per	12.0			i	cd		
68	Per	11.0		r 1	i	cd		
69	pyra				i i			
70	Per	15.0		1	i i	cd		
71 72	Per	7.0		r	i	cd		
72	Per Avic	8.0		r 1	i	cd		
73 74	Per	8.0		i	o	cd cd		
75	Per	18.0		r	0	cd		
76	unid	100000		ā	i	30 <del>0</del>		
76 77 78	Per	13.0		r	0	cd		
78	Per	11.0		r	i	cd		
79	Avic			1	0	cd		
80	Avic			1 1 ?	0	cd	-32	
81 82	Sept Per	13.0		r	o i	cd cd	x	
83	unid	15.0		1		Ca		
84	Per	10.0		?	o i	cd		
85 86	unid	N73			0			
86	Per	12.0		r	i	cđ		
87	Per	11.0		r 1 1	i i i	cd		
88	Per	6.0		1		cd		
89 90	Per Per	15.0		r	0	cu		
91	Per	5.0	x	Ţ	i i	cd		
92	Not a	fossil			<del>77.</del>	04		
92 93 94	Per	13.0		r	i	cd		
94	Per	11.0		1	i i	cd		
95 96	Per	13.0		1		cd		
96	Acan	41 0		γ.	0	cu	x	
97 98	Per Per	14.0 14.0		r11?1r11r1	0 i i	cd cd		
99	Per	15.0		ī	i	cd		
100	Avic			ī	0	cd		
101	Per	11.0		r	i	cd		
102	Per	10.0		ı	i	cd		

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103 104 105 106 107	Per Per Per Avic	13.0 9.0 16.0		r r ? ?	i i o o	cd cd cd cu cu	x x	
108 109 110 111 112 113	Avic Per Per Avic Per pyra	12.0 11.0 9.0		r r r	o i o i	cd cd cd cd		x x
114 115 116 117 118	Bell Avic pyra Per Per	17.0 12.0		? r r	i i o i i i i i i	cd cd		
119 120 121 122 123	Per Per Sept Avic pyra	8.0 12.0		1 1 1 ?	0	cd cd cd		x
124 125 126 127 128 129	pyra pyra unid Per Per Per	8.0 7.0 14.0		1	0 i i i i i	cd cd cd	x	x
130 131 132 133 134	Per Per Avic unid pyra	11.0		r 1 r ?	i i o o	cd cd cu	x x	x
135 136 137 138 139 140	unid Per Per Per unid unid	7.0 16.0		? ? r	0 i i 0	cd cd cd	x	
141 142 143 144 145	pyra Per Per Avic	6.0 8.0		r 1 r	0 i i i 0	cd cd cd	x	x
146 147 148 149 150	Per Per Per Per unid	12.0 8.0 12.0 14.0		r ? !	i i e o	ed ed eu	x	x
151 152 153	unid Per Per	10.0 13.0		? r	0 i i	cd cd	x	

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
154 155 156 157 158 159 160 161 162 163 164	Per Per Per Per Per Per Per Per Per	8.0  11.0 10.0 14.0 14.0 14.0 6.0 fossil	x	r ? ! ! ! ! r !	: ::::::::::::::::::::::::::::::::::::	cd cd cd cd cd cd cd		x
166 167	Per Per	16.0		1	i	cd cd		
168 169 170 171 172	unid Per Avic Per Per	7.0 7.0 13.0		r r l r	o i o	cd cd cd	x	x
173 174 175 176 177 178	Per Per Per Per Per pyra	12.0 7.0 5.0 17.0		r 1 ? r	i i i i i o i	cd cd cd	N	x
179 180 181 182	Per Per pyra Avic	6.0 11.0		r 1 r	i i i	ed ed ed		
183 184		9.0 fossil		r	i	cd		
185 186 187 188	Avic Per Per Per	12.0 14.0 15.0		1 1 r	o i o i	ed eu ed		х
189 190 191 192 193 194	Avic Per Per Per Per Avic	6.0 10.0 15.0 15.0		r ? r r r	0 i i i i 0	cd cd cd cd cd	x	
195 196 197 198 199 200 201	unid Per Per Per Per Per pyra	10.0 20.0 11.0 14.0 16.0		r r r	0 i i i i i i i	cd cd cd cd		
202 203	pyra Per Bell	13.0		1	i o i	cđ	x	

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
205	Per	10.0		?	i	cd		
206	unid	720			0	cu	x	
207	Per	16.0		r	i	cd		
208	Avic			?	0	cđ	x	X
209	Bell				o i i	2		
210	Per	10.0		r		cd		
211	Avic	0.0		r	0	cd		
212	Per	8.0		1	i	cd		
213	Pseu	40.0		r ?	•	cd		
214	Per	13.0			i	cd		
215 216	Per	9.0		r 1	1	cd		
217	Per Per	12.0 13.0			i	cd cd		
218	Per	10.0		r	i	cd		
219	Bell	10.0		*	i	ca		х
220	Per	10.0		1	i	cd		
221	Per	9.0		r	i	cd		
222	Per	12.0		r	î	cd		
223	Per	13.0			i	cd		
224	Per	20.0		1	i	cd		x
225	Per	11.0	x		i	i		
226	Per	17.0		1	i	cd	19	
227	Per	14.0	x		i	cd		
228	Per	14.0		r	i	cd		
229	Per	12.0		. r	i	cd		
230	Per	4.0		?	i	cd		
231	Per	16.0		r l	i	cd		
232	Per	9.0		1	i	cd		
233	Per	19.0		1 ?	i.	cd		
234	Avic			?	0	cu	x	
235	Avic	41.0		1 ?	•	cd		x
236	Per	14.0			i i	cd		
237 238	Per Avic	11.0		r		cd		
239				7	o i	cd		
240	pyra unid				0		x	
211	Per	6.0		7	i	cđ	-2-	
241 242 243 244	Avic	0.0		1?	0	cd	x	
243	unid				0	0	x	x
244	Per	6.0		r		cd	\$2.50 1	1200
245	Per	15.0		1	i i i	cd		
246 247	Per	16.0		1	i	cđ		x
247	Per	10.0		1	0	cd		
248	Per	18.0		r 1 1 1	i	cd		x
249	Per	9.0		1	i	cd		x
250	Per	8.0			i	cd		
251	Per	10.0		r	0	cd		x
252	Per	11.0		ī	i	cd		
253	Per	13.0		r 1 ? 1		cd		
249 250 251 252 253 254 255	Per	12.0			0	cu		
455	Per	16.0		r	i	cd		X

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
256	Per	9.0		1	i i	cd		
257 258	pyra Per	9.0		1	i	cd		
259	Avic			?		cd	x	
260 261	Per Bell	10.0		1	0 i i i	cđ		
262	Per	8.0		?	i	cd		
263	Per	14.0		r 1		cd		
26 <u>1</u> 265	Per Sept	11.0		r	i o	cd cd		
266	Per	13.0		r	i	cd		
267 268	Per unid	8.0		r		cd		
269	Per	11.0	x		o	cd		x
270	Per	9.0		r	i	cd		
271 272	Per Per	16.0 13.0		r	i i	cd cd		
273	unid	13.0		•	i	Cu		
274	Avic	10.0		r	0	cd		x
275 276	Per Per	12.0 12.0		r	i i	cd cd		
277	Avic	127.50		1 ?	0	cd	20	x
278	Avic	12.0		?	0	cd		x
279 280	Per Per	13.0 11.0		l r	i i	cd cd		x
281	Per	10.0		r	i	cd		
282 283	Per	10.0		r ? ! !	i	cd		
284	Per Per	10.0		?	i	i ed		
285	Avic			1	0	cd		x
286 287	Per Per	13.0 19.0		1	i	cd		
288	unid	19.0		_	0	cd		
289	Per	15.0		1	i	cd		
290 291	Per Per	16.0		7	i	cd cd		x
292	Per	10.0		i	i	cd		
293	Per	7.0		1	i	cd		
294	Per pyra	11.0		1	i i i i i i	cd	x	
295 296	Per	9.0		<u>1</u> ?	i	cd	^	
297	Per	13.0			0	cd		x
298 299	Avic Per	10.0		r	o	cđ cd		x
300	Per	9.0			i	cd		
301	Per	8.0		r ? ?	i	cd	1229	
302 303	Avic Per	11.0		r	o	cu	x	
304	pyra			F	i i i			
305 306	pyra	12.0			i	cd		
500	Per	12.0		r	1	Ga		

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
307 308 309 310 311	Avic Avic Per pyra unid	8.0		? 1 ?	0 0 i i 0	cu cd cd		x
312 313 314 315 316	Per Per Per Per Per	16.0 19.0 7.0 8.0 13.0		1 r 1	0 i i	cd cd cd cd		x
317 318 319	pyra Per Per	9.0 12.0 12.0		1	1111111	cd cd	x	x
320 321 322 323	Per Per Per unid	12.0 9.0		r r 1	0	ed ed	x	
324 325 326 327	Per Per Per pyra	8.0  9.0		r r	i 0 i	cq ca eq	a.	
328 329 330 331	Per Per Per Per	17.0 12.0 11.0 11.0		1 r ? 1	0 i i	cd cd cd		
332 333 334 335	unid Per Per Per	7.0 12.0 13.0		r 1 r	o	cu cd cd		
336 337 338	pyra pyra Avic			?	i i i o	cd	x	x
339 340 341 342 343	Per Per Per pyra	5.0 5.0 11.0		i	0 1 1 1 1	cd cd		
344 345 346 347	unid Avic Per Per Per	15.0 8.0 14.0		? 1 r r	0 0 1	cd cd cd		x
348 349 350 351	unid Avic pyra pyra			?	0 0 1 1	cđ	x	х
352 353 354 355 356	Per Avic Per Per	12.0 11.0 11.0		r 1 ? r	i 0 i	ed ed ed		x
356 357	Avic Per	12.0		?	0	ed eu	x	

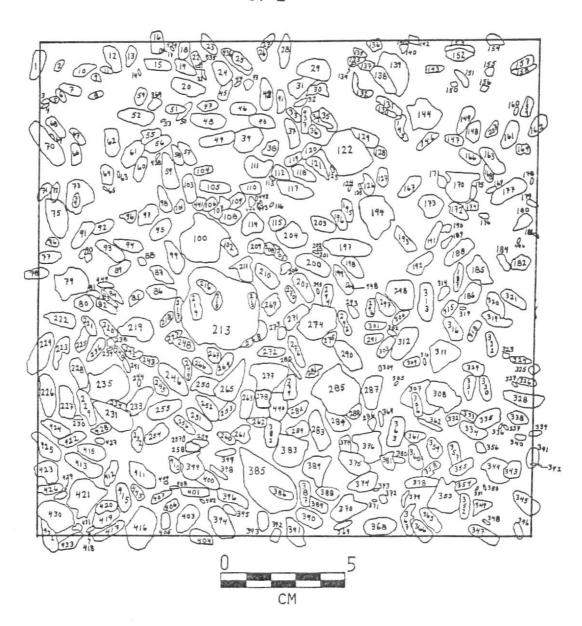
CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
358 359	Per Per	11.0		1 r	i i	cd cd		
360	Per	11.0		i	i	cd		
361	Per	11.0		ī	ī	cd		
362	Per	9.0		r	i	cd		
363	Per	13.0		ı	i	cd		
364	unid				0	cu		
365	Per	13.0		r	i	cd		
366 367	Per Avic	13.0		r ?	i	ed eu	x	
368	Per	13.0		?	i	cd	Δ.	
369	Per	9.0		ì	0	cd		x
370	Per	17.0		1	1	cd		
371	pyra				i		x	
372	Avic			?	0	cd	x	
373	pyra	41.0			i			
374	Per	14.0		r	0	cu		
375 376	Avic Avic	9.		1	0	cd	22	
377	Per	7.0		r	o	cd	x	
378	Per	17.0		r	i	cd		
379	unid	. ,		-	0	ou	R	
380	Per	6.0		ı	i	cd		
381	Avic			1	0	$\mathbf{c}\mathbf{d}$		
382	Per	9.0		r	i	cd		
383	Avic			r	0	cd		
384	Avic			1	0	cd		x
385 386	Sept	10.0		1	0	cd		
387	Per Per	12.0 11.0		r	i i	cd cd		
388	Per	9.0		r		cd		
389	Per	9.0		r	i	cd		
390	Per	12.0		r	i i	cd		
391	Per	8.0		1	i	cd		
392	Per	7.0		1	i	cd		x
393	Per	11.0		r ? ! !	11111111	cd		
394	Per	14.0		1	j	cd		
395	Per Per	8.0 13.0		;	<u>,</u>	ed		
396 397	Per	10.0		?	4	cd i		
398	Per	9.0		i	i	cd		
399	Avic	5 5 1		1		cd		X
400	Per	9.0		r 1	i	cd		
401	Per	14.0		1	ĭ	cd		
402	pyra	21. 2		2	i	9		
403	Per	14.0		1 1 ?	i	cd		
404	Per Per	12.0 9.0		2	1 1	cd		
406	Per	9.0			÷	cd cd		
407	Per	11.0		r	0 1 1 1 1 1 1 1 1 1 1 1 1	cd		
408	pyra	∵192 <b>-</b> 8070		<del></del>	i			

CT-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz	
409 410 411	pyra unid Per	13.0		r	i 0 i	cd	x		
412 413 414	Per Per Per	14.0 17.0 8.0		r r l	i : :	cd cd cd			
415	Per	15.0		r		cd		x	
416 417	Per Per	19.0 14.0		r ?	o i	cd cd		x	
418	Per	5.0		1	i	cd			
419 420	unid Avic			?	0	cd	x	x	
421 422	Avic Per	12.0		r	o i	cd cd			
423	Avic			?		cd	x		
424 425	Per Per	13.0 18.0		1	o i i	cd cd			
426 427	Per Avic	15.0		1	i o	cd cd			
428	Per	8.0		r? 1 1 1 ?? 1	i	cd		x	
429 430	Per Avic	6.0		r	0	cd cd	+1	x	
431 432	unid Per	14.0		1	O	cd	X		
433	Per	13.0		1	i i i	cd			
434 435	pyra pyra				i				
436 437	Per Per	17.0 13.0		? r	0	cd cd			
438	pyra	. 500		?	i				
439 440	Per unid					cd	x	x	
441 442	Avic pyra			?	o 1 i	cu	x		
443	pyra	6.0		,	1000				
444	Per Per	6.0 6.0		l l	i	cd			
446 447	pyra Per	15.0		r	: : : :	cd	x	х	
448	pyra unid				i				
449	anra				O			x	

## CT-1



CU-13 Area #1

#	name	size	art	val	pres	orient	frag	epiz
1234567890	Per Per pyra pyra	6.0 12.0		1	i i i	cd cd	x	
5	pyra Per	15.0		r	i i	cđ	x	
7 8 9	pyra Per pyra	7.0		1	i i	cd	x	
11 12	pyra Per Per pyra	7.0 12.0		r ?	i i i	ed ed		
13 14 15 16 17	pyra pyra pyra pyra				i i i			
18 19	Per	14.0		r	i	cd	x	
20 21	Per pyra	11.0		r	i i	cd	Α,	
22 23 24 25 26 27 28	pyra Per pyra pyra pyra	8.0		?	더러너 더러 더러 더러 이러 이러 이러 이러 이 이 이 이 이 이 이 이 이	cđ		
29 30	pyra Per pyra pyra	7.0		?	i i i	cđ		
31 32 33 34	pyra Per Avic	15.0		? r	0	cd cd		
35	Per Per	14.0 13.0		r 1	i	cd		
36 37 38	pyra pyra pyra pyra pyra				*1 *1 *1 *1		x	
41 42 43	pyra pyra pyra				i i		x	
45	pyra Per	12.0		1	i	cd		
39 44 44 44 45 44 45 45 45 45 45 45 45 45	pyra Per pyra pyra pyra pyra	10.0		r	तं भी	cd		

CU-13 Area #1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
555555555661 60 5555555555661	Per Per pyra pyra	6.0 4.0		1 r	i i i	cd cd	x	
56 57	pyra pyra				i		x	
58 59	pyra Per			?	i	ed	x	
63	pyra pyra pyra				러 너 더 더 더 더 더 더 더 더 더 더 더 더 너 너 너		x	
65 66	pyra pyra pyra				i		x	
67 68 69	Per pyra pyra	13.0		r	i i i	cd		
70 7 <b>1</b> 72	pyra pyra Per	14.0		r	i i	cd	x	
73	Per Per	6.0		?		cd cd	x	
74 75 76 77	unid Per pyra	5.0		?	o i i	cd	X	
77 78 79 80	pyra Per Per	13.0 11.0		? r	0 1 1 1 1 1 1 1 1	cd cd		
81 82 83	pyra unid unid			-	i 0 0	-	x x	
84 85 86	pyra Per	6.0		?	i i	cd	x	
86 87 88	Pe <b>r</b> ost pyra	5.0		r	i 0 i	ed ed	x	
89 90	pyra pyra	7.0		7	i		X	
91 92 93 94	Per Per Per pyra	12.0		r r	0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	ed ed ed		
93 94 95 96 97 98	pyra pyra pyra				i i		x	
99 100	pyra pyra pyra				i			
101 102	pyra pyra				i		X	

CU-13 Area #1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103 104 105 106 107 108	pyra pyra pyra pyra pyra pyra pyra				111111		x	
109 110	unid pyra				0		x	
111 112 113 114 115 116	pyra pyra pyra pyra pyra pyra				© 서러 러워 러워 너 러워		x	
117 118	pyra Per	7.0		r	i	cd	x	
119 120 121	Per pyra unid	8.0		r	i i o	cd	x	
122 123	Per pyra	14.0		1	i	cd		
124 125 126 127 128	Per pyra pyra pyra	13.0		r	111111111	cd	x	
129 130 131 132	pyra pyra pyra pyra pyra pyra				i		x	
133 134	Per pyra	10.0		?	i	cd		
135 136	Per pyra	14.0		?	1 1 1	cd		
137 138 139	pyra pyra pyra				i i		X	
140 141 142 143	Per pyra pyra pyra	10.0		r	1 1 1	cd		
144 145 146 147 148	pyra pyra Per pyra pyra	9.0		r	ਜਿਸਾਦਾਦਾਦਾਦਾਦਾਦਾਦਾਦਾਦਾਦਾਦ	cd	х	
149 150 151 152 153	pyra pyra pyra pyra pyra pyra				7 7 7 7 7		x	

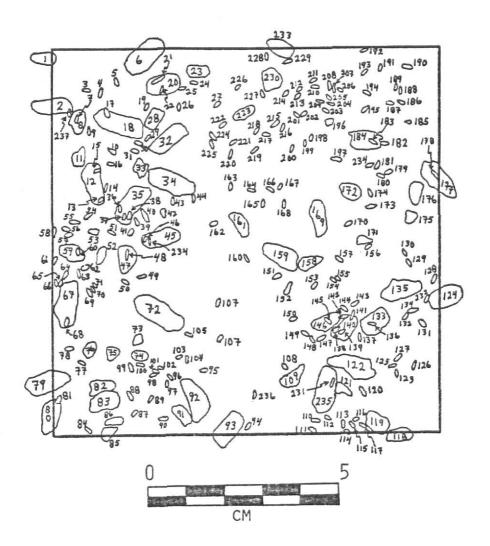
CU-13 Area #1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
154	pyra				20			•
155	pyra				i		x	
156 157	pyra pyra				치치나기 O 디디디디디디디디디디디디디디디디디		x	
158	unid				0		x	
159	Per	14.0		r	i	cd		
160 161	pyra Per	8.0		-	i	cđ		
162	pyra	0.0		r	i	cu	x	
163	pyra				i			
164 165	pyra				i			
166	pyra pyra				i			
167	pyra				1			
168	pyra	40.0		-	i			
169 170	Per pyra	10.0		1	1	cd	x	
171	unid						x	
172	Per	6.0		r	0 i i	cd		
173 174	pyra				i			
175	pyra Per	6.0		ı		cu	548	
176	Per	7.0		r	i	cd		
177	Per	12.0		1	i	cd		
178 179	pyra pyra				e i i i i i i i i i i i i i i i i i i i		x	
180	pyra				i		x	
181	pyra				i		x	
182 183	pyra pyra				ı			
184	Per	9.0	x		i	cd		
185	pyra				i			
186 187	pyra							
188	pyra pyra				i			
189	pyra				1		x	
190 191	pyra				i			
192	pyra pyra				i			
193	pyra				i			
194 195	pyra				i			
196	pyra				i			
197	pyra				i			
198	pyra				i		x	
199	pyra pyra				1111111111111111111		x	
201	pyra				i			
202	pyra				i			
203 204	pyra pyra				i		x	
	(100)						163	

CU-13 Area #1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
205 206	pyra pyra				다 저 차 더 저 저 저 저 저 서 저 저 저 저 저 저 저 저 저 저 저 저 저		x	
207	pyra				i		x	
208	pyra				i		x	
209	pyra				i			
210	pyra				i			
211	pyra				1			
212	pyra				i		x	
213	pyra				i			
214	pyra				i			
215	pyra				i			
216	pyra				i			
217	pyra				i			
218	pyra				i			
219	pyra				i		X	
220	pyra				i			
221	pyra				i			
222	pyra				i		X	
223	Per	7.0		r	i	cd		
224	pyra				i		x	
225	pyra				2		x	
226	pyra				1		x	
227 228	pyra				1		X	
229	pyra				1			
	pyra	8.0		200	-	cd		
230 231	Per	0.0		r	4	ca		
232	pyra				i			
233	pyra Per	8.0		r	4	cd		
234	pyra	0.0		-	: : : : : :	u		
235	Per	9.0		1	÷	cd		
236	pyra	7.0		-	i	ou	x	
237	pyra				ī			
**************************************	• •							

## CU-13 AREA #1



CU-13 Area #2

#	name	size	art	val	pres	orient	frag	epiz
1 2 3	Per pyra pyra	8.0		r	너저너저어서 어머니 이어 나 서어 시에 나 이어 지어 나 나 나 나 나 나 나 나 나 나	cd	x	
145	Per Per	7.0 13.0		? ! ? !	i	cd cd		
234567890	Per Per	9.0		? 1	i	cd	x	
9	pyra				i i			
11 12	pyra unid pyra				0		x	
13 14	pyra pyra				i			
15 16	pyra pyra				1		x	
17 18	pyra Per	7.0		r	i i	cd		
19 20	Per pyra	10.0		1	i i	cd		
2 <b>1</b> 22	Per pyra	5.0		r	<u>1</u> 1	cd	9	
23 24	pyra Per	6.0		?	i	cđ		
25 26	pyra pyra				i			
27 28	Per Per	7.0 7.0		? r	i	cd cd		
29 30	Per Per	6.0 6.0		r 1 ?	i	cd cd		
3 <b>1</b> 32	pyra pyra				i			
33 34	pyra pyra				i		x	
35 36	Per pyra			?	i	cd	x	
37 38	pyra pyra				i		x	
39 40	pyra pyra				i i			
41 42	pyra pyra				i			
43	pyra pyra				i i			
45	pyra pyra				i i		x	
31 31 31 31 31 31 31 31 31 31	pyra pyra				ਸਿ ਸੀ ਸੀ ਅਤੇ ਸੀ		x	
49 50	pyra pyra				i i			
51	pyra				1			

CU-13 Area #2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
53 55 55 55 55 55 56 56 56	pyra pyra pyra	2.0			i i i i	<b>.</b>		
55 56	Per pyra	8.0		r	ı i	cd		
57	ost	40.0		r	0	cd		
50	Per Per	13.0 5.0		r 1 1	i i	cd cd		
60	Per	5.0		ī		cd		
61 62	unid Per	8.0		1	o i	cu cd	x	
63	Sept	0.0		r	0	cd		
63 64 65 66	Pseu			r	0	cd		
65 66	Per Per	9.0 8.0		r	i	cd cd		
67	Per	8.0		r	i	cd		
68	Per	13.0		r	i	cd	D) de	
69 70	unid Per	7.0		r	o i	cd	x	
71	pyra	100			i		x	
72 73 74 75 76 77 78	Sept			1	0	cd	x	
7 <u>1</u>	Per unid			r	e 0	ou	x	x
75	Per	12.0		l	i	cd		27.77
76	Per Per	6.0 8.0		r	i i	cd cd		
78	unid	0.0		-	0	cu	x	
79	Per	7.0		r	į	cd		
80 81	pyra pyra				ı i			
82	Per	9.0		r	71 °1 °1 °1	cd		
83 84	Per unid	9.0		1.	i	cd		
85	pyra				o		x	
86	Per	8.0		r 1		po		
87 88	Per Per	9.0 7.0		r	i	cd cd		
89	pyra	1.0			i	- Cu		
99 91 93 93 94 95 95 97 89 99	Per			? ? 1 1 ? 1 # ? 1	i	cd	X	
91	Avic Per	12.0		i	i	cd cd	x	x
93	Per	7.0		1	i	cd		
94	Per Per	7.0		1	i	cd cd	x	
96	Per	7.0		i	i	cd	Α	
97	Per	11.0		r	i	ed		
99	Per Per	8.0 9.0		í	i	cd cd		
100	pyra			-	+ + + + + + 0 + + + + + + + + + + + + +	-		
101 102	pyra				i			
102	pyra				1		x	

CU-13 Area #2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103 104 105 106	Per pyra pyra unid			?	i i	cd	x	
107 108 109 110	Per pyra pyra	9.0		1	0 서서서서서서서서서서	cd	x	
111 112 113	pyra Per Per Per	7.0 4.0 11.0		r r 1 ?	i i	ed ed ed		
114	Per				i	cd	x	
115 116 117	Per Per pyra	12.0		r	i i	cd cd	x	
118	unid	۲ ۵		•	0		x	
119 120	Per Per	5.0 7.0		? 1	i	cd cd		
121	Per pyra	8.0		1	4444444	cd		
123 12կ	Per Per	13.0 11.0		1	i	cd cd	570	
125	Per	12.0		r	i	cd		
126 127	Per Per	9.0		? <b>r</b>	i	cd cd		
128	pyra unid	,.0		•	i	cu		
129 130	pyra				i		x	
131 132	Per pyra	6.0		r	0 1 1 1 1 1	cd		
133	pyra				i	2		
134 135	Per pyra	11.0		1	i	cd		
136	pyra	6.0		-	7.5			
137 138	Per pyra	6.0		1	i	cd		
139 140	unid pyra						X	
141	Per	10.0		r	i	cd		
142 143	Per Per	7.0		r ?	i	ed ed	X	
144	Per	9.0	x	•	i	cd		
145 146	pyra pyra				i		x	
147 148	pyra				i			
149	pyra pyra				i			
150 151 152	Per pyra	9.0		1	i	cd		
152	Per	7.0		1	0 기 귀 가 기 저 기 기 기 기 기 기 기 기 기	cd		
153	Per	7.0		r	1	cd		

CU-13 Area #2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
154	pyra				તનનનનનન			
155	pyra				i			
155 156 157	pyra				1		X	
158	pyra pyra				<u>+</u>			
158 159	pyra				i			
160	Per	5.0		r	i	cd		
161	pyra				i			
161 162	pyra				i			
163	pyra				i			
164	pyra	1 0		_	i	•		
165	Per	4.0		1	1	cd		
166 167	Per	8,0		r	1	cd	72	
168	pyra Per	200.000		?	÷	cd	x	
169	Per	12.0		? 1	ਸਿਹੀ ਸੀ	cd		
170	pyra	•		_	i			
171	pyra				1			
172	pyra				i			
173	pyra				i			
174	pyra				i		x .	
175 176	pyra				1			
177	pyra Per	5.0		r	- <del>-</del> -	cd		
178	pyra	J.0		-	i	cu		
179	Per	7.0		1	i	cđ		
180	Per	10.0		r	i	cd		
181	pyra				i		x	
182	Per			1	i	cd	X	
183	pyra				i			
184	pyra	9.0			<u> </u>			
185 186	Per Per	8.0 9.0		r	1	cd cd		
187	pyra	7.0			i	ca		
188	pyra				i			
189	bara				i			
190	pyra				i			
191	pyra				i	1		
192	pyra	0 -		_	i	-		
193	Per	8.0		1	i	cd		
194	pyra	1. 0		1	<u>.</u>	- 4		
195 196	Per Per	4.0 9.0		r	1	cd cd		
197	unid	<b>7.0</b>		1	o	Cu	x	
198	pyra				i			
199	pyra				i			
200	Per			?	i	cd	x	
201	Per	7.0		?	i	cd		
202	pyra	0 0		<u></u>	4444444440444444	- 3		
203 204	Per Per	8.0 <b>11.</b> 0		r	1	cd cd		
CUL	I GT.	11.0		T.	1	cu		

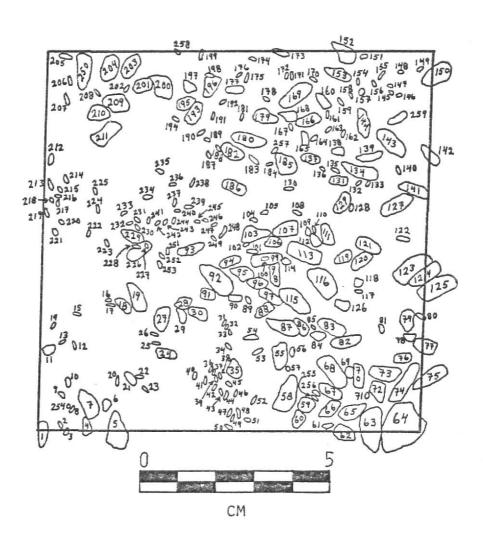
CU-13 Area #2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
205 206 207 208 210 211 212 213 214 215 216 217 218 219 220 221 222 223 224	pyra pyra pyra pyra Per Per Per pyra pyra pyra pyra pyra pyra pyra pyr	8.0		? ? ?	너 너 너 너 너 너 너 너 서 서 너 너 너 너 서 서 서 너 서	ed ed ed	x	
225 226	pyra pyra Per	6.0		r		cd	1.00	
227 228	pyra pyra			-	i			
229 230 231 233 233 235 236 237 239 241 243 244 244 244 244 245	Per pyra pyra pyra pyra pyra pyra pyra pyr	7.0		r	커 커 건 건 건 건 건 건 건 건 건 건 건 건 건 건 건 건 건 건	cď		
246 247 248 249	pyra pyra Avic pyra			?	i 0 i	cd	x	
250 251 252	Per pyra pyra	7.0		1	0 1 1 1	cu		
253 254 255	pyra ost pyra			1	i 0 i	cd	x	

CU-13 Area #2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
256 257 258 259	pyra pyra pyra Per			?	i i i	cd	x	

CU-13 AREA #2



#	name	size	art	val	pres	orient	frag	epiz
1	unid				0			
2	Avic			?	0	cd	x	
3	Avic			? ? 1	0	cd	x	
1234567890	Avic			ı	0	cd		
5	unid				0			
6	unid				0	cu		
7	Avic			?	0	$\operatorname{cd}$	x	
8	unid				0		x	
9	unid				0		x	
	Avic			1	0	cu		
11	Avic			?	0	cu	x	
12	Avic			1 ? ? ?	0	cu		
13 14	Sept			?	0	cu		
14	unid				0			x
15 16	Per	11.0		1 ?	0	cu		x
16	Sept			?	0	cd		
17 18	unid				0	cn		
18	unid				0		x	
19	unid				0		x	
20	unid			•	0		x	
21	Avic			?	0	ed	x	
22	Sept			ı	0	cd		
23	unid	0 0			0		x	
24	Per	8.0		r	i	cu		
25 26	unid				0		x	
26	unid			•	0	v2.5 <b>v</b>	X	
27 28	Avic			?	0	cd	x	
28	Avic			î.	0	cd	Table 1	
29	unid	40.0		0	0		x	
30	Per	12.0		?	0	cd		
31 32	Avic			1	0	$\operatorname{cd}$		
32	Avic			1	0	cd		
33	unid				0		x	
24	unid			?	0	cd	x	
35	Avic unid			•	0	Gu	x	
27		12 0		1	0	011	A	
28	Per Avic	13.0		r	0	cu		
20	unid			1	0	Ca	x	
10	unid				0		Α.	
1.1	unid				0			
12	unid				o			
1.3	Avic			?	0	cd		
المال	Avic			?	0	cd		
15	unid				0		x	
315 35 36 37 38 39 41 41 41 41 41 41 41 41 41 41 41 41 41	Avic			?	0	cd	x	
47	Sept			1	i	cu	STATE OF THE PARTY	
48	unid				0	- mm/07-27/		
49	Avic			1	0	cd		
50	Sept			1 1 ?	0	cd		
50 51	Avic			?	0	cd	X	
57.1								

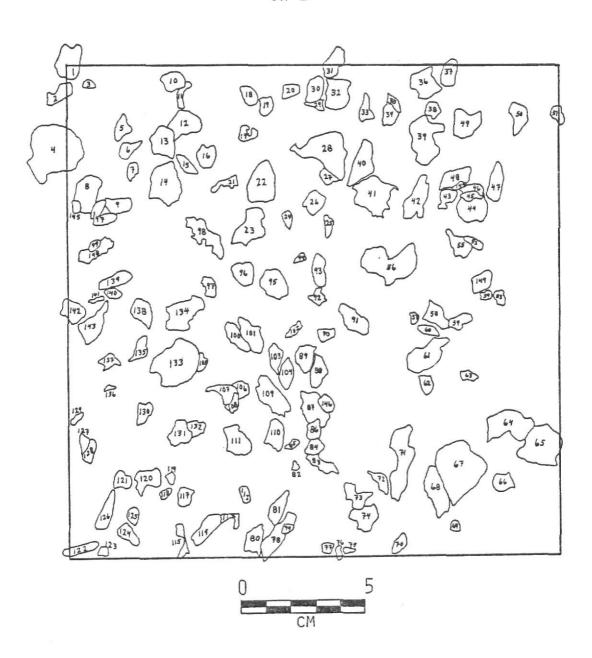
CW-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
52 53 55 55 55 55 55 56 57 89 60	Avic			?	0	cd	x	
53	Avic			?	0	cd	x	
54	Avic			?	0	cd	X	
55	unid				0		x	
56	unid				0		x	
57	Avic			?	0	cu		
58	unid				0			
59	unid				0			X
60	Per	11.0		?	i	cu		
61	Sept			?	0	$\mathbf{c}\mathbf{d}$		
62	Per	9.0		r ?	i	cu		
63 64 65 66	Avic			?	0	cd	x	
64	unid				0			
65	Avic			1	0	cd		
66	Avic			1	0	cd		
67	Sept			r	0	cd		
68	Per	22.0		?	0	cd		
69	Avic			?	0	cd	X	
70	unid				0			
71 72	unid				0		X	
72	Avic			?	0	cd	X,	
73 74	Avic			?	0	$\mathtt{cd}$	x	
74	unid				0			
75	Avic			?	0	cd	X	
76	unid			65	0		x	
77	Avic			?	0	cd	x	
78	Sept			r ?	0	cd		
79	Avic							
80	Sept			r	0	cd		
81	Sept			r	i	cd		
82	Sept			?	0	cd	x	
83	Avic			? ? ?	0	cd	x	
84	Avic			?	0	cd	X	
85	Avic				0	cd	X	
86	Avic			1	0	cd		
87	Avic			1 ? ?	0	cd	x	
88	Sept			?	0	cd		
89	unid				0			
90	Avic			? ? ?	0	cd	X	
91	Avic			?	0	cd	X	
92	Avic			?	0	cd	X	
93	unid			_	0			
94	Avic			?	0	cd	x	
95	Avic			? ? ?	0	cd		
91 92 93 94 95 96 97	Avic			?	0	cd	FUSE	
97	unid				0		x	
98	unid				0	5. 9	x	
99	Avic			r	0	cd		
100	Sept			?	0	cd		
101	unid			_	0			
102	Avic			?	0	cd	X	

CW-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
103	Avic			?	0	cd	x	
104	unid				0		x	
105	Avic			?	0	cd	x	
106	unid				0		x	
107	unid				0		x	
108	Avic			?	0	cd	x	
109	Avic			1	0	cd		
110	unid				0			
111	Avic			?	0	cd	x	
112	unid				0			
113	Avic			1	0	cu		
114	unid				0			
115	Avic			r	0	cd		
116		fossil						
117	unid				0		x	
118	Avic			?	0	cd	x	
119	Avic			?	0	cd	x	
120	Avic			1	0	cu		
121	unid				0			
122	Sept			r	0	cu		
123	unid				0		x	
124	unid				0		-	
125	unid				0		x	
126	Sept			?	0	cd		
127	Sept			?	0	$\mathbf{c}\mathbf{d}$		
128	unid				0			
129	unid				0			
130	Sept			?				
131	Avic			?	0	cd		
132	unid				0			
133	Pseu	7.9		r	0	cd		
134	Avic			3	0	cd	x	
135	Sept			?	0	cd		
136	unid				0		x	
137	unid				0			
138	unid				0			
139	unid				0	cu		
140	unid				0			
141	unid				0		x	
142	Avic			?	0	cd	x	
143 144	unid				0	cu		
144	Per	6.0		?	0	cu		
145	Avic			?	0	cd	x	
146	Avic			?	0	cd	x	
147	Per	10.0	x	ř.	i	cu		
148	Per	16.0		1	i	cu		
149	Sept			1	0	cd		
150	unid				0			

CW-1

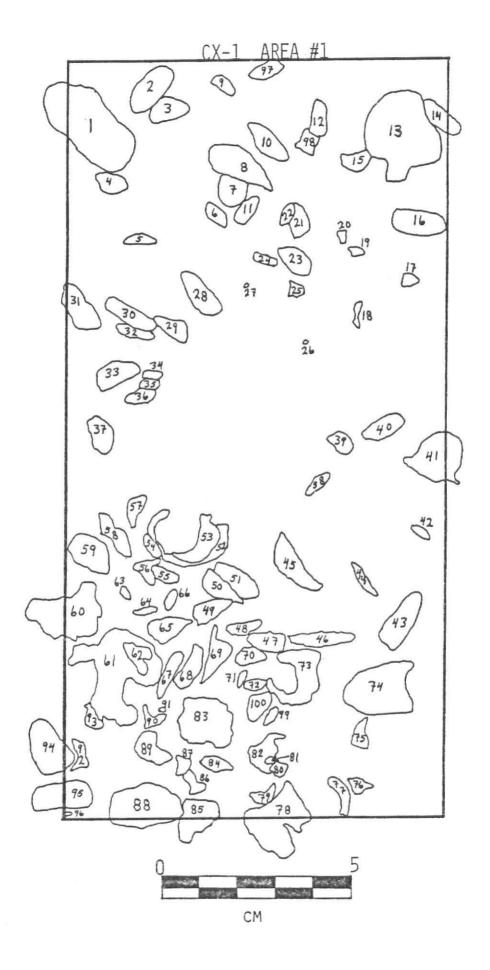


CX-1 Area #1

#	name	size	art	val	pres	orient	frag	epiz
1 2 3 4 5 6 7 8 9 10	Per Per Edm	30.0 15.0		r 1 ?	e i e	cd cu		
7	Avic			1	0	cd		
5	Sept unid			Τ	o i	cu		
7	Bell				i	Cu		
8	Sept			1	0	cd		
9	Per	9.0		1	1	cu		
	Per	13.0		r	е	cd		
11 12	Sept Sept			r	0	cu		
13	Avic			r 1 1 ?	0	cd		
14	Per	12.0		1	е	cd		
15 16	Avic				0	cd		
16	Per	18.0		r ? ?	е	cd	V225	
17 18	Avic Avic			?	0	cd cu	x	
19	unid			•	0	Cu	×	
20	Avic			?	o	cd	x	
21	Per	10.0		r	i	cu		
22	Per	7.0		r	i	cu		
23	Per	10.0		?	е	cd		
24 25	Per Avic	11.0		? 1	e o	cd cd		
26	ost			r	0	cd .		
27	ost			r	o	cd		
27 28	Per	17.0		1	е	cd		x
29	Sept			r	0	cu		
30	Per	15.0		?	е	cd		
31 32	Sept Per	12.0			o e	cd	x	
33	Per	16.0		? ? !	e	cd		
34	Per	6.0		ī	i	cu		
35	Per	7.0		r	i	cu		
36	Per	9.0		1		cu		
37	Sept	10.0		r 1 ? ? r 1 ?	0	cd	x	
30	Per Per	10.0			i e	cu cd		
1.0	Per	13.0		ī	e	cd		
41	Avic			?	0	cd		
42	unid				0		x	
43	Per	17.0		?	0	cd		
777	Per	15.0			0	i		
1.6	unid Per	18.0		ı	o i	cu		
3456789012345678901	Sept	.0.0		r	0	cu		
48	Sept			r	0	cu		
49	Sept				0	cu		
50	Sept	42.0		r	0	cu		
57	Per	13.0		r	е	cd		

CX-1 Area #1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
234567890	Avic Avic Per	6.0		? ? ?	0 0 i	cd cd cu		
55	Sept			?	0	cd		
56 57	Avic Per	14.0			o	cd cu		
58	Avic			r ?	0	cd	x	x
59	Per			1 ?	0	cd		
61	Avic Avic			?	0	cd cd		
62	alfr					-		
63 64	Per	4.0		r	i	cu		
64	unid Sept			r	0	cu		
65 66	unid			-	0	cu		
67	Per	17.0		?	i	cu		
68	Avic			?	0	cd		
69 70	Avic Sept			r	0	cd cu		
71	unid				0		x	
72	Sept			?	0	cd	x	
73 74	Avic unid			1	0	cd		
75	unid				0			
75 76	unid				0		x	
77	Avic			?	0	cd		
78 79	unid Avic			?	0	cd	x	
80	Avic			?	o	cd	x	
81	Avic			?	0	cd	x	
82 83	unid			?	0	cd	x	
84	Avic Sept			i	0	ca		
85	Avic			r	0	cđ		x
86	Avic			? ? ?	0	cd	x	
87 88	Avic Avic			?	0	cd cd	x	
89	Avic			?	0	cd		
90	Sept			r	0	cu		
91 92	unid			1	0	- 4	x	
93	Avic unid			_	0	cd	x	
93 94 95	Per			r	1	cu		
95	Per	16.0		r	е	cd		
96 97	pyra Per	9.0		٦	i	cu		
98	Avic	7.0		1	0	cd		
99	unid				0	cu		
100	Sept			1	0	cu		



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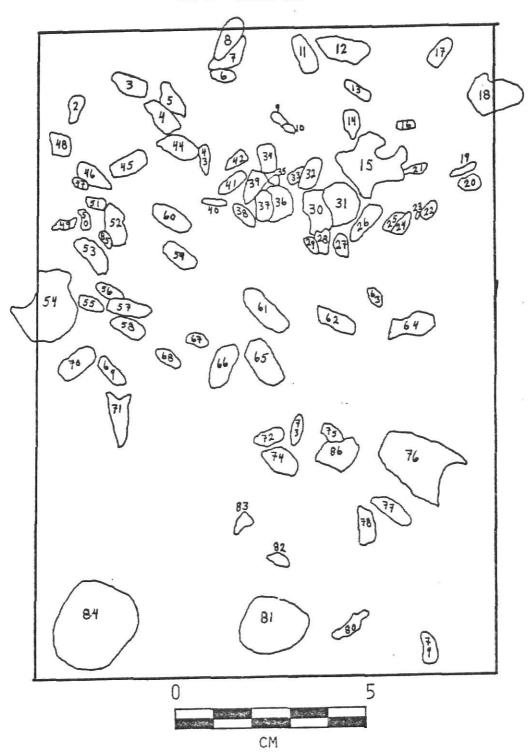
CX-1 Area #2

#	name	size	art	val	pres	orient	frag	epiz
1	Not a unid	fossil			0			
2345678	Per	13.0		r	o i	cu		
Ĺ	Per	10.0		1	i	cu		
5	Per	10.0		r	i	cu		
6	unid			0	0	•	X	
9	Avic	11.0		?	0 <u>i</u>	cd	x	
9	Per Per	5.0		2		cu		
10	Per	5.0		1 ? ?	i i i	cu		
11	Per	5.0 12.0			i	cu		
12	Per	14.0		r 1	i	cu		
12 13 14 15 16 17	unid			_	0		x	
14	Per Avic	9.0		1	i	cu		
16	Per	6.0			3	cd cu		
17	Per	9.0		r	o i	cu		
18	Avic	,		?	0	cd		
19	unid	_			i i	cu		
20	Per	8.0		?		cn		
21	Sept	6.0		?	0	cu		
22 23	Per	6.0		1	i i i	cu		
211	pyra Per	8.0		?	+	cu		
24 25 26	Per	8.0		?	i	cu		
26	Per	13.0			0	cd		
27	Per	9.0		r 1 ?	i i	cu		
28	Per	7.0		?	i	cu		x
29 30	Per unid	7.0		?	i	cu		X
31	Edm			?	o e	cd		
32	Per	10.0		r	i	cu		
32 33	Per	5.0		1	i	cu		
34	Per	10.0		1	i	cu		
35	unid	0 0			0		x	
36	Per	8.0		ı	i	cu		
37 38	Sept Per	8.0		r 1 ? ?	o i i	cu		
39	Per	11.0		ī	1	cu		
40	Sept			?		cu		
41	Per	9.0		?	o i	cu		
42	Per	9.0				cu		
43	Sept			r	0	cu		
39 40 41 43 44 45 6	unid Per	11.0		70	o i	cu		
46	Sept	. 1 . 0		r 1		cd		
47 48	Per	6.0			e i i	cu		
48	Per			?		cu		
49	Avic			?	0	cd	x	
50 51	Avic	6.0		r ? ?	0	cd	x	
וכ	Per	6.0		1	i	cu		

CX-1 Area #2 Cont.

#	name	size	art	val	pres	orient	frag	epiz
55555555666666666690772	Sept Per Avic Per Sept Per	13.0 8.0 11.0 9.0 8.0 11.0 16.0 17.0 14.0 7.0 8.0 10.0 14.0		?1111		cu cd		
73 74 75 76 77 78 79 80 81 82 83 84 85 86	Per Per unid unid Per Per Per unid Edm Sept unid Edm Sept Avic	10.0 11.0 13.0 11.0		rr rrl l?	00000110000100	ed ed eu ed ed ed	x x	

CX-1 AREA #2



				150 50 50				
#	name	size	art	val	pres	orient	frag	epiz
1 2 3 4 5 6 7 8 9 10	pyra pyra pyra pyra pyra pyra pyra pyra				ਜਿਜਕਜਿਜਜਜਿਜ			
11 12 13 14	pyra Per Avic Per Per Per pyra	6.5 7.5 6.5 9.5		r 1 1	0 1 1 0 1	ed ed ed	x	
15 16 17 18 19 20	Avic Per Per Per Per pyra	14.0 9.0 9.5 5.5		r r r	0 i i i i i i	cd cd cd cd	x	
21 22 23 24 25 26	Avic Avic Per Per Per	9.0 11.0 11.0 6.5		1 1 1 r	0 0 1 1 1 1 1	cd cd cd cd cd	x	x
27 28 29 30 31	pyra Avic Per pyra Avic	13.0		1 r ?	0 i	cd cd cd	х	
32 33	unid Per pyra Per Pseu Per Per Per	8.0 9.5 7.5 31.0 9.0		r r l	0 i i	cd cd cd cd cd	x	X
345 35 36 37 38 39 41 42 34 44 45 46 47 48	Avic Per Per Avic Per pyra Per Per Per	12.5 6.5 8.0 11.5 11.0		rilir?i rir	0 1 1 1 1 1 1 1 1 0 1 1 1 1 1 1 1 1 1 1	ed ed ed ed ed ed		х
49 50 51	Per Per pyra	9.0 13.5		r 1 1	i i i	cd cd		

J-1-1 Cont.

#	name	size	art	va1	pres	orient	frag	epiz
52 53 54 55 56 57 58 59 60 61 62 63	Per Per Per Per Per Per Per Per Per Per	8.5 10.0 6.5 10.5 10.0 8.0 6.0 4.5 6.0 3.5 5.0		1 r r r 1 1 r	i i i i i i i i i i i i i i i i i i i	cd cd cd cd cd cd cd cd		
64 65 66 67 73	Per alfr alfr alfr unid	5.0		1	i	cd		
114 P P P P A	Sept Per Per Per Per	11.0 12.0 9.5 12.5		1 1 r 1	o i i i	cu cd cd cd	х	
P P S S	Avic Per Per pyra pyra	14.0		1 1 1	o i i i	cd cd		
P C C S	Per alfr alfr pyra pyra	6.0		1	i i i	cd		
P S S S	Per pyra pyra pyra pyra	9.0		1	i i i i	cd		
P P P S	Per Per Per pyra Per	13.0 6.0 8.5		1 1 1	e i i i	cu cd cd		
P S P P	Per pyra Per Per	7.5 14.0		1 r 1	i i i i	cd cd		
P P P P	Per Per Per Per Per	8.5 14.5 7.5 10.5 9.0		r 1 1	i i i i	cd cd cd cd		x

J-1-1 Cont.

45 ac c

#	name	size	art	val	pres	orient	frag	epiz
C P P P P P P P P P	alfr Per Per Per Per Per Per	15.5 12.5 10.0 13.0 13.5 7.5 10.5 8.5		1 1 1 1 r 1 r	i i i i i i	cd cd cd cd cd cd		х
S S P P P S S C	pyra pyra pyra Per Per Per pyra	10.0 11.0 10.5		r 1 1	i i i i i i	cđ cđ cđ		
C S P P S	alfr alfr alfr pyra Per Per pyra	14.0 10.0		r 1	i i i i	cd cd		
P P P P A P	Per Per Per Per Avic Per Per	5.5 8.5 13.0 8.0 12.0		1 r r 1 1 1	i i i i o i	cd cd cd cd cu cd	x	
P S S P P P	Per pyra pyra Per Per Per	13.0 10.0 11.0 17.0 9.5		1 1 r r	i i i i i	cd cd cd cd		
P P S A S C P	Per Per pyra Avic pyra alfr Per	12.5 11.5	х	1 1	i i o i	cd cd		
P P P	Per Per Per Per	11.0 10.5 5.0 11.5		1 r r 1	i i i	cd cd cd		

J-1-1 Cont.

#	name	size	art	val	pres	orient	frag	epiz
P	Per	10.5		r	i	cd		
P	Per	10.0		1	i	cd		
P	Per	10.0		1	i	cd		
S	pyra				i			
S	pyra				i			
P	Per	10.0		1	i	cd		
P	Per	14.0		r	i	cd		
P	Per	9.0		r	i	cd		
P	Per	12.0		r	i	cd		
P	Per	14.5		1	i	cd		
P	Per	7.0		r	i	cd		
P	Per	11.0		r	i	cd		
P	Per	13.5		r	i	cd		
P	Per	11.5		r	i	cd		
P	Per	14.5		1	i	cd		
P	Per	8.0		1	i	cd		
P	Per	6.5		1	i	cd		
P	Per	9.5		r	i	cd		
P	Per	11.0		1	i	cd		
Α	Avic			1	0	cd		
S	pyra				i			
P	Per	12.0		r	i	cd		
Α	Avic			1	0	cd		
S	pyra				i			
P	Per	9.5		r	i	cd		
S	pyra				i			
A	Avic			?	0	cd	x	x
P	Per	6.5		1	i	cd		
A	Avic			r	0	cd		
P	Per	8.0		1	i	$\operatorname{cd}$		
P	Per	5.5		r	i	cd		
S	pyra				i			
P	Per	7.5		r	i	cd		
Α	Avic			1	0	cd		
C	alfr							
C	alfr							
A	Avic			1 ?	0	cd		x
A	Avic			?	0	cd		
C	alfr							
A	Avic			r	0	cd		
A	Avic			?	0	cd	X	
A	Avic			r	0	cd		
C	alfr							
S	pyra				i			
P	Per	11.5		1	i	cd		
A	Avic			r	0	cd		
A	Avic			r	0	cd		
A	Avic	0 5		r	0	cd		
P	Per	9.5		1	i	cd		
P A	Per	8.5		r	i	cd		
A	Avic			r	О	cd		

J-1-1 Cont.

#	name	size	art	va1	pres	orient	frag	epiz
A P P	Avic Per Per	6.5 8.0		r r 1	o i i	cd cd cd		

## APPENDIX IV

Thin sections from 99 samples were examined. Rock names were assigned according to Folk's (1962) classification of limestones and grain size scale for carbonate rocks. The stratigraphic positions from which the samples were collected are shown in Appendix I. Sample numbers with an "a" or "b" at the end (e.g. CQ-7a, CQ-7b) designate a slide made from the bottom or top parts respectively, of a sample.

- \* indicates thin sections point counted as representative of the four major rock types
- \*\* indicates other thin sections point counted

Sample	Rock Name
**C0-1	Medium calcarenite: immature, recrystallized bivalve, ostracode, gastropod, foraminifer silty Biomicrite
CO-2	Fine calcarenite: immature, recrystallized ostracode, foraminifer, silty, ferruginous Biomicrite
CO-3	Fine calcarenite: immature ostracode, foraminifer, Osagia, ferruginous Biomicrite
co <b>-</b> 5	Ferruginous Dismicrite
CO-7	Medium calcarenite: immature recrystallized ostracode, foraminifer, bivalve, gastropod, echinoderm, pelloid-bearing Biomicrite
CO-8	Medium calcarenite: immature recrystallized bivalve, foraminifer, gastropod, ostracode, algal, echinoderm Biomicrite
୯୧-2	Fine calcirudite: immature recrystallized bivalve, gastropod, ostracode, foraminifer, echinoderm, Osagia Biomicrudite
CQ-3	Coarse calcarenite: immature recrystallized bivalve, gastropod, ostracode, Osagia Biomicrite
CQ-4	Coarse calcarenite: immature recrystallized gastropod, bivalve, ostracode, foraminifer, bryozoan, Osagia Biomicrite
CQ-5	Coarse calcarenite: immature recrystallized gastropod, ostracode, bivalve, pelloid-bearing Biomicrite
ପଦ୍-6	Medium calcarenite: immature recrystallized gastropod, ostracode, foraminifer, bivalve, ferruginous Biomicrite
CQ-7a	Fine calcarenite: immature recrystallized ostracode, foraminifer, gastropod, bivalve, ferruginous Biomicrite
ପ୍ୟ-75	Medium calcarenite: immature recrystallized ostracode, gastropod, foraminifer, bivalve, Osagia, ferruginous Biomicrite
CQ-9	Dolomitized silty Micrite

Sample	Rock Name
CQ-10	Dolomitized ostracode, foraminifer, silty Micrite
CQ <b>-11</b>	Dedolomitized ostracode, foraminifer, silty Micrite
CQ-12	Recrystallized ostracode, bivalve, foraminifer, laminated silty Micrite
CQ-19	Fine calcirudite: immature recrystallized clast- bearing, ostracode, vertebrate, charcoal-bearing Biomicrudite
CQ-19+	Fine calcirudite: immature recrystallized clast- bearing, ostracode, vertebrate, charcoal-bearing ferruginous Biomicrudite
CR-1	Coarse calcarenite: immature recrystallized bivalve, gastropod, foraminifer, ostracode, echinoderm, Osagia Biomicrite
CR-2	Medium calcarenite: immature recrystallized bivalve, gastropod, foraminifer, ostracode, Osagia, ferruginous Biomicrite
CR-2+	Medium calcarenite: immature recrystallized ostracode, gastropod, bivalve, foraminifer, Osagia, silty Biomicrite
CR-4	Fine calcirudite: immature recrystallized gastropod, bivalve, ostracode, foraminifer, Osagia Biomicrudite
GR-5	Fine calcirudite: immature recrystallized gastropod, bivalve, foraminifer, ostracode, ferruginous Biomicrudite
CR-7	Fine calcarenite: immature recrystallized ostracode, silty Biomicrite
CR-8	Fine calcarenite: immature recrystallized ostracode, foraminifer Biomicrite
CR-9	Medium calcarenite: immature recrystallized ostracode, foraminifer, ferruginous Biomicrite
**CS-1	Fine calcirudite: immature recrystallized bivalve, gastropod, algal, ostracode, <u>Osagia</u> , pelloid-bearing Biomicrudite
CS-5	Fine calcirudite: immature recrystallized gastropod, bivalve, foraminifer, ostracode, pelloid-bearing Biomicrudite

Sample	Rock Name
*CS-6	Coarse calcarenite: immature recrystallized gastropod, bivalve, foraminifer, ostracode, ferruginous Biomicrite
CS-7	Fine calcarenite: immature recrystallized ostracode, foraminifer Biomicrite
CS-11	Dolomitized ostracode Micrite
CS-14	Dedolomitized laminated silty Micrite
*CS-15	Dolomitized ostracode, foraminifer silty Micrite
*CS-16	Dedolomitized ostracode laminated silty Micrite
CT-2	Fine calcirudite: immature recrystallized bivalve, gastropod, ostracode, algal, echinoderm, Osagia, silty Biomicrudite
CT-5	Coarse calcarenite: immature recrystallized gastropod, bivalve, ostracode, foarminifer, pelloid-bearing Biomicrite
CT-6	Fine calcirudite: immature recrystallized bivalve, gastropod, ostracode, foraminifer Biomicrite
CT-7	Medium calcarenite: immature recrystallized ostracode, foraminifer, gastropod, bivalve, pelloid-bearing Biomicrite
CT-8	Fine calcirudite: immature recrystallized gastropod, bivalve, foraminifer, ostracode, Csagia, ferruginous Biomicrudite
CT-9	Medium calcarenite: immature recrystallized ostracode, gastropod, foraminifer, ferruginous, silty Biomicrite
CT-12	Recrystallized ostracode, foraminifer, silty Micrite
CT-14	Dedolamitized ostracode, silty Micrite
CT-15	Dolomitized ostracode, silty Micrite
CT-16	Dolomitized ostracode, laminated, ferruginous, silty Micrite
CT-17	Dedolomitized ostracode, silty Micrite

Sample	Rock Name
CU-1	Coarse calcarenite: immature recrystallized bivalve, gastropod, ostracode, foraminifer, silty Biomicrite
CU-2	Coarse calcarenite: immature recrystallized bivalve, ostracode, foraminifer, silty Biomicrite
CU-3	Coarse calcarenite: immature recrystallized gastropod, ostracode, bivalve, foraminifer, silty Biomicrite
CU-7	Fine calcirudite: immature recrystallized gastropod, bivalve, foraminifer, ostracode, echinoderm, ferruginous Biomicrudite
OU-5	Dedolomitized ostracode, silty Micrite
CU-6	Fine calcarenite: immature dedolomitized ost- racode, foraminifer, ferruginous, silty, Bio- micrite
CU-7	Dedolomitized ostracode, silty Micrite
CU-8	Dedolomitized ostracode, laminated, silty Micrite
CU-11	Recrystallized ostracode, foraminifer, gastropod, silty Micrite
CU-13	Coarse calcarenite: immature recrystallized gastropod, bivalve, algal, ostracode, foraminifer, Osagia, ferruginous Biomicrite
*CV-1	Coarse calcarenite: immature recrystallized gastropod, bivalve, ostracode, foraminifer, algal, Osagia, ferruginous Biomicrite
CV-2	Medium calcarenite: immature recrystallized bivalve, ostracode, foraminifer, Osagia Biomicrite
CV-3	Medium calcarenite: immature recrystallized ostracode, bivalve, foraminifer, Osagia, ferruginous, silty Biomicrite
CA-7	Fine calcirudite: immature recrystallized bivalve, gastropod, ostracode, foraminifer, Osagia, ferruginous Biomicrudite
CV-6	Silty, argillaceous Micrite

Sample	Rock Name
CV-7	Dedolomitized ostracode, foraminifer, silty Micrite
CV-8	Dedolomitized ostracode, foraminifer, silty Micrite
CV-9	Recrystallized ostracode, silty Micrite
CV-10	Recrystallized ostracode, foraminifer, silty Micrite
CV-11	Medium calcarenite: immature recrystallized ostracode, foraminifer, bivalve, silty Biomicrite
CV-13	Dedolomitized recrystallized ostracode laminated Micrite
CV-14	Recrystallized ostracode, ferruginous, silty Micrite
**CU=2	Coarse calcarenite: immature recrystallized bivalve, gastropod, ostracode, algal, foraminifer, Osagia, ferruginous Biomicrite
CW-3	Coarse calcaremite: immature recrystallized bivalve, gastropod, ostracode, algal, foraminifer, Osagia, ferruginous Biomicrite
CW-14	Medium calcarenite: immature recrystallized gastropod, foraminifer, ostracode, bivalve, algal, Osagia, ferruginous, Biomicrite
CW-5	Silty, argillaceous Micrite
CW-7	Dolomitized silty, argillaceous Micrite
CW-8	Dolomitized silty Micrite
CM-9	Recrystallized ostracode, foraminifer, silty, Micrite
CW-10	Dedolomitized foraminifer, ostracode, silty Micrite
CW-11	Dolomitized recrystallized ostracode, laminated, silty Micrite
CW-12	Fine calcarenite: immature recrystallized ost- racode, foraminifer, ferruginous, silty Micrite

Sample	Rock Name
CW-14a	Recrystallized ostracode, laminated, bioturbated, silty Micrite
CW-14b	Dolomitized, recrystallized ostracode, laminated, silty Micrite
CW-15	Fine calcarenite: immature dedolomitized, recrystallized ostracode, foraminifer, echinoderm Biomicrite
CX-1	Fine calcirudite: immature recrystallized gastropod, ostracode, foraminifer, bivalve, algal, Osagia, ferruginous Biomicrite
**CX=2	Medium calcarenite: immature recrystallized foraminifer, ostracode, gastropod, algal, bivalve, Osagia Biomicrite
CX-3	Coarse calcaremite: immature recrystallized gastropod, ostracode, bivalve, algal, foraminifer, echinoderm, Osagia, ferruginous Biomicrite
CX <b>-</b> 6	Fine calcarenite: immature recrystallized ostracode, foraminifer, silty Biomicrite
CX-7	Medium calcarenite: immature dedolomitized recrystallized ostracode, foraminifer, ferruginous, silty Biomicrite
CX-8	Dedolomitized, recrystallized ostracode, foram- inifer, laminated, silty Dismicrite
**CX-9	Recrystallized ostracode, foraminifer, laminated Micrite
**CY-1a	Coarse calcarenite: immature recrystallized bivalve, ostracode, gastropod, echinoderm, foraminifer, Osagia, pelloid-bearing, ferruginous, silty Biomicrite
CY-1b	Coarse calcarenite: immature recrystallized gastropod, bivalve, ostracode, echinoderm, foraminifer, algal, Osagia, ferruginous, silty Biomicrite
CY-2	Fine calcarenite: immature recrystallized foraminifer, ostracode, echinoderm, bivalve, pelloid-bearing, intraclast-bearing Biomicrite

Sample	Rock Name
CY-3	Fine calcarenite: immature recrystallized ostracode, foraminifer, echinoderm, bivalve, silty Biomicrite
CY-4	Coarse calcarenite: immature recrystallized gastropod, bivalve, ostracode, foraminifer, algal, Osagia, ferruginous Biomicrite
CY-7	Dolomitized silty Micrite
CY-8	Dolomitized, recrystallized ostracode, silty Micrite
CY-10	Dolomitized ostracode, echinoderm, foraminifer, silty Micrite
CY-11	Fine calcarenite: immature dolomitized, dedolomitized ostracode, foraminifer, bivalve, <u>Osagia</u> , silty Biomicrite
CY-12a	Fine calcarenite: immature recrystallized ostracode, foraminifer, bivalve, Osagia Biomicrite
CY-12b	Medium calcarenite: immature recrystallized gastropod, foraminifer, ostracode, bivalve, echinoderm, Osagia Biomicrite

#### APPENDIX V

Approximately 1400 points were counted on each of ten thin sections according to the methos described by Chayes (1949). These data are summarized on the following pages. The stratigraphic positions of the samples from which these thin sections were made are shown in Appendix I.

- \* indicates thin sections point counted as representative of the four major rock types
- \*\* indicates other thin sections point counted

\*\*Slide CO-1

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar	860 123 25	62.4 8.9 1.8	
Allochems (Grains) Bivalves Ostracodes Osagia Gastropods Encrusting Foraminifers Free-living Foraminifers Phylloid Algae Unidentifiable Skeletal Fragments	55 140 6 48 10 13 1	4.0 10.2 0.4 3.5 0.7 1.3 0.1	15.0 38.3 1.6 13.1 2.7 4.9 0.3
Terrigenous (Grains) Quartz Silt	5	0.3	
Total	1379	100.0	99•9

\*\*Slide CS-1

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix)			
Micrite	455	32.3	
Microspar	126	8.9	
Spar	78	5.5	
Limonite	7	0.5	
Allochems (Grains)			
Bivalves	169	12.0	23.1
Ostracodes	343	24.4	46.8
Osagia	69	4.9	9.4
Gastropods	47	3.3	6.4
Echinoderms	2	0.1	0.3
Encrusting Foraminifers	11	0.8	1.5
Free-living Foraminifers	21	1.5	2.9
Phylloid Algae	5 3	0.2	0.7
Pelloids	3	0.2	-
Unidentifiable Skeletal			
Fragments	66	4.7	9.0
Terrigenous (Grains)			
Quartz Silt	6	0.4	
Total	1408	99.9	100.1

\*Slide CS-6

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar Limonite	227 149 325 1	16.7 11.0 23.9 0.1	
Allochems (Grains)  Rivalves Ostracodes Osagia Gastropods Echinoderms Encrusting Foraminifers Free-living Foraminifers Unidentifiable Skeletal Fragments	114 90 6 266 7 8 5	8.4 6.6 0.4 19.6 0.5 0.6 0.4	18.8 14.8 1.0 43.8 1.2 1.3 0.8
Porosity	49	3.6	
Total	1358	100.0	100.0

\*Slide CS-15

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Nicrite Microspar Spar	1203 139 15	85.3 9.9 1.1	
Allochems (Grains) Ostracodes Free-living Foraminifers	11 1	0.8 0.1	91.7 8.3
Terrigenous (Grains) Quertz Silt Porosity	14 28	1.0 2.0	
Total	1411	100.2	100.0

\*Slide CS-16

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar	1116 219 44	79.1 15.5 3.1	
Allochems (Grains) Ostracodes Bivalves Free-living Foraminifers	10 2 3	0.7 0.1 0.2	66.7 13.3 20.0
Terrigenous (Grains) Quartz Silt Porosity	12 4	0.9	
Total	1410	99•9	100.0

\*Slide CV-1

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar	526 50 178	38.6 3.7 13.0	·
Allochems (Grains) Bivalves Ostracodes Osagia Gastropods Echinoderms Encrusting Foraminifers Free-living Foraminifers Unidentifiable Skeletal Fragments	239 197 21 25 2 9 9	17.5 14.4 1.5 1.8 0.1 0.7 0.7	39.8 32.8 3.5 4.2 0.3 1.5 1.5
Terrigenous (Grains) Quartz Silt	2	0.1	
Porosity	8	0.6	
Total	1364	99•9	99•9

\*\*Slide CW-2

Type of Particle	Number of	Percent of	Percent of
	Points	Rock	Skel, Grains
Orthochems (Matrix) Micrite Microspar Spar	182 255 336	13.6 19.1 25.1	
Allochems (Grains) Bivalves Ostracodes Osagia Gastropods Echinoderms Encrusting Foraminifers Free-living Foraminifers Unidentifiable Skeletal Fragments	194	14.5	34.4
	42	3.1	7.4
	134	10.0	23.8
	51	3.8	9.0
	3	0.2	0.5
	11	0.8	2.0
	7	0.5	1.2
Total	1337	99.8	99•9

Grain/Matrix R<sub>a</sub>tio = 0.730

\*\*Slide CX-2

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar	492 5 <b>1</b> 458	35•7 3•7 33•2	
Allochems (Grains)  Bivalves Ostracodes Osagia Gastropods Encrusting Foraminifers Free-living Foraminifers Unidentifiable Skeletal Fragments	152 1 11 64 41	0.3 11.0 0.1 0.8 4.6 3.0	1.1 40.9 0.3 3.0 17.2 11.0
Terrigenous (Grains) Quartz Silt	5	0.4	
Total	1378	100.0	100.1

\*\*Slide CX-9

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar	837 124 182	65.5 9.7 14.2	
Allochems (Grains) Bivalves Ostracodes Encrusting Foraminifers Free-living Foraminifers Unidentifiable Skeletal Fragments	2 73 5 4	0.2 5.7 0.4 0.3	1.5 54.1 3.7 3.0
Terrigenous (Grains) Quartz Silt	10	0.8	
Total	1278	100.0	100.1

\*\*Slide CY-1a

Type of Particle	Number of Points	Percent of Rock	Percent of Skel. Grains
Orthochems (Matrix) Micrite Microspar Spar	384 366 92	28.3 27.0 6.8	
Allochems (Grains) Bivalves Ostracodes Osagia Gastropods Echinoderms Encrusting Foraminifers Free-living Foraminifers Pelloids Unidentifiable Skeletal Fragments	134 139 120 6 8 19 9 12	9.9 10.2 8.8 0.4 0.6 1.4 0.7 0.9	27.8 28.8 24.9 1.2 1.7 3.9 1.9
Terrigenous Quartz Silt	14	0.3	
Porosity  Total	18 	1.3	100.0

# APPENDIX VI

Weight percent insoluble residue was determined and these data are listed in this appendix. The stratigraphic position from which the samples were collected is shown in Appendix I.

Sample	Weight Percent Insoluble Residue
CO-1	14.42
CO-2	13.13
CO-3	16.11
CO-1;	30.83
CO-5	10.82
CO-6	28.60
CO-7	8.34
CO-8	6.16
CQ-1 CQ-2 CQ-3 CQ-4 CQ-5 CQ-6 CQ-7 CQ-8 CQ-9 CQ-10 CQ-11 CQ-12 CQ-13 CQ-14 CQ-17 CQ-15 CQ-16 CQ-16 CQ-18 CQ-19 CQ-19+ CQ-20	38.28 14.28 8.82 10.87 4.12 4.30 10.22 35.47 22.28 21.99 12.21 10.78 10.10 11.01 76.83 84.58 83.47 78.28 9.50 23.54
CR-1	11.41
CR-2	10.16
CR-2+	24.98
CR-3	56.89
CR-4	8.18
CR-5	8.65
CR-6	37.97
CR-7	11.97
CR-8	7.95
CR-9	9.89
CS=1	11.88
CS=2	7.69
CS=3	13.56
CS=14	59.59
CS=5	7.89
CS=6	13.38
CS=7	10.33
CS=8	35.30
CS=9	26.93
CS=10	38.05

Sample	Weight Percent Insoluble Residue
CS=11	18.68
CS=12	26.80
CS=13	28.47
CS=14	22.82
CS=15	18.39
CS=16	12.14
CS=17	15.88
CT-1 CT-2 CT-3 CT-4 CT-5 CT-6 CT-7 CT-8 CT-9 CT-10 CT-11 CT-12 CT-13 CT-13 CT-14 CT-15 CT-16 CT-17 CT-18 CT-19	7.84 5.16 17.30 64.66 13.00 9.45 6.14 5.53 14.86 38.47 34.87 23.61 26.81 23.78 12.24 12.46 19.19 20.48 30.58
CU=1	12.26
CU=3	6.44
CU=13	3.83
CU=4	5.43
CU=9	34.74
CU=10	31.90
CU=11	13.94
CU=12	25.69
CU=5	10.28
CU=6	7.65
CU=8	10.75
CV-1	12.43
CV-2	9.66
CV-3	9.95
CV-4	5.77
CV-5	14.34
CV-6	38.40
CV-7	29.76
CV-8	27.37
CV-9	28.95
CV-10	9.88
CV-14	10.94

Sample	Weight Percent Insoluble Residue
CV-11 CV-12 CV-13	8.15 15.65 12.60
CW-2 CW-3 CW-4 CW-5 CW-6 CW-7 CW-8 CW-9 CW-10 CW-15 CW-12 CW-13 CW-11	3.09 2.73 3.40 37.73 32.14 27.20 33.26 26.27 10.23 9.53 16.87 23.08 18.53 18.85
CX-1 CX-2 CX-3 CX-1, CX-5 CX-6 CX-7 CX-8 CX-9 CX-10	6.93 4.84 7.98 35.39 24.49 9.59 9.26 14.24 12.25
CY-1 CY-2 CY-3 CY-4 CY-5 CY-6 CY-7 CY-8 CY-9 CY-10 CY-11 CY-12	9.53 13.41 16.51 11.17 38.59 38.91 22.49 19.86 21.83 11.02 11.71 7.76

# APPENDIX VII

This appendix contains data obtained from X-ray diffraction analyses. The stratigraphic positions from which the samples were collected are in Appendix I and (section E) in West, et al. (1972, p. 46-49).

Sample	Calcite Dolomite	Calcite Quartz	Mole % MgCO <sub>3</sub> in calcite	Mole % Ca in dolomite
CO=2 CO=14 CO=5 *CO=6 CO=7	0.0	11.5 3.5 32.3  24.0	2.25 1.35 1.05  2.55	53.10
CQ-2 CQ-3 CQ-4 CQ-5 CQ-6 CQ-7 CQ-8 CQ-10 CQ-11 CQ-12 CQ-13 CQ-15 CQ-16 CQ-17 CQ-18 CQ-18 CQ-19 CQ-19 CQ-19	9.6 3.6 16.4	9.0 24.0 11.5 24.0 99.0 11.5 2.6 5.2 6.3 11.5 24.0 32.3 0.2 0.1 0.6 24.0 5.7	2.25 2.55 1.62 0.895 1.95 1.95 2.55 2.562 2.80 1.62 0.40 1.95 1.95 0.60 1.95 1.95	56.85 56.15 56.77
CR-4 CR-6 CR-7 CR-9		24.0 2.6 11.5 49.0	1.95 1.95 1.62 1.95	
CS=1 CS=2 CS=3 CS=14 CS=5 CS=6 CS=7 CS=8 CS=9 CS=10 CS=11 CS=12 CS=13 CS=13 CS=15 CS=16 CS=17	4.8 3.6 16.5 2.1 15.0 6.8 38.5 10.4 44.0	11.5 32.3 11.5 1.2 13.3 6.7 13.5 2.5 6.1 3.8 3.7 3.7 8.8 6.1	1.95 1.95 1.00 1.62 2.50 1.95 1.62 1.95 1.62 2.50 1.35 1.95 1.95 1.95	55.55 56.48 56.48 56.17 56.48 56.17 57.05 54.32 56.18
CT-7 CT-11		24.0 2.6	1.05 1.95	

Sample	Calcite Dolomite	Calcite Quartz	Mole % MgCO <sub>3</sub> in calcite	Mole % Ca in dolomite
CT=13 CT=14 CT=15 CT=18	2.7 3.9	4.0 6.7 9.7 3.9	1.95 1.95 2.55 1.95	 56.48 53.43
OU-1 CU-6 CU-10 CU-11 CU-12 CU-13		11.5 11.5 4.3 15.7 6.1 49.0	1.95 1.62 1.62 1.62 1.95 1.35	
CV-3 CV-6 CV-8 CV-9 CV-12 CV-14		8.1 3.8 7.3 5.7 4.6 13.3	2.20 1.62 1.95 2.20 1.35	
CW-3 CW-6 CW-7 CW-8 CW-10 CW-11 CW-12 CW-13 CW-14	7.5 4.7 18.8  5.4  4.3 2.8	49.0 5.0 4.7 3.6 24.0 4.1 5.7 3.3 5.4	1.95 2.25 1.95 1.62 1.95 1.62 2.25 1.95 1.62 1.35	55.55 56.17 55.55 54.62 54.32 54.32
CX-2 CX-4 CX-5 CX-6 CX-8		19.0 3.2 4.3 24.0 7.3	2.55 1.95 1.00 1.35 1.95	
CY-1 CY-2 CY-3 CY-4 CY-5 CY-6 CY-7 CY-8 CY-9 CY-10 CY-11	45.0 9.6 4.9 2.0 2.2 2.0 5.7	11.5 7.3 4.3 11.3 2.6 2.0 3.8 3.6 3.6 3.2 6.0 5.7	2.20 1.62 2.20 1.35 1.80 2.20 1.35 0.70 0.70 1.05 1.62 2.55	55.25 56.18 54.90 55.87 56.17 54.95 55.87
E-6 E-7 E-8	6.0	2.2 2.9 5.7	1.35 2.25 1.95	55 <b>.</b> 55

\* no calcite in this sample

# FOSSIL ASSEMBLAGES, STRATIGRAPHY, AND DEPOSITIONAL ENVIRONMENTS OF THE CROUSE LIMESTONE (LOVER PERMIAN) IN NORTH CENTRAL KANSAS

by

## ROXIE LYNN VORAN

B. S., Kansas State University, 1975

AN ABSTRACT OF A MASTER'S THESIS

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Department of Geology

KANSAS STATE UNIVERSITY Manhattan, Kansas

#### ABSTRACT

A study of the fossil assemblages and stratigraphy of the Crouse Limestone in north central Kansas was undertaken to (1) test a previously proposed depositional model which suggested that the Crouse Limestone represents a regressive tidal flat sequence deposited in a lagoonal environment and (2) expand the area previously studied to include outcrops to the north where evidence suggested open marine conditions and the southwest where evidence favoring subaerial exposure was observed.

Careful stratigraphic correlation indicated that the rocks previously examined at sections in southern Marshall County were not the Crouse Limestone. Study of the proper stratigraphic interval at those outcrops revealed that they also support a regressive carbonate tidal flat environment of deposition.

Detailed study of a unique exposure in northern Geary County showed that an erosional surface and rock units above it--conglomerate, black shale, limestone, and gray mudstone--are consistent with the regressive interpretation. These units are interpreted as tidal channel and swamp deposits which are to be expected in a tidal flat environment.

Biotic data obtained from bedding surfaces, washed residues, and thin sections support the interpretation of the Crouse Limestone as a regressive carbonate tidal flat sequence. Fossil assemblages decrease upward in (1) density, (2) diversity, (3) mean size of individuals, and (4) size range of individuals. Petrologic data from the study of thin sections, insoluble residues, and X-ray diffraction also support the regressive interpretation. Dolomite and gypsum found in the middle shale and upper limestone occur in intertidal to supratidal deposits.

The area represents a protected lagoon and/or tidal flat. This tidal

flat and/or lagoon occupied the Irving Syncline, bounded on the west by the Abilene Anticline and on the east by the Nemaha Anticline.